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Chemical weathering of serpentinite rocks and implications for atmospheric CO2 carbonation

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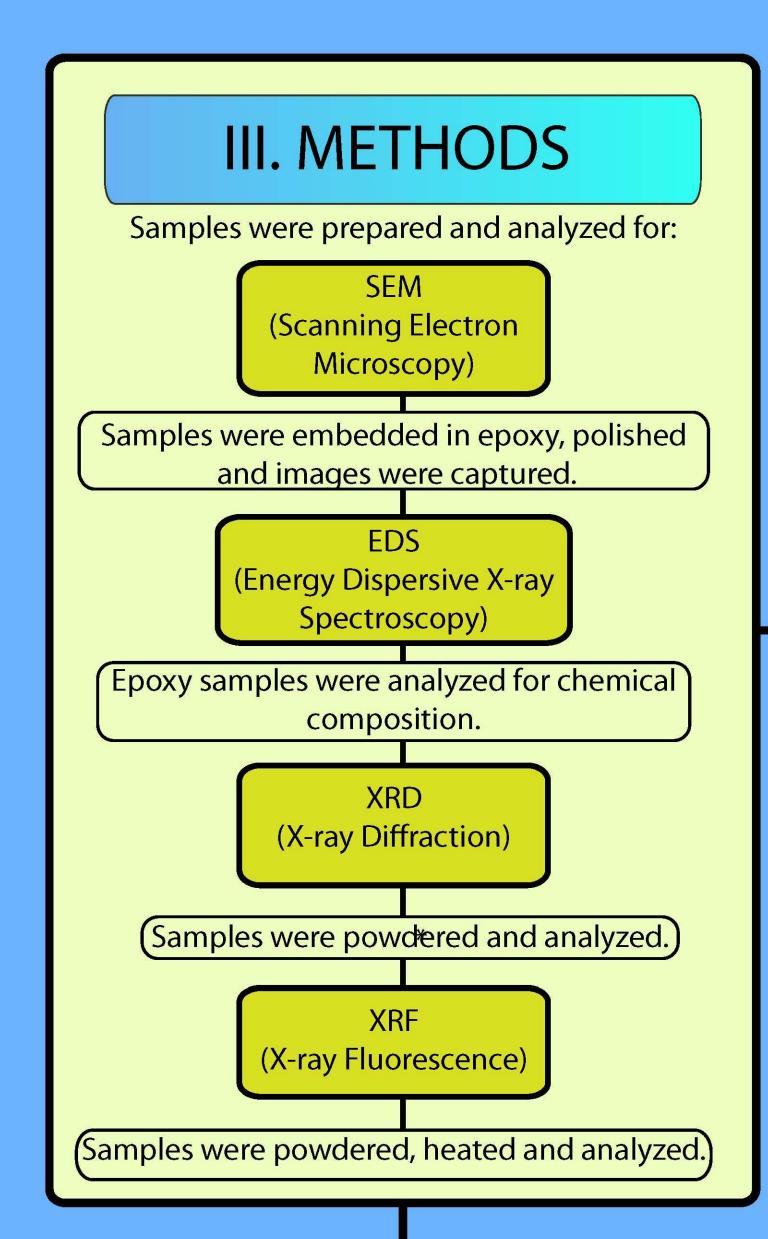
Chemical Weathering of Serpentinite Rocks and Implications for Atmospheric CO, Carbonation

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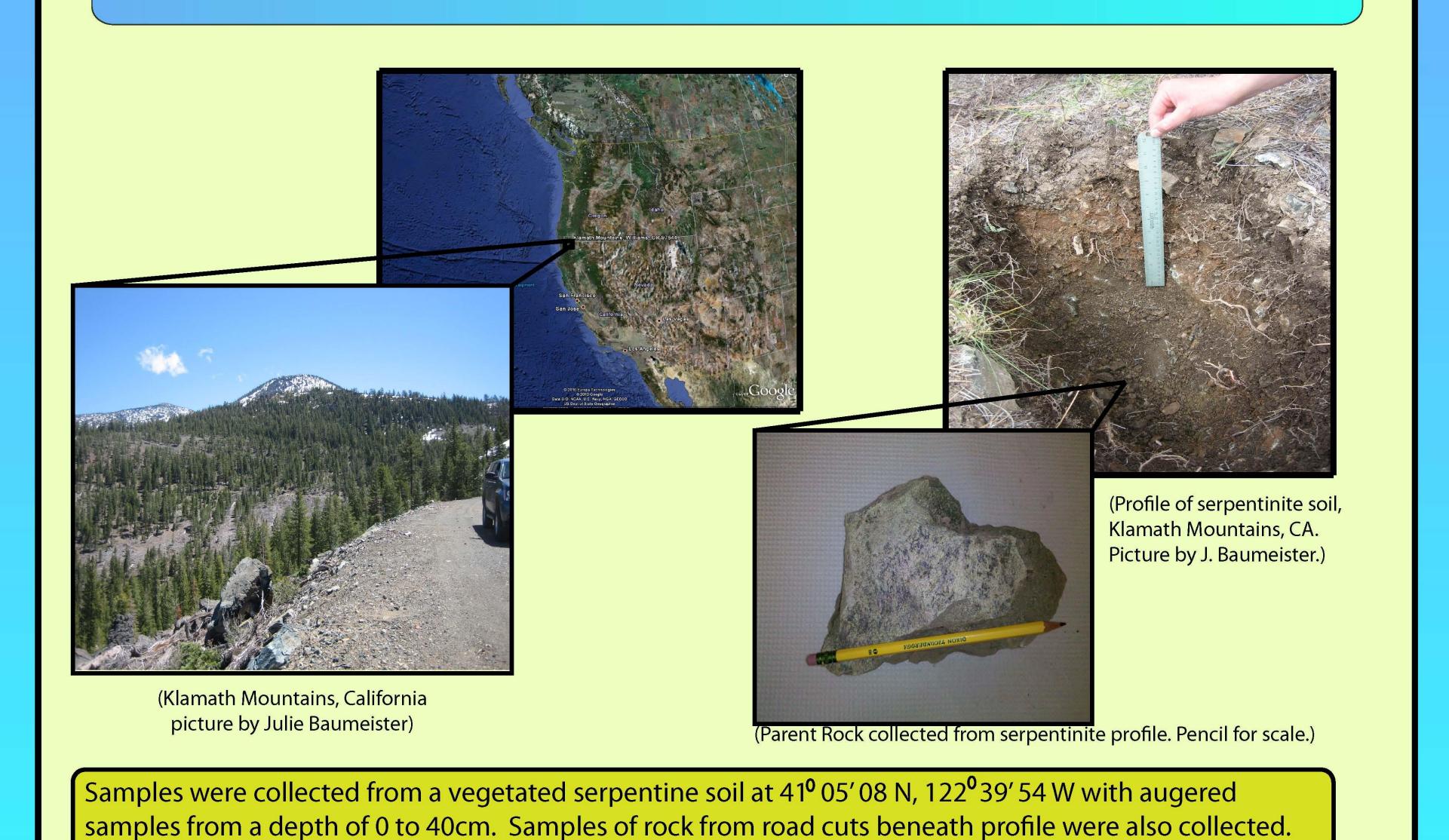


. INTRODUCTION

Concerns regarding global climate change have arisen as CO₂ emissions steadily increase at rates much larger than the consumption rates of CO₂ in nature (Kerrick et al., 1995). Chemical dissolution of weathering serpentinite soils, coincident with the formation of carbonate minerals, is a method by which carbon dioxide waste is sequestered naturally. In this study we will conduct mineralogical and chemical analyses to examine natural weathering of serpentinite soils. Quantifying dissolution rates of serpentine minerals could contribute to an understanding of weathering processes and the potential to accelerate them in industrial mineral carbonation. Although ultramafic rocks occur throughout the world, one locality, the Klamath Mountains in California, contain extensive serpentine deposits. A better understanding of serpentine mineral dissolution limits may have important implications for their use in mineral carbonation and their contributions to the carbon cycle.



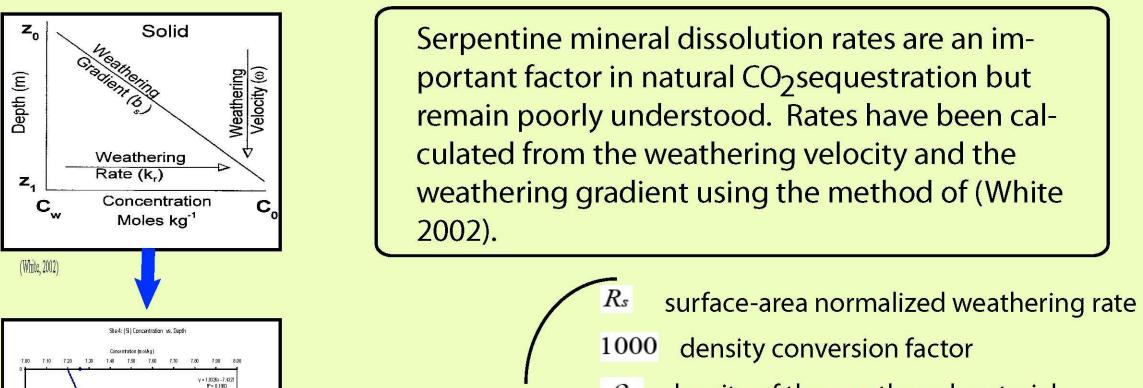
II. FIELD AREA AND SAMPLE COLLECTION



IV. RESULTS XRD SEM/EDS (X-ray Diffraction) SEM images of polished weathered Preliminary results suggest rock suggest the the presence of Lizardite, formation of Alsmectites, and Fe oxides in the bearing secondary parent material and weathminerals on the ered soils. (SEM image of serpentinite grain with weathering rind) (X-ray Fluoresence) Depth vs. Tau Plot Bulk chemistry of the soils were normalized to parent material to calculate the dimensionless mass transfer coefficient tau (1)*. normalized concentration of the immobile or mobile element $\tau = 0$, normalized concentraconcentration of the mobile ele-Cj, w ment in weathered material tion is equal to parent rock. $\tau_{i, j} = \frac{C_{j, w}}{C_{i, p}} \frac{C_{i, p}}{C_{i, p}} - 1$ $C_{j,\,p}$ concentration of an ment in parent material concentration of the mobile ele- τ < 0, normalized concentra $C_{j,p}$ $C_{i,w}$ tion is depleted. concentration of immobile ele- $C_{i,\,p}$ concentration of the ment in parent material $\tau > 0$, normalized concentration tion is enriched. concentration of immobile element in weathered material

V. DISCUSSION

Calculation of Serpentinite Dissolution Rate



Pw density of the weathered material reactive surface-area stoichiometric coefficient weathering advance rate weathering gradient

Kerrick, D. M., McKibben, M. A., Seward, T. M., and Caldiera, K., Convective hydrothermal CO2 emission from high heat flow regions: Chem. Geol., v. 121, p. 285, 1995.

VI. CONCLUSIONS

- Natural weathering rate: 5.76064 x 10⁻¹⁷ mol/m²s
- Lizardite appears to be altering to smectite.
- Chemical weathering appears to be occurring below the point of refusal.

VII. FUTURE WORK

We are continuing to analyze XRD results, constrain variables and characterize weathering in this environment to better interpret implications for atmospheric CO₂.

VIII. ACKNOWLEDGEMENTS

We would like to thank NSF EPSCoR (climate change) for funding and Dr. R. Metcalf, J. Cornell, R. Johnson, and Dr. S. Mulcahy for their assistance.

IX. REFERENCES

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