UNIVERSITY

[UNLV Retrospective Theses & Dissertations](https://digitalscholarship.unlv.edu/rtds)

1-13-2009

Relationships between the groundwaters of Ash Meadows, Death Valley, Pahranagat Valley and the Nevada Test Site based on statistical analysis and modeling of trace element data

Sara Michelle Cox University of Nevada, Las Vegas

Follow this and additional works at: [https://digitalscholarship.unlv.edu/rtds](https://digitalscholarship.unlv.edu/rtds?utm_source=digitalscholarship.unlv.edu%2Frtds%2F3148&utm_medium=PDF&utm_campaign=PDFCoverPages)

Repository Citation

Cox, Sara Michelle, "Relationships between the groundwaters of Ash Meadows, Death Valley, Pahranagat Valley and the Nevada Test Site based on statistical analysis and modeling of trace element data" (2009). UNLV Retrospective Theses & Dissertations. 3148. <http://dx.doi.org/10.25669/h6ok-5jcf>

This Thesis is protected by copyright and/or related rights. It has been brought to you by Digital Scholarship@UNLV with permission from the rights-holder(s). You are free to use this Thesis in any way that is permitted by the copyright and related rights legislation that applies to your use. For other uses you need to obtain permission from the rights-holder(s) directly, unless additional rights are indicated by a Creative Commons license in the record and/ or on the work itself.

This Thesis has been accepted for inclusion in UNLV Retrospective Theses & Dissertations by an authorized administrator of Digital Scholarship@UNLV. For more information, please contact [digitalscholarship@unlv.edu.](mailto:digitalscholarship@unlv.edu)

INFORMATION TO USERS

This manuscript has been reproduced from the microfilm master. UMI films the text directly from the original or copy submitted. Thus, some thesis and dissertation copies are in typewriter face, while others may be from any type of computer printer.

The quality of this reproduction is dependent upon the quality of the copy submitted. Broken or indistinct print, colored or poor quality illustrations and photographs, print bleedthrough, substandard margins, and improper alignment can adversely affect reproduction.

In the unlikely event that the author did not send UMI a complete manuscript and there are missing pages, these will be noted. Also, if unauthorized copyright material had to be removed, a note will indicate the deletion.

Oversize materials (e.g., maps, drawings, charts) are reproduced by sectioning the original, beginning at the upper left-hand comer and continuing from left to right in equal sections with small overlaps. Each original is also photographed in one exposure and is included in reduced form at the back of the book.

Photographs included in the original manuscript have been reproduced xerographically in this copy. Higher quality $6" \times 9"$ black and white photographic prints are available for any photographs or illustrations appearing in this copy for an additional charge. Contact UMI directly to order.

A Beil & Howell Infonnation Company 300 North Zeeb Road, Ann Arbor MI 48106-1346 USA **313/761-4700 800/521-0600**

Reproduced with permission of the copyright owner. Further reproduction prohibited without permission.

Relationships Between the Groundwaters of Ash Meadows, Death Valley, Pahranagat Valley and the Nevada Test Site Based on Statistical Analysis and Modeling of Trace Element Data

> to be submitted in partial satisfaction of the requirements for the degree of

> > Master of Science

in

Water Resources Management

by

Sara Michelle Cox

Department of Geoscience University of Nevada, Las Vegas August 1996

Reproduced with permission of the copyright owner. Further reproduction prohibited without permission.

Copyright 1997 by Cox, Sara Michelle

All rights reserved.

mVH Microform 1381011 Copyright 1996, by UM I Company. All rights reserved.

This microform edition is protected against unauthorized copying under Title 17, United States Code.

> **UMI 300 North Zeeb Road Ann Arbor, M I 48103**

The thesis of Sara M. Cox, written in pursuit of the Master of Science Degree in Water Resources Management, has been approved.

Advisory Committee:

Vernon Hodge

Chairperson, Vernon Hodge, PhD

Klaus Letzenbach
Examining Committee Member, Dr. Klaus Stetzenbach, PhD

Ken H. H. Konsesson
Examining Committee Member, Dr. Kevin Johannesson, PhD

Graduate Faculty Representative, Dr. Evangelo Yfantis, PhD

Dean of the Graduate College, Ronald W. Smith , PhD

University of Nevada, Las Vegas August 1996

ABSTRACT

Relationships between the groundwaters of the Nevada Test Site, Ash Meadows, Pahranagat, and Death Valley have been studied by many people over many years. Using hydrogeochemical data from these areas (sampled and analyzed by the Harry Reid Center for Environmental Studies) hydrologie subbasins were classified on the basis of trace element concentrations and compare these to previous classification systems. In addition, previously suggested flow relationships were examined on the basis of trace element concentrations, in particular those elements which are thought to behave conservatively in oxidizing environments. These efforts were made with the aid of statistical analyses such as principal component analysis and contouring within ARC/INFO (a geographic information system). The spring and well water chemistry data includes major ion and trace element chemistry and was obtained from the Harry Reid Center for Environmental Studies at UNLV. In most analyses only trace elements were examined. Results of principal component analysis yielded logical results that reflect differences in geology and location. Perched waters on the Nevada Test Site have the most unique chemistries. Waters from the Furnace Creek region of Death Valley, Ash Meadows, and Pahranagat Valley seem to have many similarities with respect to trace elements. Contour modeling reveals that at least three processes or parameters control the behaviour of trace elements within the study area.

Table of Contents

 $\frac{1}{2} \frac{1}{2} \left(\frac{1}{2} \left(\frac{1}{2} \frac{1$

List of Tables

 \overline{a}

List of Figures

Acknowledgements

I have been incredibly fortunate throughout my entire life to have found wonderful people to work and surround myself with. My family has always provided amazing amounts of support. I thank my parents and siblings first for their unconditional love and unfathomable patience. I would also like to thank all of my committee members, Dr. Vernon Hodge, Dr. Evangelo Yfantis, Dr. Kevin Johannesson, and Dr. Klaus Stetzenbach for all of their time and patience. The Harry Reid Center for Environmental Studies provided both a research assistantship and the hydrogeochemical data examined in this study. Without these two important components, this study would not have materialized. I appreciate the time and resources of Jason Johnson, Kevin Campbell, Kathy Lau, and Bob Taylor (Bureau of Land Management) which were spent working in Ash Meadows. I would also like to thank Mr. Bob Puterski and the Las Vegas Valley Water District for their help in the ARC/INFO work of this thesis. I am also thankful to Dave Kreamer for his efforts to keep me out of trouble. Lastly I thank my friends, both here and at home, for their ability to listen to me and make me laugh when I need to.

CHAPTER 1:

INTRODUCTION. SIGNIFICANCE, AND RELATED STUDIES

According to Claassen (1983), "ground water chemical data can help define ground water sources and pathways when combined with hydraulic data." Ground water within the study area is thought to react with tuffaceous rocks, carbonate rocks, and or carbonate valley fill (a mixed lithology of both carbonate and tuffaceous rocks). The major ion chemistry of the ground water is determined by the lithology of the rocks through which they flow and, cosequently, the trace elements are also likely influenced by this lithology (Claassen, 1983). Concentrations of elements in solution are further controlled by speciation and complexation as well as pH (Morel and Hering, 1993). In general, waters that discharge directly from the regional carbonate aquifer have certain chemical similarities, waters from local felsic volcanic rocks have separate chemical characteristics, and water flowing through different alluvial deposits should also share common chemistry because of the processes acting on the waters.

Trace element hydrochemistry is used in this study to analyze statistically how closely different springs are related and to identify differences between them. For this reason, trace element hydrogeochemistry may be used to classify hydrologie subbasins. In addition trace elements may possibly be used to support other efforts by Johannesson et al. (1996) which have hypothesized ground water mixing perhaps on both large and small scales. Ground waters of Ash Meadows, Death Valley, the Pahranagat

Valley and the Nevada Test Site (obtained from various springs and wells and named in Chapter 2) are examined in this project.

It is important to understand the pattern of ground water flow in southern Nevada and southeastern California. Much of the interest focused on the ground water of this region has been stimulated by Site Characterization studies of the high level nuclear waste site proposed for Yucca Mountain. Beyond these concerns, relationships throughout the arid basin and range province are important due to the dramatically increasing need for water. In the Las Vegas area alone the human population has grown from below 600,000 in 1985 to almost 1,000,000 in 1994 (Clark County Dept, of Comprehensive Planning). With this growth, demands on water resources have increased. As water table elevations fall, subsidence of the land surface and degradation of ground water quality may occur. Spring discharge is also affected by overtaxation of aquifers as illustrated by Figure 1 after Dudley and Larson (1976) (on page 3), posing possible problems for endangered species and other inhabitants in some subbasins.

The study by Dudley and Larson (1976) indicated that although not all springs in Ash Meadows were affected by pumping, some were. One result of Dudley and Larson's work indicated that pumping may divert flow from Davis Spring to Jackrabbit Spring. Although this example (represented in Figure 1) is at a relatively small scale, restricted to Ash Meadows, the ideas are the same at any scale.

In addition to the work of Claassen (1983), other authors have contributed chemistry data for the rocks and waters of southern Nevada as well as some knowledge of mechanisms involved in the alteration of ground water chemistry due to the host rock. These include: Winograd and Thordarson (1975), Raker and Jacobson, (1987), Schoff and Moore (1964), Johannesson et al. (1995), McKinley et al, (1991), and Broxton et al. (1989). Other scientists have contributed work over many years to further delineate and define ground water flow in southern Nevada. Some of these are: Dudley and Larson (1976), Hess (1992), Dettinger (1989), Burbey and Prudic (1991), Czarnecki and Waddell (1984), and Camera and Westenberg (1994).

Yelken (1996), and Farmer (1996) are concurrent research efforts at the University of Nevada, Las Vegas which is also related to ground water and rock chemisty in southern Nevada. Perfect (1994) has also defined subbasins in southern Nevada in work at the Colorado School of Mines. The work used a data set (compiled from numerous sources) including thousands of wells and springs, but seemed to classify on the basis of major elements.

Using hydrogeochemical data from these areas (sampled and analyzed by the Harry Reid Center for Environmental Studies) an attempt was made in this study to classify hydrologie subbasins on the basis of trace element concentrations and compare these to previous classification systems. In addition, previously suggested flow relationships are examined on the basis of trace element concentrations, in particular those elements which are thought to behave conservatively in oxidizing environments. These efforts were made with the aid of statistical analyses such as principal component analysis and contouring within ARC/INFO (a geographic information system). The spring and well water chemistry data includes major ion and trace element chemistry and was obtained from the Harry Reid Center for Environmental Studies at UNLV. In most analyses only trace elements were examined. As stated in the abstract the results of principal component analysis yielded logical results that reflect differences in geology and location. Perched waters on the Nevada Test Site have the most unique chemistries. Waters from the Furnace Creek region of Death Valley, Ash Meadows, and Pahranagat Valley seem to have many similarities with respect to trace elements. Contour modeling reveals that at least three processes or parameters control the behaviour of trace elements within the study area.

PLEASE NOTE

Page(s) not included with original material and unavailable from author or university. Filmed as received.

 $\mathcal{L}_{\mathcal{A}}$

 \sim .

UMI

Reproduced with permission of the copyright owner. Further reproduction prohibited without permission.

 $\sim 10^7$

ARC/INFO and is included in the Geographic Information Systems section of Chapter 5 to show precise locations of each sample location.

Figure 2: Schematic of Study Area Showing Locations of Clusters of Springs and Wells

Table 1: Locations of Springs and Wells (data from Harry Reid Center for Environmental Studies, McKinley et al, 1991, Camera and Westenburg 1994)

Geology of Southern Nevada

Lithology

During the Precambrian and Paleozoic more than 13,000 meters of marine sediments were deposited on the ancient marine depositional slope, which through time and geological deformation became exposed continental crust and part of the region studied in this thesis (Winograd and Thordarson, 1975, Wicander & Monroe, 1989). Both Cenozoic sedimentary and volcanic rocks are common throughout the study area (see Figure 3, page 9). Table 2 simplifies stratigraphie information from Winograd and Thordarson (1975). The table describes geology from the surface downward through successively older geology. Winograd and

Age	Stratigraphic Unit	Lithology	Thickness
Quaternary	valley fill	alluvial fan fluvial fanglomerate lakebed mudflow deposits	2000 feet
Tertiary	Basalt of Kiwi Mesa, Rhyolite of Shoshone Mountain, Basalt of Skull Mountain, Piapi Canyon Group, Wahmonie Formation, Salyer Formation, Indian Trail Formation, Calico Hills tuffs, Tuff of Crater Flat, Horse Springs Formation	tuffs & flows (non- welded to welded) sandstones, limestone, basalts	$2*104$ feet
Cretaceous to Pennsylvanian	Granitic stocks, Tippipah Limestone,	granodiorite, limestone	>3600 feet
Mississippian to Silurian	Eleana Formation, Devil's Gate Limestone, Nevada Formation	argillite, quartzite conglomerate, limestone, dolomite	$\sqrt{2^{4}10^{3}}$ feet
Ordovician	Ely Springs Dolamite Eureka Quartzite Pahrump Group	dolomite, quartzite, limestone, claystone	\equiv 3000 feet
Cambrian	Nopah Formation Dunderberg Shale Bonanza King Formation Carrera Formation Zabriskie Quartzite Wood Canyon Formation	dolomite, limestone, shale, siltstone, quartzite	\equiv 10000 feet
Precambrian	Stirling Quartz Latite Johnnie Formation	quartzite, siltstone, sandstone. limestone, dolomite	$=$ 5000 feet

Table 2: Stratigraphy: Nevada Test Site (Winograd and Thordarson, 1975)

Thordarson state that although the information is based on geology from the Nevada Test Site, it applies to a defined area which approximates the area included in this study. The lithology may prove quite important when attempting to draw conclusions from trace element data. As mentioned in Chapter 1, ground water may obtain a chemical signature from the rock it flows through (Claassen, 1983).

Figure 3: Generalized Geologic Map of Area (after Claassen, 1983)

Structure

The area of study has had a geologic history. The region was deformed during the late Mesozoic by folding, thrusting, and strike-slip faulting (Winograd and Thordarson, 1975). During the Miocene normal faulting (associated with volcanism) created the basin and range topography (Winograd and Thordarson, 1975). Figure 4 (after Stevens, et al 1991) describes in part the structural geology of the study area in very schematic

fashion. The figure is not to scale nor does it include every known structural feature in the study.

Geologic structure is important in many portions of the study area because of relationships that are sometimes observed between faults and springs. Ground water is thought to be discharged along a fault line in the Paleozoic carbonates below Ash Meadows, after which it percolates up through Quaternary deposits (Dudley and Larson, 1976). This scenario is visually explained in Figure 5 by a schematic cross section after Dudley & Larson, (1976).

Figure 5; Control of Ash Meadows Spring Line (after Dudley & Larson, 1976)

Hydrology of Southern Nevada

Aquifers

The principal aquifers in the region are the lower carbonate and valley fill aquifers. Winograd and Thordarson (1975) discuss other aquifers, including the bedded tuff aquifer, the lava flow aquifer, and welded tuff aquifers, which are not focused on in this study.

The saturated thickness of the Paleozoic lower carbonate aquifer varies from hundreds to thousands of feet (Winograd and Thordarson, 1975). This unit is fractured by faults as well as three sets of joints. Many caves are contained within this same unit, two of the largest ones being Devils Hole and Gypsum Cave (Winograd and Thordarson, 1975). The permeability of the carbonate unit within the area of the Nevada Test Site ranges from 0.00002 gpd/ft² to 0.1 gpd/ft², the mean being 0.01 gpd/ft²(Winograd and Thordarson, 1975). The carbonate aquifer discharges 1,300 gpm of the 1,430 gpm flowing from springs associated with the Spring Mountains (Winograd and Thordarson, 1975). The waters of the Spring Mountains serve to highlight the importance of the carbonate aquifer in southern

Nevada since such a high percent of the flow in the springs there is from the lower Paleozoic carbonate aquifer.

The valley fill aquifer is composed of alluvial-fan, fluvial, fanglomerate, lakebed, and mudflow deposits and overlies the Paleozoic carbonate rocks in many places (Winograd and Thordarson, 1975). Because less drilling is required, most wells in Las Vegas pump water from the valley fill aquifer, but it is important to note that these waters are related through intrabasin movement of ground water (Winograd and Thordarson, 1975). The valley fill aquifer layer happens to be more than 1000 ft thick in many locations but generally, the saturated thickness is only a small percent of the aquifer(Winograd and Thordarson, 1975).

Aquitards

The most important aquitards, in terms of aerial distribution are the lower clastic aquitard and the *tuff* aquitard (Winograd and Thordarson, 1975). The lower clastic unit is the lower boundary for ground water in the study area, and the tuff defines water in the Cenozoic aquifer versus the water in the Paleozoic aquifers(Winograd and Thordarson, 1975). The lower clastic aquitard has permeabilities much lower than those of the carbonate and valley fill aquifers ranging from 0.0000007 gpd/ft 2 to 0.0001 gpd/ft 2 with a mean of 0.00001 gpd/ft² (Winograd and Thordarson, 1975).

Springs

Most springs of the study area issue from bases of ridges of Paleozoic carbonate rocks, alluvium, lake beds, and tufa mounds (Winograd & Doty, 1980); although Scotty's Castle and Surprise springs emanate from volcanic tuffs as well as some springs in the vicinity of the Nevada Test Site. Spring

discharge is variable in terms of flow throughout the study area and occurs through a number of different processes. In the case of Travertine and Texas Springs, ground water must flow from the carbonate and through Quaternary gravels or Tertiary lacustrine deposits before surfacing while Nevares Spring flows from a travertine mound (Winograd and Thordarson, 1975). Jackrabbit Spring, Big Spring, Crystal Pool, Longstreet, Rogers, and Fairbanks are all solution or cavern type springs, created by dissolution of soluble rock (Hughes 1966). All springs and wells included in this study are listed in Table 1 as well as Table 3 on page 14. Table 3 includes the aquifer and source material (what the spring discharges from at the surface) if known, for the spring and well locations in the study. References are indicated by number and listed below the table.

Regional Flow

Within the study area water is thought to move through the earth by three different kinds of ground water movement: movement of perched water, intrabasin movement, and interbasin movement (Winograd and Thordarson, 1975). The hydraulic gradient (change in hydraulic head per unit distance) seems to increase from the Test Site (0.3 feet per mile) to the southwest (5.9 feet per mile) (Winograd and Thordarson, 1975). It may be important to note that over larger time scales the water table has not remained constant (as climates have changed) (Winograd & Doty, 1980). From the oversimplified water table contour map in Figure 6 on page 15, one can see that the flow in the study area is generally to the south and southwest (after Burbey and Prudic, 1991).

Table 3: Aguifer and Source Rock of Spring or Well

Region	Spring Name	Aquifer	Source
AM	Big Spring	carbonate ²	lake bed or travertine ²
AM	Bradford	carbonate ²	lake bed or travertine ²
AM	Cold Spring	carbonate ²	lake bed or travertine ²
AM	Crystal Pool	carbonate ²	lake bed or travertine ²
AM	Fairbanks	carbonate ²	lake bed or travertine ²
AM	Jackrabbit	carbonate ²	lake bed or travertine ²
AM	Kings Pool	carbonate ²	lake bed or travertine ²
AM	Longstreet	carbonate ²	lake bed or travertine ²
AM	Pt of Rocks NE	Aquifer	carbonate rock ²
AM	Pt of Rocks NW	carbonate ²	playa ⁷
AM	Rogers	carbonate ²	lake bed or travertine ²
AM	Scruggs	carbonate ²	playa ⁷
UDV	Upper Grapevine	X	QTal
UDV	Middle Grapevine	X	QTal
UDV	Lower Grapevine	X	QTal
UDV	Mesquite	carbonate ⁴	alluvium ⁴
MDV	Nevares	carbonate ²	travertine ⁴
LDV	Saratoga	carbonate ¹	carbonate ¹
UDV	Scotty's Castle	carbonate ³	Tvolcanics ³
UDV	Surprise	carbonate ³	Tvolcanics ³
MDV	Texas	carbonate ²	Qgravels, Tlacustrines ⁶
MDV	Travertine A	carbonate ²	Qgravels, Tiacustrines ⁶
MDV	Travertine B	carbonate ²	Qgravels, Tlacustrines ⁶
PAH	Ash	carbonate [/]	carbonate rocks ⁷
PAH	Crystal	carbonate [/]	carbonate rocks ⁷
PAH	Hiko	carbonate ⁷	carbonate rocks ⁷
NTS	Cane	t uff 1	t uff 1
NTS	Well J13	tuff ¹	NA (well)
NTS	Tippipah	tuff ¹	tuff ¹
NTS	Topopah	tuff ¹	tuff ¹
LDV	Soda Well	χ	NA (well)
NTS	Saga Well	χ	NA (well)
IS	Hardrock Well	χ	NA (well)
IS	Colliseum Well	X	NA (well)
NTS	Coffer Well	X	NA (well)
NTS	Cinderlite Well	X	NA (well)
NTS	Army Well	carbonate ³	NA (well)
NTS	Airport Well	valley fill ⁷	NA (well)
NTS	J12	volcanic ³	NA (well)
NTS	Lathrop Well	valley fill ⁴	NA (well)

Perfect, Faunt, & Steinkampf, 1994¹, Winograd and Thordarson, 1975², Camera & **Weslenburg, 1992^, Kreamer et al, 1996^, Winograd,1971®, Johannesson et al, 1996^**

Reproduced with permission of the copyright owner. Further reproduction prohibited without permission.

Winograd and Thordarson (1975) give several reasons which support the hypothesis that waters throughout the study area are hydraulically related. The first is that the saturated zone of the lower carbonate aquifer is probably at least 4,000 feet thick and laterally extensive in the study area; therefore, movement of ground water through the aquifer from basin to basin is likely (Winograd and Thordarson, 1975). In addition, the chemistries (major ion) of the lower carbonate aquifer beneath Yucca and Frenchman Flats are similar to that of the springs in Ash Meadows (Winograd and Thordarson, 1975). Lastly, a regional potentiometric map indicates that ground water at the test site contributes flow to both Ash Meadows, Death Valley, and the Oasis Valley (Winograd and Thordarson, 1975). According to a study by Hunt, et al (1966), the waters discharging in Death Valley and Ash Meadows are also related to the waters of Pahranagat Valley, with 35% of the flow in Ash Meadows fed by this source (Winograd and Thordarson, 1975).

Figure 6: Water-Table Contour Map of southern Nevada (after Burbey & Prudic, 1991)

Reproduced with permission of the copyright owner. Further reproduction prohibited without permission.

Rainfall/Evaporation

As the designation of "desert" suggests, the area has a high evaporation/precipitation ratio. The average annual rainfall over the region is three to six inches while potential evaporation ranges from sixty to eighty-two inches per year(Winograd and Thordarson, 1975). The driving force for these high potential evaporation rates is high temperatures. The average high temperatures in the area range from 40.5°C in Las Vegas, to 24.5°C in central Yucca Flat, to 49°C in Death Valley (Winograd and Thordarson, 1975). In contrast, higher elevations receive more rainfall (Linsley et al, 1992). In fact, the Spring Mountains and Sheep Range generally receive the greatest amounts of rainfall(Winograd and Thordarson, 1975).

Surface Drainage

Some precipitation in the study area is delivered as runoff to the Colorado River (Winograd and Thordarson, 1975). Both the Las Vegas Valley and the Pahranagat Valley contribute runoff to the Colorado (Winograd and Thordarson, 1975). Jackass Flats and the Amaragosa Desert are up gradient topographically from Death Valley and are drained by the Amaragosa River. Water from most of the remaining valleys in the vicinity of the Nevada Test Site (NTS) flows to playas (Winograd and Thordarson, 1975). The western Amaragosa Valley is recharged to a large extent by surface runoff and channel flow (Claassen, 1983).

Chemistry in Ground water and in Rock

Water within the study area will have reacted within tuffaceous rocks, carbonate rocks, or carbonate valley fill- a mixed lithology of both carbonate and tuffaceous rocks. The major ion chemistry of the ground water is

determined by the rock types of the subsurface and surface (Claassen, 1983). For this reason some information is given in Table 4 and 4 on the chemistry of carbonate and igneous rocks.

Trends of Trace Element Concentrations in Rock

Carbonates- Limestones and Dolomites comprise approximately 1/6 of the crustal sedimentary mass (Veizer, 1983). The average trace element chemical composition of carbonates is presented in the Table 4 on page 18. An "x" indicates that the value is on the order of magnitude indicated. In general, the chemistry of sedimentary carbonates is determined by provenance or source. Trace elements become a part of the crystal lattice in minerals by substitution for Ca^{2+} , interstitial substitution, addition of trace elements at defect sites, and adsorption induced by ionic charge (Veizer, 1983).

Igneous Rocks- The chemical composition for igneous rocks vary widely; however, there are to be relationships for trace elements in different igneous rock types. Mafic rocks typically have more chromium and cobalt, whereas felsic rocks tend to be richer in barium, rubidium, lead and beryllium (Salomons and Forstner, 1984). Table 5 on page 18 gives average trace element values for the Timber Mountain-Oasis Valley Caldera Complex, Nevada (Broxton, et al, 1989). For individual values for concentrations of each unit refer to the original cited reference.

From the works of Vezier (1983) and Broxton et al. (1989) it can be seen that carbonates on average have greater amounts of strontium than some of the NTS volcanics. Both rock types have similar concentrations of vanadium and antimony. The NTS volcanics have higher concentrations of arsenic rubidium, zirconium, and barium.

Reproduced with permission of the copyright owner. Further reproduction prohibited without permission.

Element	$\frac{1}{2}$ Carbonate	Deep Sea	, Element	Carbonate	Deep Sea
	(ppm)	Carb (ppm)		(ppm)	Carb (ppm)
Li	5	5	Ge	0.2	0.2
B	20	55	Cd	0.035	0.0x
$\overline{\mathsf{F}}$	330	540	Mo	0.4	$\overline{3}$
Na	400	2000	Sb	0.2	0.15
Mg	47000	4000		1.2	0.05
AI	4200	20000	Cs	0.x	0.4
$\overline{\mathsf{Si}}$	24000	32000	Ba	10	190
\overline{P}	400	350	La	\boldsymbol{x}	10
\overline{s}	12000	13000	Ce	11.5	35
CI	150	21000	Pr	1.1	3.3
Κ	27000	2900	Nd	4.7	14
Ca	302300	312400	Sm	1.3	3.8
Ti	400	770	Gd	1.3	3.8
$\bar{\mathbf{V}}$	20	20	Tb	0.2	0.6
Cr	11	11	Dy	0.9	2.7
Mn	1100	1000	Ho	0.3	0.8
Fe	3800	9000	Tm	0.04	0.1
Ga	4	13	Yb	0.5	1.5
As	1	$\overline{\mathbf{1}}$	Lu	0.2	0.5
Se	0.08	0.17	Hf	0.3	0.41
Br	6.2	70	Ta	0.0x	0.0x
Rb	$\mathbf{3}$	10	W	0.6	0.x
Sr	610	2000	Hg	0.04	0.0x
Y	30	42	Pb	$\overline{9}$	9
Zr	19	20	Th	1.7	\mathbf{x}
Nb	0.3	4.6	U	2.2	0.x

Table 4: Chemistry of Carbonates (Veizer, 1983)

Tables : Trace Element Chemistry: Timber Mountain-Oasis Valley Caldera Complex (concentrations are in parts per million for trace elements and

 \sim \sim \sim

Chemistry in Ground water of Southern Nevada

In general, waters that discharge directly from the carbonate aquifer have certain chemicalsimilarities, waters from volcanic rocks have separate chemical characteristics, and water flowing through different alluvial units probably have some common chemistry. Other parameters such as the regional flow pattern or the effects of man may then overprint these chemistries. Winograd and Thordarson (1975) expanded a classification scheme designed by Schoff and Moore (1964) and list five hydrostratigraphic facies of ground water in and around the NTS in southern Nevada (Table 6).

Class	Characteristic Source	Examples		
Calcium magnesium bicarbonate facies	lower carbonate aquifer or valley fill aquifer where carbonate rich	Spring Mountains, Pahranagat Valley		
Sodium potassium bicarbonate	tuff, myolite, valley fill where rich in volcanics	Yucca Flat, Frenchman Flat, Jackass Flats, west and northwest of NTS		
Calcium magnesium sodiumm bicarbonate	lower carbonate aquifer	Ash Meadows, eastern NTS		
playa	where ground water is removed by evapotranspiration, rather than by fluid flow discharge			
Sodium sulfate bicarbonate	Furnace Creek Wash and west-central Amargosa Desert			

Table 6: Ground water Classification Scheme by Major Ions
(Minograd&Thordarson, 1975)

Some water within the study area may be significantly impacted by contact with volcanic tuffs due to dissolution of metastable glass as it alters to other minerals (where tuffs are present). A study by White, Claassen, and Benson (1980) indicates that in the Rainier Mesa area deeper water, both interstitial and in fractures, is richer in sodium and depleted of calcium and magnesium with respect to shallower waters. For the most part, these increasingly sodium rich waters move through the tuff through the porosity and are changed chemically through the processes of dissolution.

precipitation, sorption, and ion exchange (White, et al, 1980). It is possible that these same processes are important for all ground water in the study area. Table 7 shows major ion chemistry for ground water in Rainier Mesa (White et al, 1980). Concentrations are in millimoles per liter and are averages of numerous values (all from units of the Rainier Mesa) in the original cited references.

Table 7: Chemistry of Tuff Waters of Rainier Mesa (concentrations are in millimoles/Liter) (White et al, 1980)

r. I	Bicarbonate		
0.12	Sulfate	0.15	
	Chloride	0.24	
0.06	Flouride	0.01	
	0.21		

CHAPTER 3:

SPRING AND WELL HYDROGEOCHEMISTRY DATA DOCUMENTATION

Reconnaissance and Sampling for Trace Elements

On reconnaissance sampling excursions different physical and chemical parameters were measured. Latitude and longitude were determined using a Panasonic brand global positioning satellite system. Other measurements taken in the field included: pH, TDS, conductance, and temperature. Four liters of spring water were collected in acid washed polyethylene bottles, after they were filtered through a $0.45 \mu m$ polysulfane filter. Samples were analyzed within one week for trace element chemistry determinations and within 2 to 4 days for anion results. The above sampling procedure is described in Stetzenbach et al. (1994). Waters were sampled for trace element concentrations on the dates shown in Table 8.

Ash Meadows springs	July 1992, January, May, October 1993, March 1994
Death Valley springs	June 1992, March, July, November 1993, March 1994
Pahranagat Valley springs	September 1993
Nevada Test Site springs	December 1994, January, February 1995
wells	May, June 1994

Table 8: Sampling Dates (Harry Reid Center for Environmental Studies)

Reproduced with permission of the copyright owner. Further reproduction prohibited without permission.

Analytical Procedures

Major anions were analyzed by ion chromatography and major cations by atomic absorption spectrophotometry (Harry Reid Center for Environmental Studies). Rare earth elements (REEs) and trace elements were analyzed by inductively coupled plasma- mass spectrometry(Harry Reid Center for Environmental Studies). Concentrations as low as parts per trillion were determined by ICP-MS (Harry Reid Center for Environmental Studies). The ICP-MS machine used is the Perkin Elmer Sciex Elan 5000 ICP- MS with an active film multiplier detector. The samples entered the ICP- MS via an ultrasonic nebulizer. Further specifics may be obtained from the Harry Reid Center for Environmental Studies.

CHAPTER 4:

METHODS OF INVESTIGATION & PROBLEM SOLUTION

Data Selection

Although approximately 55 elements were analyzed to produce the data set obtained by HRC, only a portion of this data was used. One reason that only a portion of the data set is used is that trace elements were to be the focus of this study. A second reason is that some element concentrations were found to be nondetect values more often than not. Others were not measured in all locations. Lastly it is widely held that certain elements are most difficult to obtain concentration values for because of problems inherent to sampling. For example iron, copper, lead, zinc, and cadmium concentrations have historically been difficult to measure (Windom et al, 1991, Runnells et al, 1992). For the majority of the research presented in this thesis, the following chemical element concentrations were utilized: selenium, vanadium, arsenic, tungsten uranium, molybdenum, rhenium, manganese nickel, gallium, rubidium, cobalt, strontium, cadmium, cesium, barium, thallium, tin, antimony, titanium, germanium, tantalum, lithium, chromium, and zirconium. Measured elemental concentrations used not only meet internal quality control standards of the Harry Reid Center for Environmental Studies but also were selected on the basis of the number of nondetects. Generally if an element was analyzed as a nondetect or not measured more often than not then the element was generally excluded from

23

Reproduced with permission of the copyright owner. Further reproduction prohibited without permission.
analysis. Some elements which were measured many times as non detects were included to represent columns of the periodic table with minimal weighting of each column in the periodic table. Only the three lightest REEs consistently have discrete measured values. Since the heavier REEs could not be represented in analysis, no REEs were considered. In Chapter Five, exploratory analysis shows that exemption of REEs does not hide differences between waters.

As mentioned previously the author obtained data from the Harry Reid Center for this study. The data are trace element concentrations from springs in Ash Meadows National Wildlife Refuge, Death Valley National Park, Pahranagat Valley and wells and springs on the Nevada Test Site as well as wells in Shadow and Ivanpah Valley. Because the data has already been collected, the thesis research did not include development of any experimental model. It is likely that a better understanding of the regional groundwater system could be obtained if more sampling and analysis were done at appropriate locations.

Principal Component Analysis

The author analyzed the data statistically. The statistical analysis included Principal Component Analysis (PCA) of the entire data set, as well as other analyses (such as correlation plots, dendograms, and icicle plots). Different scales and kinds of analysis were performed and are described in this paragraph. In the principal component analysis the elemental concentrations are variables and the cases are the spring or well water names. In the most inclusive analyses all springs and wells were included and the largest number of variables were considered. The most exclusive analyses considered a limited number of both cases (springs) and variables

(chemical elements). Different kinds of analyses were conducted so that smaller geographical scales could be focused on, and the importance of waters which were thought to be perched could be minimized (since perched waters are probably less important to regional flow). These methods were used in an attempt to gain a better understanding of the physical parameters of the study area. Relationships between springs and differences between springs were noted if there there were any trends of changing trace element concentrations with changing season by making simple time vs. concentration plots for different springs. The results were then discussed.

Factor analyses like principal component analysis (PCA) has been used for numerous geological problems in the past (Joreskog, et al., 1976). Examples are: (1) using trace elements in sediments to determine the sediment origins; (2) using chemistry and structure of ore bodies to find likely sites for additional ore bodies; (3) and using the relationships between sediments samples and measurements of organism populations to determine preferences of organisms for certain kinds of sediments (Joreskog, et ai., 1976). Okuda,et al (1995) used principal component analysis to classify pyroclastics based on chemical composition and make inferences about age and spatial correlations. Nash, et al (1993) clustered soils using principal component analysis into groups with two factors which explained from 53- 60% of the variance within the sample.

PCA is a statistical method which is capable of reducing an unmanageable number of variables into a smaller number of composite variables called factors. The method is based on linear algebra and the use of matrix manipulation. Principal component analysis is a form of multivariate statistical theory which makes the assumption of normal distributions; however, principal component analysis is considered robust enough

Reproduced with permission of the copyright owner. Further reproduction prohibited without permission.

mathematically to be appropriate for this study (Yfantis, 1996). A good factor analysis solution will weight variables fairly, explain observations with a minimum number of factors, be meaningful, be simple, and will be interpretable (Norusis, 1994). Using commands to extract factors, PCA determines orthogonal factors based on an uncorrelated matrix and create linear combinations of the variables (Norusis, 1994). For example, the first principal component might be described by the linear equation " $Y_1 = a_1 X_1 +$... + $a_{p1}X_{p}$ " if the covariance matrix is used (Morrison, 1967). In this case the first term of the equation represents the product of the vector and the scalar values of the first variable's contribution to the first principal component. The first principal component explains the largest amount of variance in the data set; additional principal components account for smaller amounts of variance. One should note that for this research, correlation matrices were used which likely define relationships in a similar fashion but require a much more complex equation to express principal components; however, the relationship between the vector portion of the first term is easy to define (s=variance, $z = new vector$ term for principal component created using correlation matrix).

 z_{ii} = (xij - xmean_i) / s_i (Morrison, 1967)

Both Statistical Products and Service Solutions (SPSS) (Chicago, Illinois, 1994) and Statistica (Statsoft, 1993) were used for the Principal Component Analysis (since both software packages are based on the same fundamental principles, the results should be the same). Principal Component Analysis has a general procedure outlined below (based on procedure as user progresses through analysis as performed by SPSS (Norusis, 1994)). First the operator must decide howto deal with problem of missing values. Missing values can be dealt with by exclusion of cases.

variables, or replacement. Next the user determines the number of factors necessary to explain a certain amount of variance. This can be done by examining the scree plot. A scree plot is a two dimensional graph relating the percent of variance explained (or sometimes indicated by an eigenvalue) to the number of factors that explains the variance in the principal component analysis. The option of rotating data may be considered. The software can then generate a correlation matrix and provide the user with factor scores. These factor scores are measures of how the data are now described by the principal components. For example. Case A may originally have been described by four variables with values of 1, 2, 3, and 4. After principal component analysis. Case A may be described with two principal components with values of 2 and 5. These factor scores can then be graphed in scatterplots with orthogonal axes which correspond to the different principal components.

The possibility exists that as many principle components as variables can be generated; however in this study 3 was the largest number of principle components needed to explain \equiv 50% of the variance in the sample (39 cases, 23 variables for entire working data set). For ease of display only the first two principal components were illustrated in two dimensional scatterplots. In addition, the factor analysis was not rotated and the factor scores were computed by regression. For the majority of analyses, the mean value of each element was used to fill empty cells in the original data sets (empty cells due to non detects). These means are calculated from the chemical concentrations of the same element in all other cases.

Agglomerative Cluster Analysis

Another way of analyzing data is to perform a cluster analysis, in SPSS the squared Euclidean distance (sum of the squared differences of each elem over all of the variables) coefficient matrix is the first element of forming clusters. The Euclidean distance coefficient matrix is determined by calculating differences between corresponding variables between each combination of cases, such that the differences between each variable are weighted equally (Norusis, 1994). Plots are then generated to illustrate relative similarities between waters. Some software normalizes Euclidean Distances to a certain value, but some do not. Both Dendograms and Icicle Plots can be generated after the Euclidean matrix is created. Both dendograms and icicle plots are useful, but do express some common information. Icicle plots are read from the bottom up and best exhibit (spring to spring) which cases are most similar.

Histograms

Histograms were constructed by using Microsoft Excel 4.0 and SPSS Exploratory Analysis. In Excel the histograms were assigned bin sizes after examination of the ranges in the chemical concentration data. In this case bins are divisions in concentration axis of concentration histograms. Excel was used to make sets of histograms with common bin sizes so that different subbasins may be compared. The histograms created in SPSS Exploratory are in "leaf and stem" format so that readers can verify frequencies without any effect of bin size. In this study frequency refers to the number of springs or wells having concentrations that have a given concentration. The leaf and stem format allows for reconstruction of the data set from which the histogram was built. If simple histograms alone are created and analyzed.

there may be problems related to bin size. Software may choose bin sizes for the scientist. The default selection may not be appropriate for the study and may hide certain things to a small extent. Figure 7 is an example which illustrates how this problem might affect representations of data from this study.

Figure 7: Effects of Default Bin Size on Histograms a) carefully selected bin size

Geographic Information Systems

Geographic Information Systems (GIS) is a kind of spatial database which can be used for modeling and analyzing spatial data in a variety of ways. Environmental Systems Research Institute (ESRI) has developed a software package for use with personal computers and UNIX systems entitled *ARC/INFO.* The ARC part of the program is responsible for locating the features, for example, springs; the INFO part handles the information and

descriptions of each spatial feature (ESRI, 1990). GIS is described in simple terms as an a database with two components. Part is an information database containing descriptive information, and the other part is responsible for locations in X-Y space of data points which corresponds to the descriptive information in the information part of the database.

GIS is superior to other forms of spatial analysis because it allows spatial operations on the data set. Refer to Figure 8 on page 31 when reading the following explanation. GIS can answer many important questions. The system can define what features or conditions exist at certain locations prompted by queries regarding a certain location. The system can identify changes over time as well as spatial patterns. In addition ARC/INFO is able to perform these previously mentioned tasks after certain elements of the spatial database are changed (ESRI, 1990). The ability to answer queries, and perform calculations, means that GIS is much more than just a map making tool. It may be used to present chemical data in such a form as to make certain relationships more clear and even to manipulate the data. From these capabilities, more information will be found which will either support or counter previous efforts to define ground water flow through southern Nevada.

Figure 8: Schematic of a Geographic Information System after ESRI, 1990

ARC/INFO can be a powerful tool for analysis of spatial data. For a variety of reasons only a minute portion of the software's capabilities were utilized in this study for many reasons. Foremost among these reasons is the disparate distribution of data points in the study area. For this reason the data was contoured using the kriging method available in ARC/INFO. Kriging is based on the idea that when one considers the value of a data point to determine nearby values, one should also consider the values of neighboring points. Kriging is based on the idea of a regionalized variable with three components; drift (trend), spatially correlated randomness, and noise (Clarke, 1990).

The first step in the spatial analysis was to prepare the database. INFO was used to create a template data file (chemcov) containing the concentrations of: arsenic, antimony, cesium, cobalt, selenium, vanadium, uranium, molybdenum, rhenium, rubidium, nickel, thallium, gallium, tungsten, and germanium in groundwaters from springs and wells in southern Nevada. The table was then filled from an ascii file using the ADD FROM command.

Using the ARC command PROJECT the locations of all data points were converted to a standard of Nevada State Plane (Fipszone 2702, GRS80, Datum NAD83) from the latitudes and longitudes obtained from the Harry Reid Center. The ARC command GENERATE was used to create the point coverage <location>. The location.pat file was joined to the file chemcov to form the database all work in this section is based on.

The next step was to create a map with the truest representation of space possible. This map is included in the GIS section of Chapter 5. Springs and wells are labeled to correlate with a listing at the bottom of the map. In areas such as Ash Meadows and the Grapevine springs locale, the numbers may overprint one another due to their proximity. While this does

make identification of individual springs in these areas difficult, coordinates for each data point are included in Chapter 2.

ARC/INFO has several tools for modeling surfaces and contouring spatial parameters; TIN, GRID, and kriging. TIN and kriging are examined in this study. The macros used to generate the TINS and kriged contours (Appendix E). The macros are written in ARC Macro Language and are all designed to be run from the ARC prompt.

Typing the command Kriging at the ARC prompt initiates dialog to establish the conditions of the model. A contour interval is specified (different for different elements) and the name of the contour coverage and variance coverage are defined. In each analysis, the entire study area was kriged. For this analysis the spherical distribution (default) was selected for kriging technique because neither the gaussian nor universal kriging techniques is necessarily appropriate for these models. The contour coverage created will preserve the integrity of the original data set as modeling with TIN does.

ARC/INFO makes analysis of spatial variance relatively simple by using the kriging command. Variance in this study is a measure of how each water sample concentration for each element varies from the mean of the mean for that element for all waters in the study. Kriging in this software creates a coverage of variance which can then also be contoured which in essence provides a map of variance of the data. Areas of higher variance are areas where more data points are needed to improve the quality of the chemical concentration contour map. It is probably obvious that many more data points would be useful in this analysis just from looking at the location map.

32

Reproduced with permission of the copyright owner. Further reproduction prohibited without permission.

It is important to note that contouring any parameter is to some degree affected by the technician's individual choices. All of the results obtained are based on contour intervals, techniques, and base contours defined by the operator; however, in ARC/INFO one should always obtain the same result, provided the parameters defined in the kriging dialog remain constant.

Chapter 5:

EXPLORATORY EFFORTS AND DISCUSSION

ELEMENTAL CONCENTRATION PLOTS

The elemental concentration plots are simple X-Y and 3D bar graphs which illustrate concentrations of elements in different springs and wells. They are particularly useful to provide a visual foundations for some of the results of principal component analysis and hierarchical analysis of spring and well waters. Simple bar charts can also be used to illustrate change over time. Examples illustrating these statements follow.

Simple plots of elemental concentrations of the different water samples are useful to illustrate the actual concentrations of the elements which seem to provide the most variance between subbasins (as indicated in following principal component analysis section). Figures 9-13 on pages 36-38 compare springs and wells on the basis of several different element concentrations: uranium, rhenium, molybdenum, antimony, and cesium. From inspection of these simple plots of concentrations, one can observe that there are distinct groups of waters which consistently are different from the rest. Death Valley generally has lower concentrations of rhenium. Usually the perched springs of the NTS have either higher or lower concentrations of elements compared to other waters.

34

Reproduced with permission of the copyright owner. Further reproduction prohibited without permission.

 \mathcal{L}

ပ္ပ
ဟ

Reproduced with permission of the copyright owner. Further reproduction prohibited without permission.

³⁶

As mentioned in Chapter 3 and Chapter 4, the springs of Death Valley and Ash Meadows were sampled six times by the Harry Reid Center (HRC). During the study attempts were made to identify any marked changes over time in chemical concentrations. Figure 24 was included on page 44 in this section for illustration of lithium concentrations in select Ash Meadow springs. Although there are some differences, these differences are more likely to represent improved techniques in sampling and analysis, than changes that occur over time (K. J. Stetzenbach, 1996 pers comm, V. F. Hodge, 1995 pers comm). For springs that were sampled multiple times, trace element concentration values were taken from the fifth sampling date, unless the spring was not sampled on that date. In the case of missing values in the fifth sampling date of springs sampled multiple times, the fourth sampling date was used.

Figure 14: Lithium Concentrations in Ash Meadows and Death Valley (ppb)

HISTOGRAMS

It is interesting to note how the histograms of some elements differ between subbasins. Figure 15 on page 38 shows the distribution of molybdenum in carbonate waters from Ash Meadows springs (in ppb) versus the distribution of molybdenum in carbonate waters from Death Valley springs (in ppb). This means that although there may be normal distributions when one examines the entire data set and when one examines an individual subbasin, there may also be a great enough difference between the mean of each chemical element in each different subbasin to create distributions with more than one mode. The histograms of Figure 15 make use of all of the first five sampling dates' data.

From an analysis of the entire data set simple histograms were created. Medians, maximums, and minimums were also included with stem and leaf plots in Appendix A. The histograms were used to determine contour intervals in Chapter 6 where contour maps of different element concentrations were made. The histograms are also useful for visual reference to gain a simple understanding of the abundance of elements in the ground water throughout the study area. Additional histograms are in Appendix A.

Reproduced with permission of the copyright owner. Further reproduction prohibited without permission.

PRINCIPAL COMPONENT ANALYSIS

In science the output of a model or program is dependent on input. It is also true in science that specific parameters are measured, not neccessarily every possible one. Sometimes items are measured and upon analysis it is obvious that there was some error in the measurement process, as is often the case with certain trace metals. In any case one must always be aware of what information an analysis is based on, and understand as well as possible the effect of each piece of data on the result. This section is included to illustrate that although there are differences resulting from certain changes in the variables of the data set analyzed, certain relationships are almost always found in this study area.

Effects of Different Variable Sets on Analvsis: Entire Data Set vs Malor Elements and Anions vs Trace Elements (no REEs)

The first principal component analysis is based on trace element data from springs and wells in addition to majors and anions to create Figure 22. Although it might be hypothesized that standard deviations of element concentrations measured at part per trillion levels might be higher and in effect weight the results of principal component analysis, this is not necessarily the case. If one examines water sample collected in Ash Meadows and Death Valley this is evident. The variables included in this analysis are indicated in materials within Appendix D. This analysis will then be compared with other variable sets in this section to get an idea of what effect major elements and anions, and the trace elements had in the PCA of Ash Meadows and Death Valley data. The graphical results of using different variable sets follow in Figures 22-24 on the following pages. Each graph is followed by discussion. For each scatterplot of waters as described by principal components, the author has included a variable list, eigenvalues, and factor loading matrix in Appendix

Reproduced with permission of the copyright owner. Further reproduction prohibited without permission.

D. These materials are not needed to visualize or discuss results but do allow for inspection of information in the analysis.

Figure 22: Scatterplot of Ash Meadows & Death Valley (entire variable set)

In Figure 22 outliers include Cold, Saratoga, Surprise, and Scotty's Castle. The Furnace Creek springs cluster togetherand the majority of Ash Meadows cluster closely together and near the Furnace Creek or middle Death Valley springs. The Grapevine springs and Mesquite Spring plot near each other.

It is interesting to note that in the second plot in this section. Figure 23 (p43), generated using only major elements and anions, there is the greatest amount of clustering. In other words, the outliers. Cold, Saratoga, Scotty's, and Surprise are more isolated from the remaining springs (Ash Meadows and middle Death Valley) which cluster together more closely. Note that there are still two separate "centers of mass" one for Ash Meadows and one for the carbonate springs of upper and middle Death Valley (Furnace Creek Region). The only difference in relationships seems to be that the Grapevine Springs do

not plot as close to Mesquite Spring and actually plot between Ash Meadows and the Furnace Creek springs.

If the analysis includes the trace elements of the data set a slightly different picture emerges (Figure 24, page 43). Once again, Saratoga, Surprise and Scotty plot away from the clustered group. The "clustered group" in this instance is not quite as clustered as in Figure 23 though, and Cold Spring is as near the "clustered group" as the Mesquite and Grapevine Springs are. This plot is important for several reasons. The difficulty of finding accurate values of concentrations on the part per trillion level might cause some speculation of results. Specifically, one might guess that the problem might exaggerate PCA plot results. From the following figure one may now see that this problem is not likely.

Figure 24: Scatterplot: Ash Meadows and Death Valley (trace elements)

Substitution of Missing Values: mean, zero, detection limit

Many of the elements examined have at least some values of "not measured" or "not detected" in at least one case. Only two dimensional scatterplots of spring classifications based on first and second principal components are shown here. The data, eigenvalues, and factor loadings are including in Appendix D.

From the graphs on page 44 and 45 one can see that although there may be slight differences in exact position with respect to certain springs, the groupings are essentially very similar. Saratoga plots alone, the upper Death Valley springs (indicated as Death Valley I on page 7) group together. Cold is separate from Ash Meadows and the Furnace Creek springs (middle Death Valley- indicated as Death Valley II on page 7) plot together near the Ash Meadows springs, in each analysis. These results would suggest that substituting means for empty cells in the data set is not an unreasonable method.

Figure 25:

Figure 26:

Reproduced with permission of the copyright owner. Further reproduction prohibited without permission.

Figure 27:

Scatterplot; Ash Meadow and Death Valley Substitute Zero

Geographic information Systems

Map

As explained in Chapter 4, the first graphical product generated was a location map. The location map is included as Figure 28 on page 47. Springs and wells are labeled to correlate with a listing at the bottom of the map. For the Ash Meadows and the Grapevine springs locale the spring symbols used to identify each spring overprinted one another due to their proximity. While the proximity actually does make it difficult, coordinates for each data point are included in Chapter 2 to assist in identification. All spring and well locations are projected in the Nevada State Plane Coordinate System. The Fipszone 2702, the spheroid is GRS80, the datum is NAD83, and the units are in meters.

Figure 28:

Reproduced with permission of the copyright owner. Further reproduction prohibited without permission.

⁴⁷

TINS

According to Clarke's Analytical and Computer Cartography, TINS are best suited for modeling overland flow, stream hydrology, and erosion. From the TIN contour map in Figure 29 one might guess that this is not an appropriate method to create contours for chemical concentrations. The corners and straight lines, while mathematically exact and a true representation of raw data, are not necessarily repeated in nature. Please note that one major difference between TIN modeling and kriging is that TIN modeling does not incorporate the influence of neighbors into the analysis. Because of this the TIN results (although created using all point locations) is cropped to show only Ash Meadows, some of the Nevada Test Site, and the Furnace Creek region of Death Valley. As a result these maps are represented at a different scale in hard copy than those created by kriging.

Figure 29**: Uranium Concentrations (TIN)**

Chapter 6: RESULTS AND DISCUSSION

Results From PCA and Agglomerative Cluster Analysis

Only the two dimensional Principal Component Analysis results are illustrated here. Supporting parts of the analysis are presented in Appendix D. In this section, the mean values have been substituted in for empty cells. The analyses in this section are performed in different ways based on inclusion. The first considers all cases (water sampling locations) and is called All Springs and Wells. The second eliminates perched waters and waters flowing out of volcanics and is called Carbonate System Only. Lastly, only Ash Meadows and Death Valley are examined in Ash Meadows and Death Valley Only. The variables included in all of these analyses are trace elements (excluding REEs) and are listed in Appendix D.

Results All Springs and Wells

Before looking at the results of the first analysis, (Figure 30 and 31 on pages 50 and 52) one should know that since the rock through which the water flows is important to determine water chemistry (Claassen, 1983), the signature of rock type might be as well represented on the graph as the signature due to groundwater flow paths; however, one may still gain a broader understanding of southern Nevada aquifer systems by studying this graph if one considers waters discharging from similar rocks.

Figure 30: Scatterplot: All Springs and Wells

The first relationship to notice is that the three springs which issue from volcanic tuffs on the test site. Cane, Tippipah, and Topopah plot uniquely. These springs are outliers , plotting not only separate from the other groups of springs, but from each other as well (especially Cane and Tippipah springs). There are several possible reasons why these springs are different in addition to their spatial distance from one another. As mentioned previously, these springs issue from volcanic tuff. Perhaps the spring water flowing through this part of the aquifer system is not related to the groundwater of the other hydrologie subdivisions within southern Nevada such as the Amargosa Desert or Death Valley. Since this water is perched (Thordarson, 1965) so this explanation seems adequate to account for differences between the rest of the samples and differences between Tippipah, Topapah, and Cane Springs. If the water is perched then the water chemistry of these three springs is likely more dependent on the effect of water movement through local rock or possibly an ancient water

table which is no longer continuous and is behaving as closed system with respect to horizontal flow. In other words the perched waters are not hydraulically connected to the rest of the regional groundwater flow system. Since this is true, trace element geochemical concentrations of perched waters might prove less useful in understanding regional horizontal flow relationships.

Figure 31: Scatterplot: All Springs and Wells (changed scale-Cane and Tippipah cropped out by cutting off the graph for Factor 1 values greater than 2 and less than -1.5)

There are similarities of the Ash Meadows group to the Pahranagat group when one examines Figure 31 in terms of first and second principal component values. First principal component values overlap and second principal component values are quite similar for Pahranagat and Ash Meadows waters. Since 60% of the water in Ash Meadows is hypothesized to come from the Pahranagat Valley, it is understandable that the waters would seem similar in principal component analysis (Winograd & Thordarson, 1975). There is a third

group which plots along with Pahranagat Valley and Ash Meadows waters. This is the middle Death Valley group which contains Texas, Nevares, and two of the Travertine springs. All of these springs issue from carbonate rock or alluvium composed of principally carbonate rock debris and consequently are thought to be from the lower Paleozoic carbonate aquifer. This group is slightly different from the Ash Meadows and Pahranagat group in terms of average first and second principal component values, but still has principal component one and two values which overlap those of Ash Meadows and Pahranagat Valley. Saga and Army wells are located on the Nevada Test Site and are clustered by PCA of trace element concentrations closely around three of the Ash Meadows springs and the Pahranagat Valley springs.

The upper Death Valley springs, Scotty's Castle, Surprise, and Mesquite. Scotty's and Surprise flow from volcanic rock are clustered together by PCA of trace element concentrations. Mesquite spring flows from alluvial material (Johannesson et al., 1995). At least two of the springs in this cluster are quite similar in chemistry to a cluster of wells which includes Cinderlite, Airport, and Lathrop. The other upper Death Valley springs plot together with factor two values similar to those of the group just described, but the factor one values are quite different. This is likely because the waters of the Grapevine springs have a carbonate aquifer source in comparison to volcanic rock. Both Soda and Coffer well waters plot apart from all other groups. Topopah Spring plots with Coliseum well waters.

In this particular analysis molybdenum, tungsten, titanium, zirconium, and germanium seem to be most important to Factor 1. This means that most of the 21% of the variance in this particular data set is caused by different amounts of these elements. Nickel, arsenic, and vanadium are most important to Factor 2 (refer to Factor Loadings Matrix in Appendix D). Factor 3 is commposed

Reproduced with permission of the copyright owner. Further reproduction prohibited without permission.

primarily of selenium, cobalt, and tin. Molybdenum and tungsten are in Group VIA and germanium is from IVB (Greenwood & Eamshaw, 1984). These elements have ionic radii ranging from 59 to 147 pm (Greenwood & Eamshaw, 1984) which is fairly similar in size to the range of ionic radii involved in substitution within the crystal lattices of clay minerals (Hurlbut & Klein, 1977). In Factor 2, arsenic, nickel and vanadium are important. In the periodic table of elements arsenic is in group VB. Vanadium is in group V. Nickel is in VIII.

Figure 32 on page 53 illustrates the hierarchical classification of all springs and wells (these figures use the same data sets that the Principal Component Analyses did in the previous section). One can see that Cane and Tippipah are by far the most different waters in comparison to the rest. Saratoga seems to be like Coffer Well. These two groups are both very different from the remainder of the waters. Again these waters are classified as being different based on trace element concentrations.

With the exception of the four outliers (Cane, Tippipah, Saratoga, and Coffer), construction of dendograms reveals that the bulk of the waters divides itself into two groups: 1)Ash Meadows, and the carbonate springs of Death Valley 2)Army, Lathrop, Cinderlite, Airport, Colleseum, Hardrock wells, Pahranagat Valley, and waters from volcanic rocks. In addition:

- a) The bulk of Ash Meadows springs are together.
- b) Saga well and the carbonate springs of upper Death Valley are grouped together.
- c) Mesquite is different from the neighboring springs in volcanics and carbonates (the Grapevine Springs, Scotty's Castle, and Surprise Spring).

The Pahranagat waters are clustered with the remaining wells .With the exception of Cold Spring, Nevarres Spring, and Mesquite Spring, most of the

clustering agrees with what a common geology and proximate locations would demand. In other words, waters near each other, which flow from similar surficial geology, are chemically similar with respect to trace element concentrations.

Discussion: All Springs and Wells

Vertical Movement of Groundwater on the Nevada Test Site:

Other studies have previously indicated that in this local area water leaks downward from the shallow aquifer into the regional groundwater system (Winograd & Thordarson, 1975, Peterman and Stuckless, 1993). The results in Figure 30 support the hypothesis that ground water is moving vertically. The carbonate waters of Ash Meadows, the waters from springs in volcanics, and the waters from wells in volcanics on the Nevada Test Site, each have different trace element chemical characters. Some of the NTS wells plot between the carbonate waters and the waters flowing from volcanic rock. The intermediate

composition of waters like J12 and J13 may indicate that there is an upward component of flow from the carbonate aquifer which mixes with waters more typical of the tuff aquifers in some portions of the Nevada Test Site. Perhaps the chemistry differences are also related to changes in solubility which might occur in vertical movement of ground water from one rock type to another.

There are other reasons to believe that vertical flow is important in this area. Most likely the influence of glass dissolution and the replacement of divalent ions with monovalent ions changes the waters' chemistry as the waters move down through the subsurface (White, Claassen, & Benson, 1980), which might then change once again if these waters mixed with carbonate waters. This is supported by high loading of lithium, rubidium, and cesium (monovalent in ion form) in the factors of some principal component analyses (Appendix D) which is important in defining the differences between the Nevada Test Site cluster and the Ash Meadows & middle Death Valley clusters. In other words, since these elements are appropriated high loadings in the factors, most likely it is differences in these concentrations which account for a large part of the difference in both the "X & Y Direction" on the plots. Generally elements in the same family behave similarly. Since the wells are cased at different depths and the springs are obviously from a far different datum, these difference may be due to how much the waters of the tuff aquifer waters are mixing with the carbonate aquifer at certain locations and depths. Since the waters of Topopah spring appear to be much more similar to the well waters of the Nevada Test Site, but are assumed to be perched in this study, one might conclude that there is some reason other than the fact that the waters are perched which makes Tippipah and Cane springs plot so uniquely.

Perched Water on the Nevada Test Site:

The most dramatic observation to be made of the PCA : All Springs and Wells, is the distance of outliers, Tippipah and Cane springs (of the Nevada Test Site, from the other clusters. According to Thordarson, 1965, all springs in Yucca Flat are perched waters retarded by tuffs. From Winograd and Thordarson (1975), it is known that ground water tables of Frenchman Flat, Yucca Flat, and Jackass Flat differ by only 9 feet in elevation. It is because of these two studies that the springs of the Nevada Test Site are assumed to perched waters (in addition to chemical characteristics of the springs). The very different chemistry of Tippipah and Cane springs supports this classification but Topopah appears to be slightly different. Topopah plots closer to some of the Nevada Test Site wells and so it is not so obviously different from the rest of the clusters. This might be explained simply by the likelihood that perched waters which are not hydrologically related might have different chemistries dependent on local conditions. The springs are located in flats so perhaps there are constituents of the playa deposits (dried lake beds) which determine the positions of these outliers.

According to principal component analysis of trace element data (Figures 30 and 31), the waters from wells of the NTS loosely gather themselves into a staggered, widespread cluster, while Cane and Tippipah springs (also on the Nevada Test Site plot some distance from them. Perhaps the reason for the existing range in Factor 1 and Factor 2 for these waters is influenced by chemical changes that occur or do not occur over certain flow paths. In other words chemical heterogeneities between different tuffs may be strong enough to influence ground water chemistry and possibley cause noticeable dispersion of the clusters of waters. For the perched waters, which are less important to regional flow of groundwater, local conditions are most likely more important in determining the trace element chemistry of the waters.

Results: Carbonate System Only

This analysis includes only those springs which are thought to be the result of springs flowing from the lower carbonate aquifer through alluvial deposits or directly to the surface and all wells. Supplementary information is included in Appendix D for the following Figure 43.

Again the Ash Meadows group and the mid death valley group plot together. Pahranagat appears just as different from these two groups as Mesquite does (Mesquite plotted nearer Scotty's and Surprise in first analysis: All Springs and Wells). Groups of wells plot together, one with the Ash Meadows and middle Death Valley springs while others are more isolated.

In this analysis Factor 1 accounts for approximately 20% of the variance in this study is most substantially composed of tungsten rubidium, lithium, arsenic, antimony, and germanium. Factor 2 explains about 17% of the variance

and is most heavily influenced by the variables rhenium, and cobalt. Lithium, chromium, uranium, and antimony are important to Factor 3 (15% of the variance) Tungsten chromium, uranium, and rhenium are thought to form conservative oxyanion species in this system. Perhaps it is best to focus on the these elements that occur in this conservative oxyanion species when looking at regional flow patterns; if the entire groundwater system of southern Nevada were well mixed, and oxidized the concentrations of these oxyanions would be relatively uniform and not so important to the first three principal components which explain \approx 50% of the variance within this analysis. Since these factors seem to be important perhaps this supports mixing.

This analysis is useful because it may illuminate which waters would yield the most reliable result when put into a model. This is not to say that certain data points should be ignored and cut from the study, but there may be a need to eliminate some data from a spatial analysis in order to better understand regional flow. This anomalous waters may be due to problems with sampling and analysis, or simply a small scale heterogeneity not representative of the locale. An example of one of this might be the exclusion of Bradford data from a set of data points chosen to represent Ash Meadows.

Discussion: Carbonate System Only

Groundwater Flow Between Sut)-basins

Many studies have indicated that water from the Pahranagat Valley is a large component of the flow at Ash Meadows by using isotope and other kinds of data(Winograd & Thordarson, 1975). The small distance between the centers of mass of clusters from my study would intuitively support this, but now some attempts at modeling mixing of consen/ative oxyanions should be made. Modification of the mixing model PHREEQE is suggested. This would require making changes in both the program and the database to make use of trace

element data (particularly those elements who are thought to behave conservatively in southern Nevadan and southeastern California groundwaters).

From Peterman and Stuckless (1993)one can see the trend of increasing *587* Sr from the Nevada Test Site to Ash Meadows to the middle Death Valley springs. This can possibly be correlated to the trend of the centers of mass of the corresponding clusters with respect to Factor 2 of several principal component analyses. Important variables for this factor from different analyses are below, many of which behave conservatively:

Analysis of entire data set: lithium, vanadium, uranium Analysis of Carbonate System Only (+wells): cobalt, rhenium Analysis of Ash Meadows and Death Valley only: chromium, uranium, rhenium, and strontium, lithium, molybdenum

Czarnecki & Waddell (1984), modeled groundwater in southern Nevada through the use of finite element simulation. The results of this study indicate that there is a component of middle Death Valley flow which is from the Nevada Test Site and a component of the more northern reaches of Death Valley which is from the Oasis Valley. Both of these observations are supported by PCA of trace elements by the fact that the center of mass of the Nevada Test Site cluster lies between the Ash Meadows and middle Death Valley clusters with respect to Factor 1 (not just simple trend in Factor 1 from Nevada Test Site to Ash Meadows to middle Death Valley).

This support is not specifically for an additional source from the Oasis Valley, but is indicative of some additional source (other than Ash Meadows and the Nevada Test Site). This support is based on the idea that without another source one would think that middle Death Valley waters would plot between Ash Meadows and the Nevada Test Site. It is certainly possible that this observation could be due to the effects of source rock on groundwater geochemistry
signatures, but the springs in both Ash Meadows and the Furnace Creek region (middle Death Valley) are all assumed to flow from the same carbonate aquifer and issue from carbonate rocks or carbonate rich sediments.

Principal component analysis also reveals a cluster of northern Death Valley waters (with considerably large differences in Factor 1-in most analyses) that is further removed from the other clusters. If the Oasis Valley does contribute to Death Valley perhaps the percentage of Oasis Valley water is greater in the upper reaches of Death Valley which could statistically detach this group from other waters. Possibly the upper Death Valley springs are not only chemically unique because of source rock, but also due to mixing of waters. If one notes, only Scotty's and Surprise springs have volcanic source rock. Perhaps source rock does still play a key role and mixing is occurring between the waters of the volcanic rock springs and the Grapevine and Mesquite springs. The variables of higest loading in Factor 1 may provide some clues as to which effects are more dominant. Again there seem to be some chemical elements which consistently are combined into Factor 1. These are explained below:

Analysis of entire data set: Tungsten, molybdenum, thallium, germanium, zirconium

Analysis of Carbonate System Only (+wells): Lithium, arsenic, tungsten, rubidium, antimony, germanium

Analysis of Ash Meadows and Death Valley only: Lithium, selenium, vanadium, tungsten, uranium, rubidium, barium

Using this knowledge of what variables are important in defining the waters of southern Nevada (based on trace element geochemistry) should prove useful in future efforts to model mixing between hydrologie sub-basins.

This analysis is very similar to the one performed using the entire data set and is located in Appendix D. In the first run it can be seen that basically there

is an Ash Meadows/Death Valley group and a second Well Data and Pahranagat group (ignoring outliers- Cold, Saratoga, and Coffer). After deleting Cold, Saratoga, and Coffer, the carbonate waters generally divide themselves into a Mid Death Valley group, an Ash Meadows group, an Upper Death Valley group, and several groups made up of wells, and Pahranagat waters. Mesquite clusters uncharacteristically for an Upper Death Valley water and Bradford (AM), and Ash (PAH) also seem to stray from the expected clustering trends of their geographic groups.

Sub-basin Classification:

In this study sub-basins are classified different ways, a technique used in Mifflin, 1968 which defines studies as being "regional," "local," and "small local". In this study analyses were conducted on the regional scale with waters from all source rocks considered, on the regional scale with only waters obtained from carbonate sources and wells, and on the local scale. Waters that are thought to be perched (Tippipah, Topopah, and Cane Springs) are omitted from classification. Principal component analysis classifies waters on the regional scale as follows: Amargosa Desert- all Ash Meadows springs and all of the middle Death Valley springs from the Furnace Creek region, and some Nevada Test Site wells including Saga, Army, and usually Coffer (depending on list of elements in analysis); lower Death Valley- Soda Well and Saratoga Spring; Volcanics and Nevada Test Site-Cinderlite, Mesquite, Scotty's, Surprise, Colleseum, Lathrop, and Airport; Pahranagat-Hiko, Crystal, and Ash springs; Upper Death Valley Nonvolcanics-Upper, Middle, and Lower Grapevine springs. Hardrock plots near Topopah when a study of this scale is considered.

On the same physical scale but considering waters only thought to be carbonate and wells, the following classifications were made: Amargosa Desertall Ash Meadows springs and all of the middle Death Valley springs from the

Furnace Creek region, and some Nevada Test Site wells including Saga, Army, Cinderlite, Lathrop, and Airport; lower Death Valley- Soda Well and Saratoga Spring; IS-Colleseum, Coffer, and Hardrock; Pahranagat-Hiko, Crystal, and Ash springs; Upper Death Valley Nonvolcanics-Upper, Middle, and Lower Grapevine springs.

Due to the time restraints, the author did not study the structural geology in detail; however, efforts were made to understand the general regional and local geology of study areas, and in doing so a general correlation was noted and compares well to the results of this study. If one refers to Figure 4: Structure in the South-Central Great Basin (Stevens, Stone, & Belasky) on page 10 and is familiar with the location of the springs and wells in the study one might observe that most waters which clustered in this study were located within boundaries. Most of these areas are bounded by shear zones and mountain ranges, although one should note that in this study data points are not distributed in homogeneous grid, but rather as nature dictated (for springs) and man chose for purposes other than this study (wells).

The Pahranagat waters are north of the Las Vegas Valley Shear Zone and west of the the Sheep Range. The Nevada Test Site waters are northeast of the Furnace Creek Fault Zone and north of the Las Vegas Valley Fault Zone. It might be interesting to note that if the entire fault lengths are not represented in this map, the faults might dissect the Nevada Test Site. If this is the case, this possibility might also explain some of the differences in waters in this area. The Ash Meadows springs are all west of the Spring Mountains and near the Stewart Valley Fault.

The middle Death Valley springs are just west of the Furnace Creek Fault Zone, east of the Panamint Range, and southwest of the Funeral Mountains.

The upper Death Valley springs are also located on the Furnace Creek Fault Zone but are just east of the upper reaches of the Panamint Range.

As mentioned in Chapter One, there has already been work done to define hydrologie sub-basins in southern Nevada. The Perfect thesis from the Colorado School of Mines (1991) defines sub-basins using cluster analysis of major elements and ions. The groupings were defined in SAS and other tools and then defined by the spatially correlated surficial geology. These identified clusters were then used to delineate spatial boundaries. Each bounded area was labeled with most common cluster definitions. Included were comments pertaining to relationships between nearby clusters which would include a signature not due solely to the source rock, but also flow parameters such as mixing and chemical evolution of water over time.

Comparisons may be made for sub-basins common to both Perfect's major ion and element data and the Harry Reid Center trace element data, but only on check basins. If spatial boundaries were drawn for subbasins based on this study most boundaries would rely on inference for their placement. Where locations are common to both studies this study suports that of Perfect. In general source rock seems to be dominant when statistically clustering data by trace element chemistry, as Perfect indicated by major ion chemistry. When considering waters of similar source rock however, one might possibly be able to interpret certain relationships between subbasins. Conservative chemical species have the potential to be a powerful tool in these kinds of studies.

Principal component analysis generally clusters waters into the same groups whether one uses a balanced variable set (each column of the periodic table is fairly equally represented) or a weighted variable set (more elements from certain columns on periodic table included). It is true that outliers are further distanced from the "clusters" by a weighted column analysis. Both

analyses may be useful and neither should be ignored as a tool in statistical analysis of groundwater samples.

Results: Ash Meadows and Death Valley Only.

Since there are only two springs of these two areas flowing from volcanics it seems logical that these two, Surprise and Scotty's, cluster together and apart from the rest. Saratoga, with anomalous values for almost every variable Figure 34:

concentration analyzed, obviously plots separate from every other spring. In general however two groups of springs plot with each other. Once again these two groups are the mid and upper Death Valley grouping, and the Ash Meadows grouping. In addition to Scotty's, Surprise, and Saratoga, Cold Spring waters (from Ash Meadows) are characterized by anomalous concentrations so does not group with the rest of Ash Meadows. Nevares plots with most of the Ash Meadows springs and like Cold has a different temperature than nearby springs(Harry Reid Center). It may be interesting to note that with respect to Factor 2 both Cold and Nevarres Spring are about 1 unit more (in terms of

values of principal components) than the "center of mass" of the group in which one might guess that they would cluster (Ash Meadows and middle Death Valley respectfully). The supporting information is in Appendix D and is the same for the Figure 25.

Tungsten, selenium, vanadium, uranium, rubidium, and barium are most important to Factor 1. Lithium, chromium, uranium, molbdenum, rhenium, and strontium heavily influence Factor 2. Some of these same variables have proven to be important in preceding analyses and are probably important in defining relationships between Ash Meadows and Mid and Upper Death Valley carbonate springs.

On the local scale examined. Ash Meadows and Death Valley, three main groups may be defined: Ash Meadows-13 springs previously identified as Ash Meadows; middle Death Valley-Texas, Travertine A&B, and Nevares springs; and upper Death Valley-Scotty's, Surprise, and the three Grapevine springs.

All of the waters in this analysis have been included in previous discussions, and there are no grand scale differences between these results and those of an agglomerative cluster analysis of a carbonate waters only section. Once again Cold and Saratoga Springs are outliers. Scotty's and Surprise group together. The remaining upper Death Valley waters cluster out neatly together, and until deletion of a case the remainder of Ash Meadows and Death Valley waters remain undivided. Upon deletion of Cold Spring and the volcanic waters, Mesquite separates from the Grapevines, and the MDV waters separate from the AM waters. Both Big and two thirds of the spring waters in the Point of Rocks region pull out from the Ash Meadows cluster.

Discussion: Ash Meadows and Death Valley

This analysis is useful because it may illuminate which waters would yield the most reliable result when put into a model. This is not to say that certain

data points should be ignored and cut from the study, but there may be a need to eliminate some data from a spatial analysis in order to better understand regional flow. This anomalous waters may be due to problems with sampling and analysis, or simply a small scale heterogeneity not representative of the locale. An example of one of this might be the exclusion of Bradford data from a set of data points chosen to represent Ash Meadows.

Results from Spatial Analysis with Geographic Information Systems and Discussion

ARC/INFO makes analysis of spatial variance relatively simple in the kriging command. Kriging in this software creates a coverage of variance which can then also be contoured. This provides a map of variance. Areas of higher variance are areas where more data points are needed to improve the quality of the chemical concentration contour map. It is probably obvious that many more data points would be useful in this analysis just from looking at the location map. This topic will be further addressed in the Conclusions chapter.

With the definitions used in this study, contour maps had many interesting relationships. Contour maps seem to fall into four groups which display similar patterns. The patterns are more obvious in some contour maps than others but these differences could be minimized if contour intervals were adjusted. Almost every contour map fits into one of three classes. Groundwater movement can only be in one direction (downgradient) at any discrete point. If there is a way of analyzing flow with this data only one group of the three should be chosen to do it. Other groups should then represent the effects of other phenomenon. All conclusions are based on the major assumption that the data is sufficient to produce true concentrations throughout the study area when kriged. From

variance maps one can see that this assumption is more valid in some areas than in others.The groupings and the trends are classified and describe below:

Group 1. thallium and nickel, exhibit similar features when contoured. These are included in Figures 35-36 on pages 68-69. Each has a zone higher concentrations. The zone trends roughly east northeast-west southwest through the central portion of the study area. The zone generally trends across Ash Meadows and the Furnace Creek or middle Death Valley areas. Spring or well locations are indicated on the map by an "0". Refer to the location map on page 48. The concentration maps (thallium and nickel) illustrate these observations.

Group 2, cesium and antimony, have a different contour pattern when kriged. Concentration contour maps for cesium and antimony are included as Figures 37 and 38 on pages 70 and 71. The two maps show areas of higher concentrations south of the Pahranagat Valley springs, south of the upper Death Valley springs, and in the area around Saratoga Spring. Concentrations are generally lower and more uniform throughout the study area with respect to Cesium and Antimony. Since there are several isolated highs that do not seem to correlate with geology, water table elevation, or topography perhaps they are more related to anthropogenic effects. Again, locations are indicated on the map by an $"O"$.

Reproduced with permission of the copyright owner. Further reproduction prohibited without permission.

Reproduced with permission of the copyright owner. Further reproduction prohibited without permission.

Group 3, Uranium, Molybdenum, and Tungsten, also show corresponding concentration contour patterns. The trends of all three of thes elements are 1)highest concentrations in the vicinity of the Nevada Test Site and 2)a generally increasing trend of concentrations from the southeast portion of the study area toward the northwest. The steepest concentration gradients are in the vicinity of Saga, Cinderlite, Airport, Lathrop, J12, and J13 wells which are in or near the southern portions of the Nevada Test Site. Figures 40-42 on pages 74-76 are concentration contour maps for molybdenum, tungsten, and uranium. The chemical elements contoured in this group are thought to behave conservatively under certain conditions. Perhaps these maps show the conservative behavior in the "flat parts" and physical boundaries along the steep gradients. If these physical boundaries were fast pathways the contour maps might further support the ideas that mixing is occurring in different areas, for example between Ash Meadows and Death Valley.

Group 4, Germanium and Rubidium, share an interesting contour pattern. In general, there are higher concentrations in the southwestern portion of the study area. The region of higher concentrations stretches from the vicinity of the upper Death Valley to the southwestern comer of the map, west of the Furnace Creek-Death Valley Fault Zone. This area of higher concentrations is surrounded by an area of steeper concentration gradient. Figure 43 on page 77 is the concentration contour maps for rubidium. Rubidium and Germanium consistently have high loadings in Factor 1

Group 5 is made up of contour maps which do not fit into any pattern in particular. These include arsenic, selenium, titanium, cobalt, gallium, vanadium, and rhenium. These contour maps follow on pages 79-85 as Figures 44-50. Cobalt concentrations are fairly constant throughout the area. Gallium has concentration highs in the vicinity of the NTS, south of the Furnace Creek area.

and north of Hardrock and Colleseum wells. The contour map of Vanadium concentrations has curvilinear zones of steep chemical concentration gradients approximately ninety degrees in orientation from the trend of the Furnace Creek Fault Zone as well as the steepest hydraulic flow gradients in the study area. A more detailed map of the potentiometric surface of the lower carbonate aquifer is included below as Figure 39 for comparison. Selenium concentrations are highest in the vicinity of the Nevada Test Site and generally decrease from the

Figure 39: Potentiometric Map of Lower Carbonate Aquifer (reformat from the Las Vegas Valley Water District of Thomas and Crabtree, 1986)

Reproduced with permission of the copyright owner. Further reproduction prohibited without permission.

⁷⁶

northwest to the southeast. There is a local high of arsenic concentrations in the vicinity of the Nevada Test Site and near the Pahranagat springs, with a general trend of decreasing concentrations from the northeast to the southwest. Rhenium concentrations also generally decreas from the northeast to the southwest. Titanium concentrations are generally greater in the southern half of the study area than in the northern half with concentrations in the northeast quadrant of the study area being more uniform.

79

⁸²

Reproduced with permission of the copyright owner. Further reproduction prohibited without permission.

Chapter 7:

CONCLUSIONS AND RECOMMENDATIONS FOR FURTHER STUDY

The strongest conclusion to be made is that more sampling is needed to accurately spatially analyze data on this scale. Using the variance maps and known locations of springs the following list of waters should be used in further analysis;

- Spring Mountains
- Black Mountains
	- Montgomery Spring
	- Salsberry Spring
	- Sheephead Spring

Panamint Range

- Fivemile Spring
- Colter Spring
- Emigrant Spring
- Montezuma Range
	- Railroad Spring
	- Indian Spring
	- McNamara Spring

Death Valley and Funeral Mountains

- Triangle Spring
- Daylight Spring

It seems that trace elements support previous efforts to define hydrologie subbasins and the idea that mixing is occurring between Ash Meadows and the Furnace Creek area of Death Valley as well as between the Nevada Test Site and Furnace Creek. In addition, contouring of these trace elements indicates that there are several physical or physico-chemical parameters which are affecting distribution of trace elements concentrations throughout the study area. Possibly these are rock chemistry, flow paths, changes in pH or reduction/oxidation regimes. Because of relationships between components of principal component analysis and contouring patterns it is possible that a restricted list of elements could be analyzed for (within the study area) to gain an understanding of the hydrogeochemistry of the region.

When performing statistical analysis it is important to understand the techniques and methods the software package uses. Bin sizes should be selected carefully (or stem and leaf diagrams included). An appropriate method for substitution of missing values should be chosen for missing values. In this case either substitution of detection limit, mean, or zero are appropriate.

Other recommendations for additional study are numerous. Besides additional sampling at different locations throughout the study area, it would probably be beneficial to create or obtain additional ARC/INFO coverages; for example, surficial geology, geology at interval depths, and potentiometric maps for aquifers. If the user has enough disk space to manipulate these coverages within ARC/INFO then the contour maps should be draped or overlain over these different coverages. Then more accurate observations may be made or queries performed on the data set. There is certainly a wealth of borehole geology data on the Nevada Test Site. Much effort will be needed. If it is even possible, to gather enough information in other areas to have enough data density for valid interpretations.

Appendix A: Simple Plots

 $\frac{1}{4}$

 $\overline{1}$

 α

 \sim

 $\frac{1}{2}$

 \bar{a}

 $\frac{1}{2}$

 \sim

 $\frac{1}{2}$

§

 $\overline{}$

«

 \rightarrow

ჯ,

 $\frac{96}{2}$

 $\bar{1}$

 δ

Reproduced with permission of the copyright owner. Further reproduction prohibited without permission.

 $\overline{5}$

¹⁰²

Appendix B: Periodic Table of Elements

Appendix C: Histograms and Stem and Leaf Plots

SR 86

SB 121

n o

BA 135

Reproduced with permission of the copyright owner. Further reproduction prohibited without permission.

Ill

Reproduced with permission of the copyright owner. Further reproduction prohibited without permission.

114

Reproduced with permission of the copyright owner. Further reproduction prohibited without permission.

2.10 M-Grape# .21 HIKO

Note: Only a partial list of cases wth the value 2.10 are shown in the table of upper extremes.

 \sim \sim

 $\mathcal{A}^{\mathcal{A}}$

Reproduced with permission of the copyright owner. Further reproduction prohibited without permission.

 \perp \perp

AS75_1 SMEAN(AS75) Valid cases: 39.0 Missing cases: .0 Percent missing: .0 Mean 18.0432 Std Em 1.5909 Min .5200 Skewness .1775 Median 18.6000 Variance 98.7103 Max 44.2000 S E Skew .3782 5% Trim 17.7930 Std Dev 9.9353 Range 43.6800 Kurtosis .0930 95% CI for Mean (14.8225, 21.2638) Frequency Stem & Leaf 5.00 0 • 01223 4.00 0 . 5789 3.00 1 * 022 10.00 1. 5556678899 9.00 2 * 002233334 4.00 2 . 5678 3.00 3 * 144 1.00 Extremes (44) Stem width: Each leaf: 1 case(s) 5 Highest SPRING 5 Lowest SPRING 44.2 SODA .5 COLLESEUM 34.6 ASH 1.6 Topopah 34.0 Scotty#5 2.0
31.0 Mesquite 2.7 3.0 COFFER
Texas#5-
3.0 COFFER **28.0 Texas#5- 3.0 COFFER W182_1 SMEAN(W182)** Valid cases: 39.0 Missing cases: **Percent missing: .0 Mean 8.9097 Std Enr 7.1210 Min .0000 SkevMiess 6.2085 Median .4300 Variance 1977.642 Max 279.0000 S E Skew .3782 5% Trim 1.6257 Std Dev 44.4707 Range 279.0000 Kurtosis 38.6840 1.8700 S E Kurt .7410 95% Cl for Mean (-5.5060, 23.3255) IQR Frequency Stem & Leaf 21.00 n * ni111979??2?999?33444 2.00 0 . 68 1.00 1 * 3 3.00 1 . 788 4.00 2 * 0113 .00 2 . .00 3 * 1.00 3 . 5 2.00 4 * 44 1.00 4 . 6 4.00 Extremes (8.9), (11.9), (279.0) Stem wdth: 1.00 Each leaf: 1 case(s) Highest SPRING 5 Lowest SPRING** $\overline{\mathbf{A}}$ **279.00 Cane .00 COFFER 11.90 Tippipah .12 SAGA 8.91 HARDROCK .16 Rogers 4.60 Nevares# .17 ARMY**

U238_1 SMEAN(U238) Valid cases: 39.0 Missing cases:
Mean 4.5763 Std Err 8279 Min **Mean 4.5763 Std Err .**
Median 2.9500 Variano Median 2.9500 Variance 26.7288 Max 5% Trim 3.8511 Std Dev 5.1700 Range 29.9800 Kurtosis 15.5085 95% Cl for Mean (2.9004, 6.2522) IQR Frequency Stem & Leaf Percent missing: .0 .0200 Skevmess 3.5475 30.0000 S E Skew .3782 2.7000 S E Kurt .7410 4.00 0. 0056 2.00 1 . 17 14.00 2 . 23355677788999 7.00 3. 0004578 2.00 4. 45 2.00 5. 27 3.00 6. 009 1.00 7. 9 4.00 Extremes (8.9), (13.8), (13.9), (30.0) Stem width: 1.00 Each leaf: 1 case(s) 5 Highest SPRING 5 Lowest SPRING 30.00 SODA .02 COFFER 13.90 Saratoga .08 Topopah 13.80 COLLESEUM .52 Tippipah 8.90 Surprise .60 AIRPORT 7.90 Cold Spring 1.19 Nevares# M095_1 SMEAN(M095) Valid cases: 39.0 Missing cases: .0 Percent missing: .0 Mean 147.0789 Std Err 109.1761 Min 1.0800 Skewness 5.9512 Median 7.1000 Variance 464857.3 Max **5% Trim 23.9361 Std Dev 681.8044 Range 4225.920 Kurtosis 36.2821** 95% Cl for Mean (-73.9365, 368.0944) **IQR Frequency Stem & Leaf 6.00 0 * 112334 17.00 0 . 8.00 1 * 01222444 2.00 1 . 68 2.00 2 * 24 4.00 Extremes (147), (352), (701), (4227) Stem width: 10.00 Each leaf: 1 case(s) 5 Highest SPRING 4227.00 701.00 352.00 147.08 24.00 Cane Tippipah SODA mean Saratoga** 5 Lowest SPRING **1.08 1.99 2.23 3.00 3.64 HARDROCK AIRPORT COLLESEUM Topopah COFFER**

CS133J SMEAN(CS133) Percent missing: .0 Valid cases: 39.0 Missing cases: **.0797 Skewness 1.9479 Mean 2.9680 Std Err .4609 Min Median 2.3973 Variance 8.2843 Max 12.0000 S E Skew .3782 5% Trim 2.6526 Std Dev 2.8782 Range 11.9203 Kurtosis 3.6515 95% CI for Mean (2.0350, 3.9011) 2.0900 S E Kurt Frequency Stem & Leaf 6.00 0 * 011114 2.00 0 . 57 4.00 1 * 2334 5.00 1. 56799 3.00 2 * 003 2.00 2 . 79 9.00 3 * 123333344 3.00 3 . 556 1.00 4 * 0 4.00 Extremes (8.9), (10. Stem wdth: 1.000 Each leaf: 5 Highest SPRING 5 Lowest SPRING 125. M-Grape#** .080
L-Grape- .125 **11.000 L-Grape- .125 COLLESEUM 10.300 U-Grape# .139 Tippipah 8.862 ASH .194 SODA 4.000 Cold Spring .194 Mesquite BA135_1 SMEAN(BA135) Valid cases: 39.0 Missing cases: .0 Percent missing: .0** Mean 48.4645 Std Err 6.6610 Min **Median 43.0000 Variance 1730.392 Max 213.4020 S E Skew .3782 5% Trim 43.8038 Std Dev 41.5980 Range 213.1440 Kurtosis 5.8441 95% Cl for Mean (34.9800, 61.9490) IQR 45.3539 S E Kurt .7410 Frequency Stem & Leaf 8.00 0. 00011468 4.00 1** . **8899 .00 2 . 3.00 3. 599 6.00 4 . 001138 6.00 5. 034567 5.00 6. 13467 3.00 7. 377 1.00 8 . 1 .00 9 . 1.00 10 . 7 2.00 Extremes (146), Stem wdth: 10.00 Each leaf: 1 case(s)** 5. **Highest SPRING 5 Lowest SPRING 213.40 HARDROCK .26 Topopah 145.91 ASH .34 Tippipah 107.46 .79 Scotty#5 HIKO Point of Rx NE 1.34 CINDERLITE 81.00 1.74 AIRPORT 77.07 CRYSTAL**

TL205_1 SMEAN(TL205) Valid cases; 39.0 Missing cases: .0 Percent missing: .0 Mean .1851 Std Err .0231 Min .0290 Skewness 1.2419 Median .1730 Variance .0208 Max .6700 S E Skew .3782 5% Trim .1724 Std Dev .1442 Range .6410 Kurtosis 2.0190 95% CI for Mean (.1384, .2319) Frequency Stem & Leaf 8.00 0 * 23334444 7.00 0 . 5555699 $1 * 4$ **7.00 1 . 6677889 6.00 2 * 000224 4.00 2 . 6799 1.00 3 * 1 1.00 3 . 8 2.00 4 * 01 1.00 4 . 6 1.00 Extremes (.67) Stem width: .10 Each leaf:** Highest SPRING 5 Lowest SPRING
67 Trav-B#5 03 Scottv#5 $\overline{\mathbf{5}}$.67 Trav-B#5 .03
.46 Rogers .03 **.46 Rogers .03 Surprise 41 Cold Spring .41 Cold Spring .03 .41 HIKO .03 AIRPORT .39 SAGA .04 Mesquite SN117_1 SMEAN(SN117) Valid cases: 39.0 Missing cases: .0 Percent missing: .0 Mean .0488 Std Em .0043 Min .0200 Skewness 2.2126 Median .0435 Variance .0007 Max .1400 S E Skew .3782 5% Trim .0456 Std Dev .0269 Range .1200 Kurtosis 4.8462 95% Cl for Mean (.0400, .0575) IQR .0158 S E Kurt .7410 Frequency Stem & Leaf 3.00** 2* **000 2.00** 2 **. 78 6.00 3* 022334 5.00** 3 **. 67888 4.00 4" 1123 10.00** 4 **6888888888 4.00 5 * 0023 .00 5 .** 6* **1.00 4.00 Extremes (.102), (.110), (.130), (.140)** Stem width: .01
Each leaf: 1 case(s) **Each leaf: 5 Highest SPRING 5 Lowest SPRING .14 King's Pool .02 Trav-B#5 .13 Saratoga .02 Texas#5- .11 Nevares# .02 Surprise .10 SODA .03 AIRPORT .06 Cold Spring .03 Jackrabbit**

SB121_1 SMEAN(SB121) Valid cases: 39.0 Missing cases: .0 Percent missing: .0 Mean 60.4398 Std Err 50.0844 Min **Median .2951 Variance 97829.59 Max 1944.000 S E Skew .3782 5% Trim 4.4889 Std Dev 312.7772 Range 1943.978 Kurtosis 37.2888 95% Cl for Mean (-40.9509,161.8304) IQR .5987 S E Kurt .7410 Frequency Stem & Leaf 13.00 0 * 0000011111111 10.00 Ot 2222222333 3.00 Of 445 4.00 Os 6777 1.00 0 . 9 3.00 1 * Oil 1.00 1 t 3 4.00 Extremes (60.4), (69.5), (269.0), (1944.0) Stem width: 1.00 Each leaf: 1 case(s) 5 Highest SPRING 5 Lowest SPRING 1944.00 Tippipah .02 Saratoga 269.00 Topopah .02 COFFER 69.50 Cane .04 HARDROCK 60.44 mean .07 COLLESEUM 1.30 ASH .09 SODA TI47.1 SMEAN(TI47) Valid cases: 39.0 Missing cases: .0 Percent missing: .0 Mean 141.1462 Std Err 108.3694 Min .4000 Skewness 5.5385 Median .6700 Variance 458013.2 Max 4068.000 S E Skew .3782 5% Trim 6.5362 Std Dev 676.7667 Range 4067.600 Kurtosis 31.9233 95% CI for Mean (-78.2362, 360.5286) IQR Frequency Stem & Leaf 5.00 8.00 9.00 6.00 5.00 1.00 1.00 4 . 00568 5 . 22234466** 6. **011122799 7 . 012334** 8 . **33558 9 .** 2 **10** 0 **4.00 Extremes (2.80), (141.15), (1270.00), (4068.00) Stem width: .10 5 Highest SPRING 5 Lowest SPRING 4068.00 1270.00 141.15 2.80 Tippipah Cane mean Topopah .40 Saratoga .40 Point of .45 Fairbanks .47 HARDROCK**

123

Reproduced with permission of the copyright owner. Further reproduction prohibited without permission.

.48 Jackrabbit

1.00

Scotty#5

NI60_1 SMEAN(NI60) Valid cases: 39.0 Missing cases: .0 Percent missing: .0 Mean .5589 Std Em .0648 Min .0500 Skewness **Median .5000 Variance .1638 Max 1.4200 S E Skew .3782 5% Trim .5402 Std Dev .4047 Range 1.3700 Kurtosis 95% CI for Mean (.4278, .6901) IQR .7400 S E Kurt 95% Cl for Mean (.4278, .6901) IQR .7400 S E Kurt .7410 Frequency Stem & Leaf 11.00 0 * 00001111111 4.00 Ot 2233 8.00 Of 44445555 4.00 Os 6677 5.00 0 . 88999 4.00 1 * 0011 2.00 1 t 23 1.00 1 f 4** Stem width: **Each leaf: 1 case(s) 5 Highest SPRING 5 Lowest SPRING 1.42 Bradford .05 AIRPORT 1.38 COFFER .06 Scotty#5 1.25 Point of Rx NE .09 Surprise 1.12 Point of Rx NW 1.10 HARDROCK .11 Cane Note: Only a partial list of cases with the value .11 are shown in the table of lower extremes. MN55_1 SMEAN(MN55)** Valid cases: 39.0 Missing cases: **Mean 2.0897 Std Em 1.2020 Min Median .3600 Variance 56.3458 Max 5% Trim .8335 Std Dev 7.5064 Range 95% Cl for Mean (-.3436, 4.5230) IQR Percent missing: .0 .0200 Skewness 6.0514 47.3000 S E Skew .3782 47.2800 Kurtosis 37.3086 1.9897 S E Kurt .7410 Frequency Stem & Leaf 21.00 0 * 000000000111111222234 3.00 0 . 568 3.00 1 * 134 2.00 1 . 56 8.00 2 * 00000000 2.00 Extremes (5.2), (47.3) Stem width: 1.00 Each leaf: 1 case(s) 5 Highest SPRING 5 Lowest SPRING 47.30 COFFER .02 Cold Spring 5.16 Topopah .02 Saratoga 2.09 CRYSTAL .04 Trav-B#5 2.09 HIKO .04 Jackrabbit 2.09 ASH .05 Point of Rx NE**

Note: Only a partial list of cases with the value 2.09 are shown in the table of upper extremes.

RB85J SMEAN(RB85)

Valid cases: 39.0 Missing cases: .0 Percent missing: .0 Median 17.0000 Variance 166.9620 Max 5% Trim 17.8389 Std Dev 12.9214 Range 54.5200 Kurtosis 2.6522 95% CI for Mean (14.6806, 23.0578)

**Mean 1.6668
185.4000 SE Skew .3782**

Frequency Stem & Leaf 3.00 5.00 9.00 10.00 6.00 2.00 0 * 034 0 . 57899 1 * 012233444 1 . 5677777899 2 * 000133 2 . 67 4.00 Extremes (50), (53), (55) Stem wdth: 1000 Each leaf:

SR86_1 SMEAN(SR86)

Valid cases: 39.0 Missing cases: .0 Percent missing: .0 Mean 783.2485 Std Err 136.7031 Min 6.2000 Skewness 2.6088 Median 608.2510 Variance 728821.7 Max **5% Trim 655.9001 Std Dev 853.7105 Range 4200.270 Kurtosis 8.1621** 95% CI for Mean (506.5075, 1059.989) IQR

Appendix D: Information for Factor Graphs from PCA

 \mathcal{A}

 $\mathcal{A}^{\mathcal{A}}$

 \pm

 \bullet

 $\langle \cdot \rangle$

 \mathcal{A}

 \mathcal{A}

 $\frac{1}{3}$

 \sim

ŧ

Reproduced with permission of the copyright owner. Further reproduction prohibited without permission.

للمستردات

المناسب

 $\bar{\gamma}$

 $\frac{1}{2}$

Information to Accompany Figure 22

Number of Eigenvalues

Reproduced with permission of the copyright owner. Further reproduction prohibited without permission.

 $\Delta \sim 10$

 ~ 10

 \sim 1 $-$

Information to Accompany Figure 23

 \mathcal{A}^{max} and \mathcal{A}^{max}

 $\ddot{}$

 ~ 10

Eigenvalues and Scree Plot(amanddv2.stal-Extraction: Principal components

 \sim

 \mathbb{R}^2 .

 $\frac{1}{2}$ and $\frac{1}{2}$ and $\frac{1}{2}$ and $\frac{1}{2}$

 $\bar{}$

 $\langle \cdot, \cdot \rangle$

لتناصب

¹³²

Information to Accompany Figure 24
Eigenvalues & Scree Plot(amanddv.sta)-Extraction: Principal components

Plot of Eiganvalues

Factor Loadings (Unrotated) (amanddy.sta)-Extraction: Principal components

 ~ 10

 $\overline{}$

 \mathcal{L}_{max} and \mathcal{L}_{max}

 $\overline{}$

 $\hat{\mathcal{A}}$

 $\frac{1}{i}$

 \bar{t}

 $\mathbf{1}$

 $\frac{1}{4}$

 \mathcal{A}

Reproduced with permission of the copyright owner. Further reproduction prohibited without permission.

للمسترد والمرابط والمتواز

المستند

 $\bar{\gamma}$

 $\ddot{}$

 \sim .

 $\label{eq:reduced} \frac{1}{2} \left(\frac{1}{2} \frac{1}{2$

 $\ddot{}$

Reproduced with permission of the copyright owner. Further reproduction prohibited without permission.

 $\mathcal{L} = \{ \mathcal{L} \}$.

 $=$ $-$

.

 $\bar{\beta}$

 $\frac{1}{2}$

 \sim \sim

 $\mathcal{A}^{\mathcal{A}}$ $\frac{1}{2}$

 $\ddot{}$

 $- - -$

t,

 $\frac{1}{2} \left(\frac{1}{2} \right) \frac{1}{2} \left(\frac{1}{2} \right)$

 $\bar{\gamma}$

 \sim $-$

 \hat{f} , \hat{f} , \hat{f} , \hat{f} , and

 $\frac{1}{2}$

 $\hat{\mathcal{A}}$

 $\ddot{}$

Reproduced with permission of the copyright owner. Further reproduction prohibited without permission.

and the state of the state of

الأرابية

 \sim $-$

 $\mathcal{L}^{\mathcal{L}}$

2004.
2004. prosent

 $\hat{\mathcal{L}}$

 \sim

 $\ddot{}$

 \sim

 \mathcal{A}

 $\frac{1}{2}$

 147

÷ $\bar{\mathrm{t}}$

 $\overline{}$ $\bar{\gamma}$ $\frac{1}{2}$

 \sim

 \bar{z}

 $\hat{\boldsymbol{\beta}}$

Information to Accompany Figure 30 and 31

Plot of Eigenvalues

Number of Eigenvalues

 \mathbb{R}^2

Information to Support Figure 33

Factor Loadings (Unrotated) (a1.stal-Extraction: Principal components

151

Reproduced with permission of the copyright owner. Further reproduction prohibited without permission.

	<u> Factor Scores (a1. stal-Rotation: Unifotated-Extraction: Frincipal Components</u> Factor Factor Factor Factor				
	Factor	2	3	4	5
			.99674	-00170	-70864
CRYSTAL	.53011	-77729		.10295	-93671
HIKO	.06068	-60088	1.32111		-1.13244
ASH	-48797	-22465	1.73294	-25937	
Big Spri	.43683	.26876	-12066	.27730	-81673
Bradford	.51654	1.10608	-86959	.36635	-07476
Cold Spr	.42737	1.88338	-1.49038	1.04739	-1.41111
Crystal	.69105	.87662	-55950	.50009	-67856
Fairbank	.96909	.47858	.09653	.83665	-1.21744
Jackrabb	.53788	.60996	-44573	.09377	-54326
King's P	.25957	.52174	.35436	.89184	-71433
Longstre	.67971	1.28214	-.49059	.84782	-80102
Pt Rx NE	.94865	.12085	-03340	-1.29605	.64872
Pt Rx NW	.68864	.37007	-09932	.31143	.16543
Rogers-M	.47932	.65171	.05498	-1.46541	-09852
Scruggs-	.42036	-13011	.45677	-02586	-31465
Grape#5-	-2.01494	1.06621	.84658	.02641	.33220
U-Grape#	-2.06333	1.23267	1.03116	.21642	1.13291
M-Grape#	-2.22068	1.41688	.98314	06413	1.24569
Mesquite	-.70749	-51295	-1.01409	$-03952.$	10201
Nevares#	-0.87713	.76814	.68539	.52644	1.02265
Saratoga	-52669	.40451	-3.40758	-62173	1.98231
Texas#5-	.17487	-32562	.00207	-1.78466	-16673
Trav-A#5	.25657	-31337	.07071	-3.95320	-58393
Trav-B#5	-03128	.15891	.20979	-23463	-1.11664
SAGA	.25292	-50498	.27041	.17586	.39708
SODA	-2.40625	-2.49517	-2.13516	.26972	-1.94207
HARDROCK	1.38931	-97400	.90076	.89768	1.35153
ARMY	92481	$-.77533$	45030	.37859	.54804
CIDERLIT	-.54780	-1.16742	.70879	.11155	.00381
COFFER	1.26150	.12453	-12784	-0.81120	1.96076
AIRPORT	$-.42642$	-1.43240	.00446	.44553	.08720
COLLESEU	.95477	-1.73154	-27094	1.54929	1.72139
LATHROP	-.55059	-1.37605	-11220	.55609	.55579

Factor Scores fal.stal-Rotation: Unrotated-Extraction: Prindpal components

 \sim

 ~ 100

 $\ddot{}$

Appendix E: ARC/INFO Macros

Macro for contouring using TIN: generates TIN contour maps for list of elements. Items within <> should be replaced with variable, value, or string of user's choice.

&do i &list <chemlcai elements, space delimited> createtin %i%tin cover location point %i% end tincontour %i%tin %i%contour <contour interval «contour base line> %i% ap display 1040 %i%contour.gra mape 460000 6150000 600000 6253000 pagesize 8 10 textfont universe mapposition cen cen textsize .14 arcs %i%contour arctext %i%contour %i% # LINE # BLANK points location pointtext location %i% quit postscript %i%contour.gra %i%contour.ps Ipr -<printemame> %i%contour.ps kill %i%contour kill %i%tin &end &do I &list «chemical elements> createtin %i%tin cover location point %i% end

Macro for contouring variance: generates variance contour maps for specified element. Can be called from contouring macro. Items within <> should be replaced with variable, value, or string of user's choice.

ap display 1040 %i%variance.gra mape location pagesize 8 11 maplimits 1.5 1 7.5 9 textfont universe mapposition cen cen textsize .14 arcs %i%variance markerset vater.mrk markerpattem 411 markercolor 5 points location arctext %i%variance %l% # LINE # BLANK linesymbol 5 box 1.5 1 7.510 line 1.5,9 7.5,9

Reproduced with permission of the copyright owner. Further reproduction prohibited without permission.

textsize .18 move 3 8.5 text %i% 'variance' quit postscript %i%variance.gra %i%variance.ps Ipr-<printemame> %i%variance.ps

Macro for contouring using kriging: generates contour maps for specified elements. Items within ∞ should be replaced with variable, value, or string of user's choice.

&do I &list CO re ti sn as sb ba cs ga ge tl ni wu mo se rb &setvar int &if %i% = re &then &s int 3 &if %i% = CO &then &s int 10 &if %!% = sn &then &s int .01 &if %i% = ti &then &s int .01 &if %i% = as &then &s int 3 &if %i% = sb &then &s int .3 &if %i% = cs &then &s int .6 &if %i% = ba &then &s int 10 &if %i% = ga &then &s int .6 &if %i% = ge &then &s int .3 &if %i% = tl &then &s int .02 &if %i% = ni &then &s int .1 &if %i% = w&then &s int .3 &if %i% = V &then &s int .6 &if %i% = u &then &s int .6 &if %i% = mo &ttien &s int 5 &if %i% = se &then &s int .1 &if %i% = rb &then &s int 2.5 kriging location %i%lattice %i%var %i% # lattice 434904.062 6043319.500 630377.25 6317028.500 15 latticecontour %i%lattice %i%contour %int% 0 %i% ap display 1040 %i%contour.gra mape location pagesize 8 11 maplimits 1.5 1 7.5 9 textfont universe medium textquality proportional mapposition cen cen textsize .14 arcs %l%contour markerset water, mrk markersymbol 412 markercolor 5 points location linecolor 1 arctext %i%contour %i% # line # blank linesymbol 5

box 1.5 1 7.5 10

line 1.5,9 7.5,9 textsize .18 move 3.5 9.5 text %i% move 3.8 9.5 text 'Concentrations «concentration units>' quit postscript %i%contour.gra %i%contour.ps Ipr -«printemame> %i%contour.ps &end

Macro for location map: generates maps for specified sample sites in the location cover. Items within <> should be replaced with variable, value, or string of user's choice.

ap display 1040 mymap.gra mape location pagesize 8 11 maplimits 1 2.7 8 9.7 mapposition cen cen markerset water, mrk markersymbol 412 markercolor 5 markersize .08 points location textfont univers medium textquality proportional textsize .10 overpost .1 # .025 overpost on overpost text moveable nodelete pointtext location location-id overpost off textsize .18 textquality proportional move 4.5 9.6 text 'Locations of Springs and Wells' cc textquality constant textsize .1 1.5 2.59 Crystal 13 Point of RxNW 25 Travertine' move 1.5 2.48
text '2 Hiko **Hiko 14 Rogers 26 Saga' 1.5 2.37 Ash 15 Scruggs 27 Soda' 1.5 2.26 Big 16 LGrapevine 28 Hardrock 1.5 2.15 Bradford 17 U.Grapevine 29 Army' 1.5 2.04 Cold 18 M.Grapevine 30 Cinderiite' 1.5 1.93 Crystal Pool 19 Mesquite 31 Coffer'**

156

 \sim \sim

 $\frac{1}{2} \left(\frac{1}{2} \right) \left(\frac{1}{2} \right) \left(\frac{1}{2} \right) = \frac{1}{2}$

 $\frac{1}{2}$. The set of $\mathcal{O}(\mathbb{R}^2)$

 \sim

 $\overline{}$

 \overline{a}

 \blacksquare and \blacksquare

 \sim \sim \sim

References:

- Bedinger, M.S.. Sargent, K.A., Langer, William H., "Studies of Geology and Hydrology in the Basin and Range Province, Southwestern United States, for Isolation of the Death Valley Region, Nevada and California," USGS Prof Paper 1370-F, 1989.
- Broxton, David E., Warren, Richard G., Byers, Frank M., Scott, Robert B., "Chemical and Mineralogic Trends Within the Timber Mountain-Oasis Valley Caldera Complex, Nevada: Evidence for Multiple Cycles of Chemical Evolution in a Long-Lived Silicic Magma System," Journal of Geoohvsical Research. Vol. 94, No. 85, pp5961-8985. May 10, 1989.
- Burbey, Thomas J., Prudic, David E., "Conceptual Evaluation of Regional Ground-Water Flow in the Carbonate-Rock Province of the Great Basin, Nevada, Utah, and Adjacent States, USGS Prof. Paper 1409-D, 1991.
- Camera and Westenberg, "Selected Ground- Water Data for Yucca Mountain Region, Southern Nevada and Eastern California, Through December 1992, USGS OFR 94-54, 1994
- Czarnecki, John B., Waddell, Richard K., "Finiete-Element Simulation of Ground-Water Flow in the Vicinity of Yucca Mountain, Nevada-California," USGS Water-Resources Investigations Report 84-4349, 1984.
- Clarke, Keith C., Analytical and Computer Cartography, Prentice Hall, NJ, 1990.
- Claassen, Hans C., Sources and Mechanisms of Recharge for Ground Water in the West-Central Amargosa Desert, Nevada- A Geochemical Interpretation, USGS Open-file Report 83-542, pp1-57, 1983.
- Dettinger, Michael D., "Distribution of Carbonate-Rock Aquifers in Southern Nevada and the Potential for their Development," Summary of Findings, 1985-1988, Summary Report No.1, 1989.

Reproduced with permission of the copyright owner. Further reproduction prohibited without permission.

- Dudley, W. W., Larson, J. D., Effect of Irrigation Pumping on Desert Pupfish Habitats in Ash Meadows, Nye County, Nevada, USGS professional Paper 927, pp1-55,1976.
- ESRI, Understanding GIS: The ARC/INFO Method. ESRI, 1990
- Greenwood & Earnshaw, Chemistry of the Elements. Pergamon Press, Cambridge, 1984.
- Hess, John W., "Hydrogeology of Regional Carbonate Aquifers in Eastern Nevada," International Contributions to Hydrogeology, Vol 13, 1992
- Hodge, Vernon, Dept of Chemistry, University of Nevada, Las Vegas, personal communication 1995.
- Hughes, Jerry L, Some Aspects of the Hydrogeology of the Spring Mountains and Pahrump Valley, Nevada, and Environs as Determined by Spring Evaluation. Mackay School of Mines, January 1966.
- Hunt, C.B., Robinson, T.W., Bowles, W.A., and Washburn, A.L., "Hydrologic Basin, Death Valley, California," USGS Professional Paper 494-B, 1966.
- Klein, Cornelis, and Hurlbut, Cornelius, Manual of Mineralogy. Wiley and Sons, 1977.
- Johannesson, Kevin H., Stetzenbach, Klaus J., Kreamer, David K., Hodge, Vernon F., "Multivariate statistical analysis of arsenic and selenium concentrations in groundwater from south-central Nevada and Death Valley, Californnia." Journal of Hydrology, Vol, 1996.
- Joreskog, K.G., Klovan, J. E., Reyment, R.A., Methods in Geomathematics 1 : Geological Factor Analysis. Elsevier Scientific Publishing Company, NY, 1976.
- Kreamer, David, Hodge, Vernon, Rabinowitz, Irving, Johannesson, Kevin, Stetzenbach, Klaus, 'Trace Element Geochemistry in Water From Selected Springs in Death Valley National Monument, California," In press, 1996
- La Camera, Richard J., Westenburg, Craig L., "Selected Groun-Water Data for Yucca Mountain Region, Southern Nevada and Eastern California, through December 1992, USGS OFR 94-54, 1994.
- Linsley, Franzini, Freyberg, Tchobanoglous, Water Resources and Environmental Engineering. McGraw-Hill. USA, 1992.
- Mack, C., Essentials of Statistics. Plenum Press, New York, 1967.
- McKinley, P.W., Long, M.P., Benson L.V., "Chemical Analyses of Water from Selected Wells and Springs in the Yucca Mountain Area, Nevada and Southeastern California, USGS OFR 90-355, 1991.
- Microsoft Excel 4.0, Microsoft Corporation, 1992-1993.

Mifflin, Martin D., Hess, John W., "Regional Carbonate Flow Systems in Nevada." Journal of Hydrology. 43, 1979, pp217-237.

- Morrison, Donald F., Multivariate Statistical Methods. McGraw-Hill, 1967.
- Norusis, Marija J., SPSS Professional Statistics 6.1. SPSS Incorporated, USA, 1994.
- Okuda, Itaru, Okazaki, Masanori, & Hahitani, Takusei, "spatial and Temporal Variations in the Chemical Weathering of Basaltic Pyroclastic Materials," Soil Science Society of America Journal. 59:887-894 1995.
- Nash, M.H, Daugherty, L.A., Whitford, W.G., "Cluster and Principal Commponent Analysis of Soil Sites Along a Transect," Soil Survey Horizons. Vol 34, no. 2, p39-60, summer 1993.
- Perfect, Diana, Creation and Analysis of a Hvdrochemical Data Base for the Death Valiev Region. Nevada and California, thesis, Colorado School of Mines, March 29, 1994.
- Perfect, Faunt, & Steinkampf, Hydrochemical Data Base for the Death Valley Region, U.S.G.S. OFR 94-XXXX, 1994 (data obtained on disk from Randy Lazniak-USGS (Water Resources Division, Las Vegas, NV).
- Peterman, Z.E., Stuckless, J.S., Mahan, S.A., Marchall, B.D., Gutentag, E.D., and Downey, J.S., "Strontium isotope characterization of the Ash Meadows ground-water system, southern Nevada, USA," Water-Rock Interaction. 1992.
- Peterman, Z.E., Stuckless, J.S., "Isotopic Evidence of Complex Ground-Water Flow at Yucca Mountain, Nevada, USA," High Level Radioactive Waste Management: Proceedings of 4th Annual International Conference, Vol 2, p. 1559-1566, 1993.
- Runnells, Donald D, Shepherd, Thomas A., Angino, Ernest E., "Metals In Water," Environmental Science and Technology. Vol 26, no 12. 1992.
- Salomons, W., Forstner, U., Metals in the Hydrocycle, Springer-Verlag, New York, 1984, pp138-209
- Schoff, Stuart L., Moore, John E., "Chemistry and Movemnet of Ground Water, Nevada Test Site," USGS Report TEI-838,1964.
- SPSS 6.1, SPSS Inc., 1994.
- Statistica 4.5, Statsoft, Inc. 1993.
- Stetzenbach, Klaus, Harry Reid Center for Environmental Studies, personal communication, 1996.
- Stetzenbach, Klaus, Amano, Megumi, Kreamer, David, Hodge, Vernon, 'Testing the Limits of ICP-MS: Determination of Trace Elements in Ground Water at the Part-Per-Trillion Level," Ground Water. Vol 32, No. 6, Nov.-Dee. 1994.
- Stevens, Calvin H., Stone, Paul, & Belasky, Paul, "Paleogeographic and structural significance of an Upper Mississippisan facies boundary in southern Nevada an deast-central California," Geological Society of America Bulletin, v. 103, p876-885, July 1991.
- Stoodley, Keith, Applied and Computational Statistics, Ellis Horwood Limited, New York, 1984.
- Thordarson, William, "Perched Ground Water in Zeolitized-Bedded Tuff, Rainier Mesa and Vicinity, Nevada Test Site," USGS Report TEI 862, 1965.
- Veizer, Jan, 'Trace Elements and Isotopes in Sedimentary Carbonates," Carbonates; Mineralogy and Chemistry, Reviews in Mineralogy. Vol 11, 1983
- White, Art, Claassen, Hans, Benson, Larry, The Effect of Dissolution of Volcanic Glass on the Water Chemistry in a Tuffaceous Aquifer, Rainier Mesa, Nevada, USGS Water Supply Paper 1535-0.
- Windom, Herbert L, Byrd, James T., Smith, Ralph G, and Huan, Feng, "Inadequacy of NASQAN Data for Assessing Metal Trends in the Nation's Rivers," Environmental Science and Technology, vol 25, 1991, pp 1137-1142.
- Winograd, Issac J., "Interbasin Movement of Ground Water at the Nevada Test Site," USGS, TEI 807, March 1962.
- Winograd, Isaac, and Thordarson, William, Hydrogeologic and Hydrochemical Framework, South-Central Great Basin, Nevada-California, with Special Reference to the Nevada Test Site, USGS Professional Paper 712-C, 1-125 pp, 1975.
- Winograd, Isaac, and Doty, Gene,Paleohydrology of the Southern Great Basin, With Special Reference to Water Table Fluctuations Beneath the Nevada Test Site During The Late (?) Pleistocene, USGS OFR 80- 569, pp 1-87, 1980.