Nitrite oxidation in geothermal springs: evidence of an upper temperature limit for thermophilic nitrite-oxidizing bacteria of 60 - 65 °C

Tara Edwards

Follow this and additional works at: https://digitalscholarship.unlv.edu/thesesdissertations

Part of the Environmental Microbiology and Microbial Ecology Commons

Repository Citation

https://digitalscholarship.unlv.edu/thesesdissertations/1729

This Thesis is brought to you for free and open access by Digital Scholarship@UNLV. It has been accepted for inclusion in UNLV Theses, Dissertations, Professional Papers, and Capstones by an authorized administrator of Digital Scholarship@UNLV. For more information, please contact digitalscholarship@unlv.edu.
NITRITE OXIDATION IN GEOTHERMAL SPRINGS: EVIDENCE OF AN UPPER TEMPERATURE LIMIT FOR THERMOPHILIC NITRITE-OXIDIZING BACTERIA OF ~60 - 65 °C

by

Tara A. Edwards

Bachelor of Science in Biology
University of Nevada, Las Vegas
2010

Bachelor of Science in Chemistry
University of Nevada, Las Vegas
2010

A thesis submitted in partial fulfillment of the requirements for the

Master of Science in Biological Sciences

School of Life Sciences
College of Sciences
The Graduate College

University of Nevada, Las Vegas
December 2012
Copyright by Tara A. Edwards, 2012

All Rights Reserved
THE GRADUATE COLLEGE

We recommend the thesis prepared under our supervision by
Tara Edwards

entitled

Nitrite Oxidation in Geothermal Springs: Evidence of an Upper Temperature Limit for
Thermophilic Nitrite-Oxidizing Bacteria of ~60-65°C

be accepted in partial fulfillment of the requirements for the degree of

Master of Science in Biological Science
School of Life Sciences
Brian Hedlund, Ph.D., Committee Chair
Penny Amy, Ph.D., Committee Member
Eduardo Robleto, Ph.D. Committee Member
Elisabeth Hausrath, Ph.D., Graduate College Representative
Tom Piechota, Ph.D., Interim Vice President for Research &
Dean of the Graduate College

December 2012
ABSTRACT

Nitrite oxidation in geothermal springs: evidence of an upper temperature limit for thermophilic nitrite-oxidizing bacteria of ~60 - 65 °C

by

Tara A. Edwards

Dr. Brian Hedlund, Committee Chair
Professor of Biology
University of Nevada, Las Vegas

Despite growing evidence of the importance of nitrification in terrestrial geothermal environments, little is known about nitrite oxidation in these environments. In order to further our knowledge, this study combined cultivation-dependent and -independent approaches with measurements of nitrogen speciation along the outflow channels of two Great Basin geothermal springs. Enrichment cultures were inoculated with sediment slurries from sites ranging in temperature from 42 – 87 °C at the sources and along the outflows of >15 hot springs. While attempts to enrich nitrite-oxidizing bacteria (NOB) from sites ≥61°C were unsuccessful, NOB were enriched from five hot springs located in U.S. Great Basin, southwestern China, and Armenia at sites <58 °C. All enrichments analyzed contained organisms with ≥97% 16S rRNA gene identity to *Nitrospira calida*, regardless of origin, demonstrating the wide geographic range of this organism. In addition, enrichments from Armenia contained organisms with ≥97% 16S rRNA gene identity to *Nitrospira moscoviensis*. Physiological properties were similar for all enrichments, with an upper temperature limit between 60 - 65 °C and a temperature optimum of 45 - 50 °C While the rates of nitrite oxidation were significantly different for the Great Basin enrichments when compared to the Tengchong and Armenian
enrichments at 50 °C (5.4 ± 2.2, 11.4 ± 3.3, and 11.6 ± 1.7 for Great Basin, Tengchong, and Armenian, respectively), the much higher rates observed for the Tengchong and Armenian enrichments may be attributed to improved cultivation conditions for the primary enrichments.

Patterns of nitrogen speciation in water samples collected along the outflow channels of two springs (Sandy’s Spring West and Rick’s Hot Creek) within the Great Boiling Spring (GBS) geothermal system suggested ammonia oxidation activity at ≥75.1 °C in both springs. In contrast, nitrite oxidation activity did not appear to be present at ≥65 °C in either spring. “Candidatus Nitrosocaldus yellowstonii” 16S rRNA gene copy numbers were abundant in sediment samples from the outflow of both springs at ≤79.6 °C, but Nitrospira 16S rRNA gene sequences were only abundant at ≤57.9 °C. Thus, an apparent difference in the upper temperature limit for ammonia oxidation and nitrite oxidation exists within the GBS system, decoupling the two steps of nitrification and leading to accumulation of nitrite above ~60 °C. In addition, 16S rRNA sequences belonging to known NOB were absent from 557,076 pyrotag sequences obtained from hot springs located in the U.S. Great Basin, and Tengchong, China at temperatures ≥55 °C and no significant matches for entire NOB genome sequences were found in the ~250 Mbp of metagenomic data from GBS environmental samples at 77 to 85 °C. In addition, no significant matches for entire NOB genome sequences were found in the 557 Mbp of metagenomic data from Yellowstone National Park environmental samples obtained from 16 springs at temperatures ranging from 52.9 to 90 °C. This study presents evidence that the upper temperature limit for nitrite oxidation in geothermal systems worldwide may be similar to the upper temperature limit observed within the GBS system and demonstrates
the wide geographic range of Nitrospira spp. in geothermal environments. Finally, we propose that the temperature-driven decoupling of ammonia oxidation and nitrite oxidation leads to a high temperature nitrite shunt in the nitrogen cycle whereby nitrite produced by ammonia oxidation is used directly by denitrifiers in geothermal ecosystems. We propose that the high temperature nitrite shunt is complete at temperatures exceeding ~65 °C, with greater flow of nitrogen through nitrate with decreasing temperature below ~65 °C.
ACKNOWLEDGEMENTS

I would like to express my deepest gratitude to my advisor, Dr. Brian Hedlund for giving me the opportunity to work in his lab and for all of his guidance throughout the last couple years. I am also grateful to my other committee members; Dr. Eduardo Robleto, Dr. Penny Amy, and Dr. Elisabeth Hausrath for all of their support and guidance throughout my graduate career. I thank Jessica Cole for being a great friend and colleague and Nicole Calica and Namritha Manoharan without whom I would have never have finished all of this work. It was an honor to work with all of the Hedlund lab members and I thank them all for their support and willingness to help. I am grateful to David and Sandy Jameson for access to the springs. I also thank Keala Kiko and Nicole Booker for all their help.

My family has selflessly supported me throughout this entire process. I am especially grateful to Jason and Jesse Gage whose sacrifices have not gone unnoticed. I would like to thank my parents for their never-ending support and encouragement. Finally I would like to thank all of my siblings for their love and encouragement.

This work was supported by U.S. National Science Foundation (Grant Numbers MCB-0546865, DBI REU 1005223, OISE-0968421, and OISE-0836450), Natural Science Foundation of China (Grant number 31070007, and Armenian National Foundation of Science and Advanced Technologies and U.S. Civilian Research & Development Foundation (Grant Number TFP-12-05).
TABLE OF CONTENTS

ABSTRACT .......................................................................................................................... III

ACKNOWLEDGEMENTS ..................................................................................................... VI

LIST OF TABLES .................................................................................................................. VIII

LIST OF FIGURES ............................................................................................................... IX

CHAPTER 1 INTRODUCTION ............................................................................................. 1

CHAPTER 2 CULTIVATION OF NITROSPIRA SPECIES FROM GEOTHERMAL SPRINGS IN CHINA, ARMENIA, AND WESTERN U.S. AT ~60 - 65 °C: IMPLICATIONS FOR GLOBAL UPPER TEMPERATURE LIMIT OF CHEMOLITHOTROPHIC NITRITE OXIDATION ........................................ 6
Abstract .............................................................................................................................. 6
Introduction .......................................................................................................................... 7
Methods ............................................................................................................................... 9
Results and Discussion ..................................................................................................... 13
Conclusions ........................................................................................................................ 17

CHAPTER 3 INTEGRATED STUDY SUGGESTS DECOUPLING OF AMMONIA OXIDATION AND NITRITE OXIDATION ABOVE ~65 °C: IMPLICATIONS FOR GLOBAL UPPER TEMPERATURE LIMIT OF CHEMOLITHOTROPHIC NITRITE OXIDATION AND PROPOSAL FOR A HIGH TEMPERATURE NITRITE SHUNT ............................................................ 26
Abstract .............................................................................................................................. 26
Methods ............................................................................................................................... 29
Results and Discussion ..................................................................................................... 32
Conclusions ........................................................................................................................ 37

CHAPTER 4 CONCLUSIONS .............................................................................................. 44

APPENDIX A ADDITIONAL TABLES .............................................................................. 47

APPENDIX B ADDITIONAL FIGURES ............................................................................ 48

REFERENCES ...................................................................................................................... 50

VITA .................................................................................................................................... 55
LIST OF TABLES

Table 2.1. NO₂⁻ oxidation activity in inoculated enrichment cultures .................. 19
Table 3.1. Locations, temperatures, and pH for each sampling site. ...................... 39
Table 3.2. OTUs from the outflows of SSW and RHC at 60°C and 50°C. ............... 40
Table A.1. Concentrations of major anions and O₂ along SSW and RHC outflow systems in July, 2010. ................................................................. 47
LIST OF FIGURES

Figure 1.1. Nitrogen biogeochemical cycle. .......................................................... 5
Figure 2.1. Stoichiometric conversion of NO$_2^-$ to NO$_3^-$ in enrichment cultures from
Great Basin, Tengchong, and Armenia ............................................................... 20
Figure 2.2. Inhibition of NO$_2^-$ oxidation activities by ClO$_3^-$ in enrichment cultures
from Great Basin, Tengchong, and Armenia ....................................................... 21
Figure 2.3. Phylogenetic tree showing the relationship between NOB in enrichment
cultures from Great Basin (GB), Tengchong (TC), and Armenia (AR) and
all cultivated Nitrospira species ...................................................................... 22
Figure 2.4. Optimization of NO$_2^-$ oxidation activity with respect to temperature in
enrichment cultures from Great Basin, Tengchong, and Armenia ................. 23
Figure 2.5. NO$_2^-$ oxidation activities at 60 °C for Great Basin and Tengchong
enrichments ......................................................................................................... 24
Figure 2.6. Optimization of NO$_2^-$ oxidation activity with respect to initial NO$_2^-$
concentration in enrichment cultures from Great Basin, Tengchong, and
Armenia .................................................................................................................. 25
Figure 3.1. Abundance of “Ca. N. yellowstonii” and Nitrospira and concentrations of
dissolved inorganic nitrogen along geothermal outflow channels of SSW
and RHC in July 2010 ......................................................................................... 41
Figure 3.2. Concentrations of dissolved inorganic nitrogen along geothermal outflow
channels of SSW and RHC in September 2012 ............................................... 42
Figure 3.3. Inferred rates of metabolic activities along the outflow of SSW and
RHC ...................................................................................................................... 43
Figure 4.1. Conceptual model of nitrogen cycle activities in GBS system springs at
temperatures: A) <65 °C  B) ~65 – 82 °C  C) >~82 °C ....................................... 46
Figure B.1. Lag of NO$_2^-$ oxidation activity with respect to temperature in enrichment
cultures from Great Basin, Tengchong, and Armenia ....................................... 48
Figure B.2. Lag of NO$_2^-$ oxidation activity with respect to initial NO$_2^-$ concentration in
enrichment cultures from Great Basin, Tengchong, and Armenia ................. 49
CHAPTER 1
INTRODUCTION

As a critical element for the synthesis of proteins, nucleic acids, and other forms of biomass, nitrogen is essential for all living organisms. In nature, nitrogen exists mainly in the following oxidation states: -3 ($\text{NH}_3$), 0 ($\text{N}_2$), +1 (N₂O), +2 (NO), +3 ($\text{NO}_2^-$), +4 (NO₂), and +5 ($\text{NO}_3^-$). Due to chemical instability and transformations by living organisms, each oxidation state is transient. Considering the major nitrogen transformations together, the processes form an integrated cycle in which the product of one process becomes the substrate for another (Figure 1.1). Since many of the reactions are catalyzed exclusively by microorganisms, the nitrogen cycle is largely dependent on these organisms. The steps of the nitrogen biogeochemical pathway can be classified as assimilatory pathways, dissimilatory pathways, or waste pathways. Assimilatory pathways, such as N₂ fixation, ammonia assimilation, and coupled nitrate/nitrite assimilation and reduction, provide reduced nitrogen that is incorporated into cell material. In contrast, dissimilatory processes such as nitrification and denitrification produce free energy from the oxidation and reduction of nitrogenous compounds. Finally, the waste process of ammonification involves the biological removal and excretion of ammonia from nitrogen-rich compounds such as proteins or amino acids.

The availability of a suitable nitrogen source often limits primary productivity in natural environments. The largest reservoir of nitrogen on Earth is N₂, which is not readily available to most organisms. By microbial nitrogen fixation, N₂ is reduced to NH₃, thereby making the nitrogen accessible to other organisms for metabolic processes.
or assimilation. NH$_3$ is also produced during the decomposition of organic nitrogen compounds in a process called mineralization or ammonification. The NH$_3$ produced is then available for nitrifying microorganisms that catalyze the oxidation of NH$_3$ to NO$_3^-$ with NO$_2^-$ as an intermediate via the two step nitrification process. During the first step of nitrification, ammonia-oxidizing microorganisms use NH$_3$ as an electron donor and O$_2$ as an electron acceptor and produce NO$_2^-$ as a waste product. In the second step of nitrification, nitrite-oxidizing bacteria use NO$_2^-$ as an electron donor and O$_2$ as an electron acceptor and produce NO$_3^-$. The waste products of nitrification (NO$_2^-$ and NO$_3^-$) may be used as respiratory electron acceptors in environments with limited oxygen availability. A diverse group of microorganisms couple the anaerobic oxidation of organic carbon to the reduction of NO$_2^-$ and NO$_3^-$ producing either NH$_4^+$ in a process known as dissimilatory nitrate reduction to ammonium (DNRA) or the gaseous products NO, N$_2$O or N$_2$ during denitrification. In addition, members of the phylum Planctomycetes couple the oxidation of ammonia (electron donor) to the reduction of NO$_2^-$ (electron acceptor) to produce N$_2$ (waste product) in anaerobic ammonia oxidation (anammox).

This cycle of coupled oxidation/reduction reactions forms an electron pool that is cycled among members of the microbial community that are often spatially separated (Reviewed in Falkowski et al., 2008). While oxidative processes are carried out in oxic environments, reductive processes are often limited to anoxic environments. In addition, while many of the organisms that catalyze the reduction of N species are heterotrophic and therefore require an environment with an available organic carbon source, autotrophic ammonia- and nitrite-oxidizing microorganisms fix their own carbon thereby
alleviating the need for available organic carbon. Furthermore, environmental constraints, such as low and high temperatures or pH, can cause spatial separation since some organisms are better equipped to deal with these constraints, as has been observed and manipulated in waste water treatment operations (Li et al., 2012).

This project focuses on nitrification. The organisms that catalyze the two steps of nitrification are distinct and, with the exception of the demonstration of a spatial-temporal co-occurrence of these organisms in some habitats (Santoro et al., 2010; Xia et al., 2011), research into the relationship between the two groups in natural environments is lacking (Hatzenpichler, 2012). Therefore, greater insight into the relationship between ammonia-oxidizing microorganisms and nitrite-oxidizing bacteria is needed. In many environments the two steps of nitrification are tightly coupled and the \( \text{NO}_2^- \) produced by ammonia oxidation is immediately consumed by nitrite oxidizers (Reviewed in Bock and Wagner, 2007). However, this coupling of the ammonia oxidation and nitrite oxidation may not be universal. For example, in waste water management low \( \text{O}_2 \) conditions are used to create a nitrite shunt where the nitrification process is shortened, ending at \( \text{NO}_2^- \) production, which is subsequently reduced by denitrifiers (Ju et al., 2012). In addition, \( \text{NO}_2^- \) accumulation reported in some geothermal springs may suggest a decoupling of ammonia and nitrite oxidation in high temperature environments (Costa et al., 2009).

Despite several indications of the importance of nitrification in terrestrial geothermal environments (Costa et al., 2009, Dodsworth et al., 2011; Dodsworth et al., 2012) and the cultivation of themophilic ammonia- and nitrite-oxidizing microorganisms (de la Torre et al., 2008; Ehrich et al., 1995; Hatzenpichler et al., 2012; Lebedeva et al., 2005; Lebedeva et al., 2008; Lebedeva et al., 2011), knowledge of nitrification, and more
specifically nitrite oxidation, in these environments is still scarce. This study utilizes several approaches such as culture-dependent and -independent methods, and geochemical analyses, to study nitrite oxidation in geothermal environments. The thermal constraints as well as the organisms involved in this second step of nitrification in high temperature environments are explored.
Figure 1.1. Nitrogen biogeochemical cycle. Yellow arrows indicate oxidative reactions, and red arrows indicate reductive reactions. White arrows indicate no change in oxidation state. Black line separates reactions that occur in oxic (top) and anoxic environments (bottom) (Madigan and Martinko, 2006).
CHAPTER 2
CULTIVATION OF *NITROSPIRA* SPECIES FROM GEOTHERMAL SPRINGS IN CHINA, ARMENIA, AND WESTERN U.S. AT ~60 - 65 °C: IMPLICATIONS FOR GLOBAL UPPER TEMPERATURE LIMIT OF CHEMOLITHOTROPHIC NITRITE OXIDATION

Abstract

Despite its importance in the nitrogen cycle, little is known about nitrite oxidation at high temperatures. In order to bridge this knowledge gap, enrichment cultures were inoculated with sediment slurries from a variety of geothermal springs. While attempts failed to enrich nitrite-oxidizing bacteria (NOB) from >10 hot spring samples at temperatures ≥61 °C, NOB were successfully enriched from the outflow of 7 hot springs located in U.S. Great Basin, southwestern China, and Armenia at temperatures ≤57.9 °C. The stoichiometric conversion of nitrite to nitrate, chlorate sensitivity, and the loss of activity after autoclaving all confirmed biological nitrite oxidation in these enrichments. Regardless of their origin, all successful enrichments contained organisms with high 16S rRNA gene sequence identity (≥97%) with *Nitrospira calida*. In addition, Armenian enrichments also contained organisms with high 16S rRNA gene sequence identity to *Nitrospira moscoviensis*. Physiological properties of all enrichments were similar, with a temperature optimum of 45 - 50 °C and an upper temperature limit between 60 °C and 65 °C. The highest rates of NOB activity occurred with initial NO$_2^-$ concentrations of 0.5 - 0.75 mM; however, lower initial nitrite concentrations resulted in shorter lag times. The results presented here suggest a possible upper temperature limit of 60 – 65 °C for
chemolithotrophic nitrite oxidation and demonstrate the wide geographic range of
*Nitrospira* spp. in geothermal environments.

**Introduction**

Nitrification, a two-step process that results in the production of nitrate from ammonia, is an important component of the nitrogen cycle. In the first step of nitrification, ammonia oxidation, ammonia is oxidized to nitrite, and in the second step, nitrite oxidation, nitrite is oxidized to nitrate. Since no known organism is capable of carrying out both steps of nitrification, it can be important to consider each step separately. The majority of research on nitrification has focused on ammonia oxidation, possibly because of the excitement surrounding the recent discovery of ammonia-oxidizing archaea (Köneke et al., 2005), the widely held assumption that ammonia oxidation is rate limiting in nature (Kowalchuk and Stephen, 2001), and the difficulty of cultivating nitrite-oxidizing bacteria (NOB) in the laboratory.

Known chemolithotrophic NOB belong to the genera *Nitrobacter*, *Nitrococcus*, *Nitrotoga*, *Nitrospina*, and *Nitrospira* (Alawi et al., 2007; Ehrich et al., 1995; Laudelout and Van Tichelen, 1960; Lebedeva et al., 2008; Lebedeva et al., 2011; Spieck et al., 2006; Watson et al., 1986; Watson and Waterbury, 1971). While *Nitrotrichlora* was traditionally considered to be an important NOB, several recent studies using cultivation-independent methods have indicated that *Nitrospira* may be more abundant than *Nitrobacter* in many environments (Altman et al., 2004; Altman et al., 2003; Hovanec et al., 1998; Schramm et al., 1999). Members of the genus *Nitrospira* represent a monophyletic group of NOB within the bacterial phylum *Nitrospira* (Spieck and Bock, 2001). They are found in a wide variety of natural habitats such as freshwater sediments
(Altman et al., 2004; Altman et al., 2003; Stein et al., 2001), soils (Attard et al., 2010; Bartosch et al., 1980; Noll et al., 2005), marine water (Watson et al., 1986), and geothermal springs (Kanokratana et al., 2004; Lebedeva et al., 2005; Lebedeva et al., 2011). Despite the tremendous ecological importance of *Nitrospira*, knowledge about this group of NOB is limited.

All *Nitrospira* are hypothesized to be K-strategists with high substrate affinity and low maximum growth rate (Huang et al., 2010; Kim and Kim, 2006; Nogueira and Melo, 2006; Schramm et al., 1999). Consistent with this hypothesis, *Nitrospira* tend to be adapted to low oxygen (Schramm et al., 1999) and nitrite availability (Nogueira and Melo, 2006; Schramm et al., 1999) and their growth can be inhibited by nitrite concentrations as low as 1.5 mM (Off et al., 2010). This property might be a contributing factor to the difficulty in cultivating *Nitrospira* in the laboratory. To date, there are only five formally identified members of this genus, with several lineages detected only by cultivation-independent censuses (Lücker et al., 2010). *Nitrospira marina* is a marine mesophile originally isolated from the Atlantic Ocean off the Gulf of Maine (Watson et al., 1986). “*Candidatus Nitrospira defluvii*” is a freshwater mesophile isolated from activated sludge (Spieck et al., 2006). *Nitrospira moscoviensis* and “*Candidatus Nitrospira bockiana*”, both isolated from heating systems, were reported to be moderately thermophilic with growth temperature optima of 39 °C and 42 °C and growth temperature ranges of 33 - 40 °C and 28 - 44 °C, respectively (Ehrich et al., 1995; Lebedeva et al., 2008). Most recently, the most thermophilic NOB isolate known, *Nitrospira calida*, was isolated from Gorjachinsk Hot Spring in the Lake Baikal area of Russia and shown to
have a growth temperature optimum of 46 - 52 °C and an upper temperature for growth of 58 °C (Lebedeva et al., 2011).

Despite this progress, and more substantial progress on chemolithotrophic ammonia oxidation (de la Torre et al., 2008; Hatzenpichler et al., 2008; Jiang et al., 2010; Pearson et al., 2004; Reigstad et al., 2008; et al., 2008a), very little is known about the oxidative nitrogen cycle at elevated temperatures. This study focuses on the enrichment and characterization of thermophilic NOB from geothermal springs in the U.S. Great Basin, southwestern China, and Armenia.

**Methods**

**Sample collection**

Sediment slurries consisting of the top ~1 cm of sediment along with spring water were collected into sterile 15 mL polypropylene tubes from the source and outflow of springs located in Great Basin, U.S., Tengchong (southwestern China), and Armenia. Samples from Great Basin springs were transported to the lab without temperature control (≤ 1 week) before being inoculated into a sterile medium (described below). Sediment slurries from Armenia were kept at 4 °C before being transported to the lab without temperature control, at which time they were inoculated into sterile media. Sediment slurries from Tengchong were inoculated into duplicate serum bottles containing the enrichment medium on site with one replicate being incubated in the spring at the collection site for 3 - 5 days before being transported to the lab and the remaining replicate being maintained without temperature control until reaching the lab.

In the lab all enrichments were incubated at the temperature of their collection site.

**Media and cultivation**
All enrichments were made using a mineral medium (modified from Ehrich et al., 1995) with the following composition: 10 mg L\(^{-1}\) CaCl\(_2\) \(\bullet\) 2H\(_2\)O, 5.8 mg L\(^{-1}\) NaHCO\(_3\), 0.5 g L\(^{-1}\) NaCl, 150 mg L\(^{-1}\) KH\(_2\)PO\(_4\) and 1 mL L\(^{-1}\) of the following stock solutions: MnCl\(_2\) \(\bullet\) 4H\(_2\)O (52 mg L\(^{-1}\)), H\(_3\)BO\(_3\) (40 mg L\(^{-1}\)), ZnSO\(_4\) \(\bullet\) 7H\(_2\)O (34 mg L\(^{-1}\)), Na\(_2\)MoO\(_4\) \(\bullet\) H\(_2\)O (34 mg L\(^{-1}\)), CuSO\(_4\) \(\bullet\) 5H\(_2\)O (25 g L\(^{-1}\)), FeSO\(_4\) \(\bullet\) 7H\(_2\)O (0.97 g L\(^{-1}\)), and MgCl\(_2\) \(\bullet\) 6H\(_2\)O (47 g L\(^{-1}\)). The medium was sterilized by autoclaving 121 °C for 60 minutes.

Enrichment cultures were incubated in stoppered serum bottles with the following total and liquid medium volumes (mL): 160/40 (v/v); 160/60 (v/v); 25/10 (v/v). Primary enrichments from Great Boiling Spring (GBS; U.S. Great Basin) at 80 °C, 65 °C, or 50 °C contained 1 mM NaNO\(_2\) and had a headspace composition of either full air or N\(_2\):air (3:1, v:v). Enrichments from Armenia and China contained 0.1 mM NaNO\(_2\), and had a headspace composition of N\(_2\):air (3:1, v:v). Maintenance cultures for all springs were incubated at 50 °C. NO\(_2^-\) was monitored weekly and replenished if the NO\(_2^-\) concentration was below 10 µM. These maintenance cultures were used to inoculate enrichments for all further studies.

Rate studies measuring the consumption of NO\(_2^-\) quantitatively were used to determine the optimal conditions for metabolic activity of nitrite oxidizers in enrichment cultures. All enrichments for rate experiments were inoculated with 1% inoculum from an enrichment maintained at 50 °C with a NO\(_2^-\) concentration of 0.3 mM. The rate study enrichments were started with 0.1 mM NaNO\(_2\). Two uninoculated negative controls and three inoculated enrichments were used for each condition in each study. In order to determine the optimal incubation temperature, enrichments were incubated at the following temperatures concomitantly: room temperature (24 - 26 °C), 40 °C, 45 °C, 50
°C, 55 °C, 60 °C, and 65 °C. Temperatures in all incubators were monitored with a traceable thermometer (VWR 46610-024) throughout the experiment and were within 1 °C of the target temperature. NO$_2^-$ concentrations were determined immediately after inoculation and weekly thereafter. NO$_2^-$ was replenished, to 0.1 mM, as needed with NaNO$_2$. After 6 (GBS and Armenian enrichments) or 8 (Tengchong enrichments) weeks, NO$_2^-$ concentrations were increased to 0.3 mM. Samples (500 µL) were taken from each enrichment and control immediately and every 3 hours thereafter for 24 hours for NO$_2^-$ concentration measurements.

The optimal initial NO$_2^-$ concentration was determined in a similar manner to the optimal temperature with the following exceptions: enrichments were prepared with 0.1 mM, 0.3 mM, 0.5 mM, 0.75 mM, 1 mM, and 1.25 mM NO$_2^-$ and incubated in the same 50 °C incubator and NO$_2^-$ concentrations were checked immediately, after 2 weeks of incubation, and weekly thereafter for 6 weeks at which time the rate study was performed as described for the temperature optimization.

**Measurement of NOB activity and inhibition**

NO$_3^-$ and NO$_2^-$ concentrations were measured colorimetrically by diazotization with and without cadmium reduction, respectively, using commercial kits (LaMotte, Chestertown, MD, USA) and a Spectronic 20D spectrophotometer (Milton Roy, USA).

Biological nitrite oxidation was verified using three methods. For all three methods, cultures were inoculated with 1% inoculum from a culture maintained at 50°C. First, the stoichiometric conversion of NO$_2^-$ to NO$_3^-$ was tested. Three enrichments and two uninoculated controls were spiked with 0.1 mM NaNO$_2$ and the conversion of NO$_2^-$ to NO$_3^-$ was monitored. When all NO$_2^-$ was depleted NO$_3^-$ was measured and enrichments
were spiked with an additional 0.1 mM NO$_2^-$.

This process was repeated until 0.9 mM NO$_2^-$ had been consumed.

Second, inhibition of NO$_2^-$ oxidation activity by the inhibitor ClO$_3^-$ was tested. Ten enrichments and six controls were incubated at 50°C for 6 weeks at which time the NO$_2^-$ concentrations were increased to 0.5 mM. Samples (500 µL) were collected from each enrichment and control every 3 hours for 30 hours and after the first 15 hours five enrichments were spiked with 5 mM NaClO$_3$. Controls included: two uninoculated bottles without NaClO$_3$, two uninoculated bottles with 5 mM NaClO$_3$, and two inoculated bottles that were spiked with 5 mM NaClO$_3$ immediately after inoculation.

Finally, to test the loss of NO$_2^-$ oxidation activity after autoclaving, three active enrichments were incubated at 50 °C for 3 weeks at which time two of the enrichments were autoclaved. NO$_2^-$ concentrations were monitored for an additional 4 weeks to ensure inactivity.

**Identification of NOB in enrichment cultures**

Cells were lysed and DNA was isolated using a FastDNA SPIN kit for Soil MP- (Biomedicals, Solon, OH) according to the manufacturer’s protocol. Isolated DNA was stored at -20 °C until analysis. DNA was amplified by PCR using primers 9bF (GRGTTTGATCCTGGCTCAG) and 1512uR (ACGGHTACCTTGTTACGACTT) (Burggraf et al., 1992; Eder et al., 1999). Each 25 µL reaction contained 5 µL of 5x Go Taq buffer (Promega, Madison, Wi, USA), 400 nM dNTP (Promega), 1 µL template DNA, 400 nM each forward and reverse primers, and 0.125 U Go Taq DNA polymerase (Promega). Cycling conditions were as follows: an initial melting step of 95 °C for 3 minutes, followed by 32 cycles of 94 °C for 30 seconds, 57 °C for 30 seconds, and 72 °C for 1
minute, and a final elongation step of 7 minutes at 72 °C. Clone libraries were made using a TOPO TA cloning kit (Invitrogen, San Diego, CA) according to the manufacturer’s protocol. Forty-four clones for each spring were sequenced by Functional Biosciences (Madison, WI) using the Sanger method. Sequences were clustered with mothur (Schloss et al., 2009), using the average neighbor algorithm at the 98% level and representative sequences for each cluster were identified using BLASTn (Altschul et al., 1990). 16S rRNA gene sequences were aligned with Nitrospira reference sequences obtained from the NCBI database using Clustal W within BioEdit (Hall, 2005) and PHYLIP was used to construct maximum likelihood and maximum parsimony trees with 100 bootstraps per tree (Felsenstein, 2005). Modified parameters in PHYLIP were as follows: outgroup was set to Thermodesulfovibrio aggregans, input order of species was randomized, and the speedier but rougher analysis was turned off.

Results and Discussion

Results of thermophilic NOB enrichments

Initially, thermophilic NOB enrichments were inoculated with 50 °C, 65 °C, and 80 °C sediments from GBS and incubated at in situ temperature with 1 mM NO$_2^-$ and a fully aerobic or sub-oxic atmosphere (N$_2$:air, 3:1, v:v). The 65 °C and 80°C enrichments remained inactive throughout the ≥1 year of incubation, but 50 °C enrichments began to show activity after 4 months of incubation. Similar lag times were observed for enrichments regardless of headspace gas composition, but in subsequent enrichments growth was more reproducible under microaerophilic conditions. In addition, when NO$_2^-$ concentrations were reduced to 0.1 mM, lag times were reduced to ≤2 weeks. Based on
the results from GBS enrichments, all additional enrichments were microaerophilic, incubated at 50 °C, and with 0.1 mM initial NO$_2^-$ concentration.

Using the improved NOB enrichment procedure, samples used for inocula were obtained from a variety of geothermal features in Tengchong County, Yunnan Province, China, and Armenia. Samples from Tengchong China were collected from various locations within the Rehai (“Hot Sea”) geothermal field, which is a high temperature, granite-hosted system with a large diversity of geothermal features and a diverse population of thermophiles (reviewed in Hedlund et al., 2012), and the carbonate-hosted Ruidian system (Meixiang & Wei, 1987; Zhang et al., 2008b). Samples from Armenia were collected from a variety of circumneutral, carbonate-buffered springs (Mkrtchyan, 1969). While all enrichments inoculated from sites ≥61 °C were unsuccessful, enrichments from ≤53 °C samples were all successful, with the exception of one set inoculated from an acidic spring (pH 3) in Rehai and one from a spring in Armenia (Table 2.1). For successful enrichments, NOB activity became apparent within 2 weeks of inoculation.

Enrichments from the outflows of GBS (50 °C), Qiaobianrequan (48.5 °C), and Jermuk Spring (53 °C) were chosen as representative cultures from the U.S. Great Basin, China, and Armenia, respectively, for further study. Once enrichments were stable, two different approaches were employed to try to obtain axenic NOB cultures. First, optical tweezers were used to inoculate single cells from GBS and Tengchong enrichments (~90 cells total) into the same medium, autoclaved spent enrichment medium, or filtered spring water amended with 0.1 mM NO$_2^-$. Second, serial dilutions were used to inoculate fresh medium. To date, neither method has successfully produced pure cultures. However,
NOB are known to be extremely difficult to isolate and, to date, only three species of *Nitrospira* have been formally described as axenic species (Ehrich *et al.*, 1995; Lebedeva *et al.*, 2011; Watson *et al.*, 1986). Additionally, “*Candidatus Nitrospira bockiana*” took up to 12 years to isolate and has not yet been described formally as an axenic culture (Lebedeva *et al.*, 2008). Given the difficulty to obtain thermophilic NOB pure cultures, and recognizing both the opportunities and limitations of studying mixed microbial cultures, enrichment cultures were used to determine the optimal conditions for nitrite oxidation activity.

**Evidence of biological nitrite oxidation**

The predicted accumulation of NO$_3^-$ with the depletion of NO$_2^-$, according to the reaction NO$_2^-$ + ½ O$_2$ → NO$_3^-$, was used to verify NO$_2^-$ oxidation in enrichment cultures. The approximate 1:1 stoichiometric conversion of NO$_2^-$ to NO$_3^-$ confirmed NO$_2^-$ oxidation activity in all enrichments tested (Figure 2.1). In addition, since ClO$_3^-$ is a specific inhibitor of biological NO$_2^-$ oxidation (Belser and Mays, 1980), it was used to further confirm NOB activity. Similar results were observed for all enrichments, regardless of origin. Five mM ClO$_3^-$ completely inhibited NOB activity in freshly inoculated subcultures and amendments of 5 mM ClO$_3^-$ stopped NOB activity in active enrichments within five hours (Figure 2.2). Finally, autoclaving long-term enrichments led to cessation of nitrite oxidation activity.

**Composition of enrichment cultures**

Clone libraries revealed the presence of organisms with high (95 - 99%) 16S rRNA gene identity to *Nitrospira calida* in enrichment cultures from all three springs. 16S rRNA gene sequences for other known NOB were not found in the GBS or
Tengchong enrichments; however, the Armenian enrichments also contained close relatives of *Nitrospira moscoviensis* (Figure 2.3). The GBS enrichment contained four additional species-level groups including close relatives of *Meiothermus timidus* (99% identity), one unidentified delta proteobacterium with 84% identity to *Geobacter hephaestius*, and two different Betaproteobacteria, one with 91% identity to *Azospira restricta* and another with 89% identity to *Petrobacter succinatimandens*. The Tengchong enrichment was much less diverse with 39 of the 44 sequences identified as close relatives of *N. calida* and the remaining five sharing 93% identity with *Anoxybacillus amylolyticus*. The Armenian enrichment contained five additional species-level groups including relatives of *Anoxybacillus contaminans* (99% identity), *Ignavibacterium album* (98% identity), *Meiothermus timidus* (99% identity), *Derxia gummosa* (92% identity), and an organism with 83% identity to “*Candidatus Chloracidobacterium thermophilum*”

**Physiological properties**

Incubation temperatures of 45 °C and 50 °C resulted in the fastest rate of NO$_2^-$ consumption for all enrichments (Figure 2.4). Similarly, temperatures of 40 - 45 °C for GBS enrichments and 50 °C for both Tengchong and Jermuk enrichments resulted in significantly shorter lags than other temperatures tested (Figure B1). The temperature range for NOB activity and the rate of nitrite oxidation appeared to be significantly different for the Great Basin enrichments compared to the other two enrichments during the optimal temperature experiments. However, this difference may be due to the use of improved NOB enrichment procedures for the primary Tengchong and Armenian enrichments or to differences in the other inhabitants of the co-culture. NOB activity
occurred at 60 °C in enrichments from all three springs (Figure 2.4 and 2.5); but, despite several attempts to grow enrichments from all three locations at temperatures ≥65 °C, NO\textsubscript{2}\textsuperscript{−} oxidation activity was never observed at these temperatures. Growth at 60 °C was difficult to reproduce in all enrichments (e.g., see Figure 2.4), despite careful temperature monitoring and all attempts to maintain reproducibility. Although not conclusive, these results are in agreement with other studies (Lebedeva et al., 2005; Lebedeva et al., 2011) and suggest an upper temperature limit of 60 to 65 °C for biological nitrite oxidation.

Despite the fact that the temperature range for Nitrospira species can be dependent on NO\textsubscript{2}\textsuperscript{−} concentration (Lebedeva et al., 2008), we did not test this dependency in our enrichments. However, we did use low (0.3 mM) NO\textsubscript{2}\textsuperscript{−} concentrations and Lebedeva et al. reported that low NO\textsubscript{2}\textsuperscript{−} concentrations (0.3 vs. 1.5 mM) resulted in a wider temperature range for “Candidatus Nitrospira bockiana” (Lebedeva et al., 2008). In addition, the temperature range did not change for Nitrospira calida whether 0.3 or 2.5 mM NO\textsubscript{2}\textsuperscript{−} was used (Lebedeva et al., 2011). While the effect of NO\textsubscript{2}\textsuperscript{−} concentrations on the temperature ranges was not explored, we did determine the effect of initial NO\textsubscript{2}\textsuperscript{−} concentration on NOB activity. The experiments revealed that low concentrations of NO\textsubscript{2}\textsuperscript{−} (0.1 - 0.3 mM) resulted in the shortest lag time for all enrichments (Figure B2). On the other hand, initial NO\textsubscript{2}\textsuperscript{−} concentrations of 0.5 to 0.75 mM resulted in the fastest rate of nitrite oxidation (Figure 2.6). These results emphasize the importance of using low concentrations of NO\textsubscript{2}\textsuperscript{−} when cultivating thermophilic Nitrospira and provide a plausible explanation for the lengthy initial lag time for GBS enrichments.

**Conclusions**
As evidence of their wide geographic range, *Nitrospira* species were found in enrichment cultures from geothermal features located in Great Basin, U.S., Tengchong, China, and Armenia. The sensitivity of these enrichments to both oxygen and nitrite concentrations supports the hypothesis that *Nitrospira calida* are K-strategists and emphasizes the need for low NO$_2^\text{-}$ concentration (<1 mM) when cultivating thermophilic *Nitrospira*. Similar to previously published results (Lebedeva et al., 2005, Lebedeva et al., 2011), the upper temperature limit for growth for the nitrite-oxidizing enrichments, regardless of origin, was between 60 - 65 °C. Together the data presented here suggests an upper temperature limit for chemolithotrophic nitrite oxidation of 60 to 65 °C.
Table 2.1. NO$_2^-$ oxidation activity in inoculated enrichment cultures.

<table>
<thead>
<tr>
<th>Location</th>
<th>Spring</th>
<th>GPS location at source</th>
<th>pH</th>
<th>Temperature (°C)</th>
<th>NO$_2^-$ oxidation activity</th>
</tr>
</thead>
<tbody>
<tr>
<td><strong>Great Basin, U.S.</strong></td>
<td></td>
<td></td>
<td></td>
<td></td>
<td></td>
</tr>
<tr>
<td></td>
<td>GBS</td>
<td>N40.84139° W119.61889°</td>
<td>6.80</td>
<td>80.0</td>
<td>-</td>
</tr>
<tr>
<td></td>
<td></td>
<td></td>
<td>7.03</td>
<td>65.0</td>
<td>-</td>
</tr>
<tr>
<td></td>
<td></td>
<td></td>
<td>7.26</td>
<td>50.0</td>
<td>+</td>
</tr>
<tr>
<td><strong>Tengchong, China</strong></td>
<td></td>
<td></td>
<td></td>
<td></td>
<td></td>
</tr>
<tr>
<td></td>
<td>Zhenzhuquan</td>
<td>N24.95115° E98.43596°</td>
<td>4.30</td>
<td>91.7</td>
<td>-</td>
</tr>
<tr>
<td></td>
<td></td>
<td></td>
<td>3.20</td>
<td>87.8</td>
<td>-</td>
</tr>
<tr>
<td></td>
<td></td>
<td></td>
<td>2.77</td>
<td>74.0</td>
<td>-</td>
</tr>
<tr>
<td></td>
<td></td>
<td></td>
<td>2.81</td>
<td>66.6</td>
<td>-</td>
</tr>
<tr>
<td></td>
<td></td>
<td></td>
<td>3.09</td>
<td>53.0</td>
<td>-</td>
</tr>
<tr>
<td></td>
<td>Diretiyanqu</td>
<td>N24.95396° E98.43829°</td>
<td></td>
<td></td>
<td></td>
</tr>
<tr>
<td></td>
<td></td>
<td></td>
<td>8.29</td>
<td>83.0</td>
<td>-</td>
</tr>
<tr>
<td></td>
<td></td>
<td></td>
<td>8.33</td>
<td>74.3</td>
<td>-</td>
</tr>
<tr>
<td></td>
<td></td>
<td></td>
<td>8.39</td>
<td>69.4</td>
<td>-</td>
</tr>
<tr>
<td></td>
<td></td>
<td></td>
<td>8.56</td>
<td>61.2</td>
<td>-</td>
</tr>
<tr>
<td></td>
<td>Zimeiquan</td>
<td>n.d.</td>
<td>8.98</td>
<td>84.7</td>
<td>-</td>
</tr>
<tr>
<td></td>
<td>Gumingquan</td>
<td>N24.57060° E98.43615°</td>
<td>9.40</td>
<td>83.5</td>
<td>-</td>
</tr>
<tr>
<td></td>
<td></td>
<td></td>
<td>8.29</td>
<td>83.0</td>
<td>-</td>
</tr>
<tr>
<td></td>
<td>Direchi</td>
<td>N24.95009° E98.43807°</td>
<td></td>
<td></td>
<td></td>
</tr>
<tr>
<td></td>
<td></td>
<td></td>
<td>8.29</td>
<td>83.0</td>
<td>-</td>
</tr>
<tr>
<td></td>
<td></td>
<td></td>
<td>8.33</td>
<td>74.3</td>
<td>-</td>
</tr>
<tr>
<td></td>
<td></td>
<td></td>
<td>8.39</td>
<td>69.4</td>
<td>-</td>
</tr>
<tr>
<td></td>
<td></td>
<td></td>
<td>8.56</td>
<td>61.2</td>
<td>-</td>
</tr>
<tr>
<td></td>
<td>Jinze</td>
<td>N23.44418° E98.46004°</td>
<td>6.71</td>
<td>80.6</td>
<td>-</td>
</tr>
<tr>
<td></td>
<td>Gongxiaoshe</td>
<td>N25.44012° E98.44081°</td>
<td>7.29</td>
<td>73.8</td>
<td>-</td>
</tr>
<tr>
<td></td>
<td>Shuirebaozha</td>
<td>N24.95002° E98.43728°</td>
<td>8.27</td>
<td>72.1</td>
<td>-</td>
</tr>
<tr>
<td></td>
<td>Qiaobianrequan</td>
<td>N24.95044° E98.43650°</td>
<td>~7</td>
<td>57.9</td>
<td>+</td>
</tr>
<tr>
<td></td>
<td></td>
<td></td>
<td>~7</td>
<td>48.5</td>
<td>+</td>
</tr>
<tr>
<td></td>
<td>Sinter Apron</td>
<td>n.d.</td>
<td>9.00</td>
<td>33.6</td>
<td>+</td>
</tr>
<tr>
<td><strong>Armenia</strong></td>
<td></td>
<td></td>
<td></td>
<td></td>
<td></td>
</tr>
<tr>
<td></td>
<td>Karvachar</td>
<td>N40.17417° E46.27500°</td>
<td>7.30</td>
<td>56.0</td>
<td>+</td>
</tr>
<tr>
<td></td>
<td>Jermuk</td>
<td>N39.96639° E45.68528°</td>
<td>7.05</td>
<td>53.0</td>
<td>+</td>
</tr>
<tr>
<td></td>
<td>Jermuk G</td>
<td>N39.87944° E45.77417°</td>
<td>6.90</td>
<td>53.0</td>
<td>+</td>
</tr>
<tr>
<td></td>
<td>Arzakan</td>
<td>N40.68389° E44.74111</td>
<td>7.20</td>
<td>44.0</td>
<td>+</td>
</tr>
<tr>
<td></td>
<td>Hankavan</td>
<td>N40.63265° E44.48463°</td>
<td>7.00</td>
<td>44.0</td>
<td>-</td>
</tr>
</tbody>
</table>

Figure 2.1. Stoichiometric conversion of NO$_2$\textsuperscript{-} to NO$_3$\textsuperscript{-} in enrichment cultures from Great Basin, Tengchong, and Armenia. NO$_2$\textsuperscript{-} and NO$_3$\textsuperscript{-} were monitored colorimetrically. Errors bars, standard deviation (n=3).
Figure 2.2. Inhibition of NO₂⁻ oxidation activities by ClO₃⁻ in enrichment cultures from Great Basin, Tengchong, and Armenia. Controls included on the graph are the uninoculated controls with ClO₃⁻ added. Arrows indicate addition of 5 mM ClO₃⁻. Errors bars, standard deviation (n=5).
**Figure 2.3.** Phylogenetic tree showing the relationship between NOB in enrichment cultures from Great Basin (GB), Tengchong (TC), and Armenia (AR) and all cultivated Nitrospira species. 16S rRNA gene sequences were clustered using mothur at the 98% level. Representative OTUs are shown along with the number of clones represented in parenthesis. Accession numbers and strain designations are included for cultivated *Nitrospira* species. Bootstrap values >80% from maximum likelihood (ML) and maximum parsimony (MP) are shown.
Figure 2.4. Optimization of $\text{NO}_2^-$ oxidation activity with respect to temperature in enrichment cultures from Great Basin, Tengchong, and Armenia. $\text{NO}_2^-$ and $\text{NO}_3^-$ were monitored colorimetrically. Errors bars, standard deviation (n=3). * Samples incubated at room temperature (24 °C to 26 °C).
Figure 2.5. NO$_2^-$ oxidation activities at 60 °C for Great Basin and Tengchong enrichments. NO$_2^-$ and NO$_3^-$ were monitored colorimetrically. Errors bars, standard deviation (n=3).
Figure 2.6. Optimization of NO$_2^-$ oxidation activity with respect to initial NO$_2^-$ concentration in enrichment cultures from Great Basin, Tengchong, and Armenia. NO$_2^-$ and NO$_3^-$ were monitored colorimetrically. Errors bars, standard deviation (n=3).
CHAPTER 3
INTEGRATED STUDY SUGGESTS DECOUPLING OF AMMONIA OXIDATION
AND NITRITE OXIDATION ABOVE ~65 °C: IMPLICATIONS FOR GLOBAL
UPPER TEMPERATURE LIMIT OF CHEMOLITHOTROPHIC NITRITE
OXIDATION AND PROPOSAL FOR A HIGH TEMPERATURE NITRITE SHUNT

Abstract
A significant difference between the upper temperature limits for ammonia and nitrite
oxidation has been suggested by previously published results, but has not been explicitly
investigated. To investigate the effect of temperature on nitrification, the abundance and
inferred activity of nitrifying microorganisms was determined along the outflow channels
of two Great Basin hot springs. Concentrations of dissolved inorganic nitrogen from the
source to ~50 °C showed apparent activity of ammonia-oxidizing archaea (AOA) at
temperatures >75.2 °C and 80.3 °C for Sandy’s Spring West (SSW) and Rick’s Hot
Creek (RHC), respectively, and an apparent upper temperature limit for nitrite oxidati
activity of ~60 - 65 °C in both springs. Correspondingly, 16S rRNA gene copy numbers
of the AOA “Candidatus Nitrosocaldus yellowstonii” were abundant at temperatures
≤79.6 °C and Nitrospira 16S rRNA gene sequences were only abundant at ≤57.9 °C.
Clone libraries confirmed the specificity of primers used and showed a predominance of
organisms with high (≥95%) 16S rRNA gene identity to either Nitrospira calida or
Nitrospira moscoviensis. In addition, 16S rRNA sequences belonging to known NOB
were absent from 557,076 pyrotag sequences obtained from hot springs located in the
U.S. Great Basin and Tengchong, China at temperatures ≥55 °C. Furthermore, significant
matches with the entire genomes of known NOB were absent from the ~250 Mbp of
metagenomic data obtained from GBS environmental samples at temperatures ranging from 77 to 85 °C, and the 557 Mbp of metagenomic data from Yellowstone National Park environmental samples at temperatures ranging from 52.9 to 90 °C. The temperature-driven decoupling of ammonia oxidation and nitrite oxidation leads to a high temperature nitrite shunt in the nitrogen cycle whereby nitrite produced by ammonia oxidation is used directly by denitrifiers in geothermal ecosystems. We propose that the high temperature nitrite shunt is complete at temperatures exceeding ~65 °C, with greater flow of nitrogen through nitrate with decreasing temperature below ~65 °C.

**Introduction**

Nitrification, the sequential oxidation of ammonia to nitrate via nitrite, is a key process in the nitrogen cycle. The two steps of nitrification, ammonia oxidation and nitrite oxidation, are carried out by distinct organisms. Aerobic ammonia oxidation, the first step of nitrification, was traditionally thought to be restricted to ammonia-oxidizing bacteria within the phylum Proteobacteria (Kowalchuk and Stephen, 2001). However, ammonia-oxidizing archaea (AOA) were recently discovered within the phylum Thaumarchaeota (Könneke et al., 2005). The second step of nitrification, nitrite oxidation, is catalyzed by nitrite-oxidizing bacteria (NOB). All cultured chemolithoautotrophic NOB belong to either the phylum Nitrospira or the alpha (*Nitrobacter*), beta (*Nitrospina*), delta (*Nitrospina*) or gamma (*Nitrococcus*) classes of the phylum Proteobacteria (Alawi et al., 2007; Ehrich et al., 1995; Laudelout and Van Tichelen, 1960; Lebedeva et al., 2008; Lebedeva et al., 2011; Spieck et al., 2006; Watson et al., 1986; Watson and Waterbury, 1971). Both ammonia- and nitrite-oxidizing microorganisms have been reported in a variety of habitats, including: soil (Bartosch et
Despite the wide distribution of nitrifying organisms in geothermal environments (Costa et al., 2009; Dodsworth et al., 2011; Reigstad et al., 2008), multiple lines of evidence suggest that the two steps of nitrification have different upper temperature limits. Ammonia oxidation has been demonstrated in co-culture at 74 °C by the thermophilic archaeon “Candidatus Nitrosocaldus yellowstonii” (de la Torre et al., 2008), and biomarkers for AOA have been recovered from environments up to 97 °C (Reigstad et al., 2008). In contrast, the most thermophilic NOB known, Nitrospira calida, is only active in pure culture up to 50 °C (Costa et al., 2009) and in enrichment culture up to 60 °C (Lebedeva et al., 2005). In addition, biomarkers for NOB have not been reported above 69 °C. One report has suggested that nitrite oxidation was active at 85 °C in two geothermal springs in Iceland based on the $^{15}$NO$_3^-$ pool dilution approach; however, that study did not distinguish between ammonia oxidation and nitrite oxidation because $^{15}$NO$_2^-$ and $^{15}$NO$_3^-$ were not distinguished (Reigstad et al., 2008).

Geochemical and molecular studies of the Great Boiling Spring (GBS) geothermal system in the U.S. Great Basin have highlighted the importance of nitrification in this system. The presence of NH$_3$ as the dominant form of dissolved inorganic nitrogen at the source of several GBS springs supports a highly active nitrogen cycle in which NH$_3$ is oxidized to NO$_2^-$ at temperatures of ~80 °C (Dodsworth et al.,

al., 2002; Nicol et al., 2008; Zhang et al., 2010), freshwater (Altmann et al., 2004; Altmann et al., 2003; Cebron et al., 2003; Jiang et al., 2009), marine (Köhneke et al., 2005; Watson et al., 1986; Watson and Waterbury, 1971), and geothermal (Costa et al., 2009; de la Torre et al., 2008; Dodsworth et al., 2011; Hatzenpichler et al., 2012; Hirayama et al., 2005; Zhang et al., 2008a) environments.
2011). Furthermore, several studies reported on the abundance of “Ca. N. yellowstonii” and close relatives at <82 °C within the GBS geothermal system (Cole et al., 2012; Costa et al., 2009; Dodsworth et al., 2011). While DNRA and denitrification are both limited by ammonia oxidation in GBS (Dodsworth et al., 2011; Hedlund et al., 2001), metabolic interactions between microorganisms important for nitrogen cycling have not been assessed. In particular, the process of NO₂⁻ oxidation has been unexplored.

In this study we investigated the effect of temperature on the abundance and inferred activity of both AOA and NOB along outflow channels within the GBS geothermal system. Rates of ammonia oxidation and nitrite oxidation were calculated based on volumetric flow rates and changes in dissolved inorganic nitrogen concentration along the spring outflows. These rates were correlated with “Ca. N. yellowstonii” and Nitrospira abundance in sediment samples collected at sites along the outflow channels of two GBS hot springs.

Methods

Collection sites and samples

Sample collection was conducted in July of 2010 and September of 2012 from the source and outflow of Sandy’s Spring West (SSW) and Rick’s Hot Creek (RHC) which are part of the GBS geothermal system and have been described previously (Costa et al., 2009; Dodsworth et al., 2011; Huang et al., 2007). Temperature and pH for each sample site was measured using a handheld pH 5 meter (LaMotte, Chestertown, MD). Bulk water samples collected for inorganic chemical measurements were filtered using a 0.2 µm IC Acrodisc filter (Pall Life Sciences) and transported on wet ice to the lab. Major anions concentrations were analyzed using ion chromatography (Table A1). Sediment samples
collected for quantitative polymerase chain reaction (qPCR) were immediately frozen on dry ice, transported back to the lab, and stored at -80 °C until analysis.

Flow velocity was determined by the time it took for a neutrally buoyant 1 in. diameter rubber ball to travel a measured distance in the spring. The cross-sectional area was measured at three and four sites along the outflow channels of RHC and SSW, respectively. Volumetric flow was determined by the product of the mean flow velocity and cross-sectional area.

**Dissolved inorganic nitrogen concentrations and rates of metabolic activity**

∑NH$_3$ was measured by the Nesslerization method and NO$_2^-$ and NO$_3^-$ were measured by diazotization with and without cadmium reduction, respectively, using either a Lachat QuikChem 8000 FIA nutrient analyzer (Zellweger Analytics, Milwaukee, WI) or a Spectronic 20D spectrophotometer (Milton Roy, Rochester, NY). Rates of metabolic activity were calculated using the volumetric flow rate and changes in dissolved inorganic nitrogen (NO$_2^-$ + NO$_3^-$ for ammonia oxidation and NO$_3^-$ for nitrite oxidation) between each site.

**Quantifying AOA and NOB**

Qualitative PCR was performed on template DNA extracted from SSW and RHC using primer set CNY16S-F and -R (TAGCTGAAATCTATATGGCCC, ATTCTCCAGCCT TTTTACAGC), which was specific for the 16S rRNA gene of “Ca. N. yellowstonii” and close relatives (Costa *et al.*, 2009), and two primer sets specific for NOB. NOB-specific primers sets included: F1-norA and R2-norA (CAGACCGACGTGTCGAAAAG, CAGACCGACGTGTCGAAAAG), which were specific for the 16S rRNA gene of *Nitrobacter* (Köneke *et al.*, 2005), and NSP8-F and -
R (CGGCAGTCCCCTCGGACCTTT, ATGGGACGGGAAACCGTTCGGA), which were specific for the *Nitrospira* 16S rRNA gene (present study).

Quantitative PCR (qPCR) was performed on template DNA extracted from SSW and RHC using primer sets CNY16S-F and –R and NSP8-F and –R. The amplified regions were ~250 nt for the CNY16S primer set and ~189 nt for the NSP8 primer set. Standard curves were produced by isolating plasmids (pCR2.1-TOPO) containing the target sequence using a QIAprep Spin MiniPrep Kit (Qiagen). Plasmids were then quantified using a Nanodrop 1000 (Thermo Scientific, Waltham, MA) and diluted in 10-fold dilutions ranging from \(~10^2\) - \(10^9\) copies/reaction. Plasmid contained either the clone SSE_L4_B03 (for primer set CNY16S-F and –R) (Costa *et al.*, 2009) or amplified *Nitrospira* 16S rRNA gene sequences from GBS (for primer set NSP 8-F and –R). All standard curve reactions were prepared in duplicate.

Sample reactions were prepared in triplicate and coupled with negative controls (no template). Reactions (25 µL) were prepared in individual wells of an iQ 96-well PCR plate (BioRad) and contained 12.5 µL of 2X PerfeCTa SYBR Green SuperMix for iQ (Quanta Biosciences, Gaithersburg, MD), 400 nM each primer, and 5 µL template DNA. The following cycling conditions were used: an initial melt cycle (95 °C for 3 min) followed by 45 cycles of melting (94 °C for 15 s), annealing (57 °C for 25 s) and extension (72 °C for 45 s), with data collection using a SYBR-490 filter enabled during the 72 °C step, followed by a melt curve 55 °C to 95 °C by 0.5 °C increments with 10 s at each step. Gene copy numbers were quantified and threshold cycles and data analysis were done using an iCycler iQ Multicolor Real-Time PCR Detection System with iCycler iQ Optical System Software v3.1 (BioRad, Hercules, CA). Standard curves were log-
linear and correlation coefficients ($r^2$) for regressions and amplification efficiencies ranged from 104% to 99.3% and from 0.890 to 0.997, respectively.

**Identification of NOB**

Cells were lysed and DNA was isolated using a FastDNA SPIN kit for Soil (MP Biomedicals, Solon, OH). NOB 16S rRNA genes were amplified using the same primer sets as used in qPCR. Cloning was done with a TOPO TA cloning kit (Invitrogen, San Diego, CA), and NOB-specific clones were sent to Functional Biosciences (Madison, WI) for sequencing (Sanger method). The resulting sequences were clustered at the 97% OTU level using the program mothur v1.24.1 (Schloss et al., 2009) and the representative sequences were coarsely identified using BLASTn (Altschul et al., 1997).

**Metagenome analysis**

MUMmer v3.0 was used to evaluate the presence of known NOB genome sequences in metagenome data from GBS and Yellowstone National Park. The ~250 Mbp of GBS metagenomic data and the 557 Mbp of Yellowstone metagenomic data were queried using the default setting for MUMmer (Kurtz et al., 2004). The following reference sequences were obtained from the NCBI database: “Candidatus Nitrospira defluvii” NC_014355; *Nitrobacter winogradskyi Nb_255*, NC_007406; and *Nitrobacter hamburgensis* X14, NC_007964 (Pruit et al., 2005) and metagenome data for GBS and Yellowstone were obtained from the Department of Energy’s Joint Genome Institute (Grigoriev et al., 2012).

**Results and Discussion**

Dissolved inorganic nitrogen concentrations and calculated rates N-transformation
In order to assess zones of metabolic activity and inferred rates of nitrification, bulk water samples were collected in July of 2010 and September of 2012 at the source and at intervals throughout the outflow SSW and RHC. Sample temperature, pH, GPS location, and distance from spring source are recorded in Table 3.1. Selected ionic concentrations for samples collected during July of 2010 are shown in Table S1. In September of 2012 the volumetric flow rates were 56 ± 9.1 L s\(^{-1}\) and 157 ± 24 L s\(^{-1}\) for SSW and RHC, respectively.

Spatial patterns of dissolved inorganic nitrogen concentrations along the spring outflows were similar for both springs on both sampling dates (Figures 3.1 and 3.2). As previously noted (Dodsworth et al., 2011), the source waters of both SSW and RHC contained moderate concentrations of ∑NH\(_3\) (59 - 97 µM) and low concentrations of NO\(_2^-\) and NO\(_3^-\) (<2 µM) (Figures 3.1 and 3.2). Decreasing concentrations of ∑NH\(_3\) concomitant with increasing NO\(_2^-\) + NO\(_3^-\) strongly suggested nitrifying activity along the outflows of both springs. However, ammonia and nitrite oxidation were spatially disconnected, and appeared to be defined by different temperature ranges.

Decreasing ∑NH\(_3\) concomitant with increasing NO\(_2^-\) strongly suggested ammonia oxidation activity at temperatures >80.3 °C. Estimated rates of ammonia oxidation were highest between ~80 and 75 °C in both springs (0.61 ± 0.10 mol cm\(^{-1}\) s\(^{-1}\) and 0.30 ± 0.05 mol cm\(^{-1}\) s\(^{-1}\) SSW RHC) and were not significant at ≥84.6 °C in either spring (Figure 3.3). Thus, the optimal temperature for ammonia oxidation appears to be between ~80 and 75 °C within the GBS geothermal system. In addition, the thermal limit of ammonia oxidation within this system appears to be between 80.3 to 84.6 °C, which is consistent with the high abundance of “Ca. Nitrosocaldus” in GBS at sites ≤ 82 °C on multiple
sampling dates (Cole et al., 2012; Costa et al., 2009; Dodsworth et al., 2011) and measurement of relatively high rates of ammonia oxidation in GBS and SSW at 82 °C and 81 °C (Dodsworth et al., 2011).

In contrast, the accumulation of NO$_2^-$ along with low NO$_3^-$ (<1 µM) strongly suggested a lack of nitrite oxidation at temperatures >65 °C. Inferred rates of nitrite oxidation were highest between ~55 and 50 °C (0.73 ± 0.030 mol cm$^{-1}$ s$^{-1}$ in SSW and 0.68 ± 0.060 mol cm$^{-1}$ s$^{-1}$ in RHC) and were not significant at temperatures >65 °C in either spring. These results suggest an upper temperature limit between ~61 - 65 °C for nitrite oxidation within the GBS geothermal system, which differs significantly (~15 - 20 °C) from the upper temperature limit of inferred ammonia oxidation activity. The apparent thermal decoupling between the two steps of nitrification within GBS geothermal system provides a plausible explanation for the accumulation of NO$_2^-$ reported previously in GBS at temperatures >65 °C (Costa et al., 2009).

Many factors, such as pH and the concentration of dissolved O$_2$ (DO) covary with temperature in geothermal systems (Fouke, 2011) providing possible alternative explanations for the spatial separation observed between nitrite oxidation and ammonia oxidation. However, with the exception of the source of SSW, the pH for all sites fell within the reported range supporting growth of Nitrospira calida (pH 7.0 - 8.8), which has been cultivated previously from this site (Table 3.1). In addition, while NOB are more sensitive to limited O$_2$ than AOA (Ju et al., 2007), NOB have been shown to persist in environments with DO levels as low as 0.8 ppm (Fussel et al., 2012), which is well below the DO concentrations at sites ≤79.8 °C in both springs (Table A1). While other
unmeasured physicochemical variables cannot be ruled out, temperature remains a likely factor driving ammonia and nitrite oxidation within the GBS geothermal system.

**Quantification of AOA and NOB**

Since “Ca. N. yellowstonii” and its close relatives are the only members of the AOA known to be present within the GBS geothermal system (Costa et al., 2009; Dodsworth et al., 2011), AOA abundance was determined by qPCR using “Ca. N. yellowstonii”-specific 16S rRNA gene primers. Gene copy numbers for “Ca. N. yellowstonii”-like 16S rRNA were most abundant at 71.9 °C in SSW and 79.6 °C in RHC and were not abundant (≥10^6) at temperatures ≥84.4 °C in either spring (Figure 3.1). These results delineate an upper temperature limit for “Ca. Nitrosocaldus yellowstonii” between 79.6 °C and 84.6 °C within the GBS geothermal system, which is in agreement with the reported lack of abundant “Ca. Nitrosocaldus yellowstonii” at 86.6 °C in SSW (Costa et al., 2009) and at temperatures >82 °C in GBS (Cole et al., 2012). In addition, the upper temperature limit of “Ca. Nitrosocaldus yellowstonii” is in agreement with the upper temperature limit for inferred ammonia oxidation activity suggesting that “Ca. Nitrosocaldus yellowstonii” is likely responsible for ammonia oxidation at the upper temperature limit within the GBS geothermal system.

Because the only known thermophilic NOB belong to the genus *Nitrospira* (Ehrich et al., 1995; Lebedeva et al., 2008; Lebedeva et al., 2011), and attempts to amplify DNA using a *Nitrobacter*-specific 16S rRNA gene primers failed, NOB were quantified using a *Nitrospira*-specific 16S rRNA gene primer set. Gene copy numbers for *Nitrospira* increased along the outflow of SSW and RHC, and were not abundant (>10^6) at ≥61.4 °C in either spring (Figure 3.1). However, NO₃⁻ concentrations in SSW in July,
2010 clearly indicate the presence of \( \text{NO}_2^- \) oxidation at \( \geq 61.4 \, ^\circ\text{C} \); therefore, the upper temperature limit for \( \text{NO}_2^- \) oxidation appears to be close to but slightly \( > 61.4 \, ^\circ\text{C} \) within the GBS geothermal system.

**Identification of NOB**

In order to evaluate the specificity of qPCR primers used and to identify the NOB present in the samples, clone libraries were constructed from the July 2010 sediment samples that contained NOB (four samples; Table 2). 36 clones were sequenced from the NOB-abundant samples, resulting in a total of 144 clones. The sequences amplified by the NOB-specific primers were closely related to *N. moscoviensis* and *N. calida*, with a minimum identity \( \geq 95\% \) (Table 3.2). These results confirmed the specificity of the primers used and suggest a thermophilic clade within the genus *Nitrospira* that includes *N. moscoviensis*, *N. calida*, and yet-uncultivated relatives.

**Cultivation-independent data**

As an independent test of the thermal niche of *Nitrospira* and other NOB in geothermal environments, pyrotag datasets from GBS and Tengchong were scanned for 16S rRNA gene sequences belonging to known NOB. Known NOB were absent from 164,178 quality-filtered pyrotag sequences amplified from GBS (twelve environmental samples taken on four different sampling dates) at temperatures ranging from 87 – 62 °C (Cole *et al.*, 2012). Furthermore, significant matches with the entire genomes of “*Candidatus Nitrospira defluvii*”, *Nitrobacter winogradski*, and *Nitrobacter hamburgensis*, were absent from 250 Mbp of metagenomic data obtained from GBS environmental samples at temperatures ranging from 77 to 85 °C. Similarly, known NOB were absent from 392,899 pyrotags from 37 samples from hot springs in Rehai and
Ruidian, ranging in temperature from 55.1 - 93.6 °C and in pH from 2.46 to 9.39 (Hou et al., 2012). These results, along with the lack of any reports of successful NOB cultures at temperatures >60 °C, suggest an upper temperature limit for chemolithotrophic nitrite oxidation of ~60 - 65 °C.

**Conclusions**

The data presented here underscores the need to consider nitrite oxidation, and ammonia oxidation separately when studying nitrification. Inferred rates of metabolic activity and abundance of “Ca. N. yellowstonii” and *Nitrospira* suggested a distinct difference in the upper temperature limits for zones of ammonia and nitrite oxidation within the GBS geothermal system. Rates of ammonia oxidation and “Ca. N. yellowstonii” abundance suggested an upper temperature limit for ammonia oxidation activity of ~80 - 85 °C. In contrast, rates of nitrite oxidation and abundance of *Nitrospira* spp. suggested an upper temperature limit of ~60 - 65 °C for nitrite oxidation. In addition, 16S rRNA sequences belonging to known NOB were absent from 557,076 pyrotag sequences obtained from hot springs located in the U.S. Great Basin, and Tengchong, China at temperatures ≥55 °C. Furthermore, significant matches with the entire genomes of known NOB were absent from metagenome data obtained from GBS environmental samples at temperatures ranging from 77 to 85 °C. Together these results provide evidence of a thermal decoupling between the two steps of nitrification resulting in an accumulation of NO₂ at temperatures ≥65 °C within the GBS geothermal system.

Along with inferred activity measurements, the distribution and abundance of “Ca. N. yellowstonii” and *Nitrospira* spp. may offer clues about other nitrogen cycle processes within the GBS system. Because of their dependence on ammonia oxidation
within the GBS system, both DNRA and denitrification are likely to be inhibited at temperatures above >84.4 °C. While the effect of the lack of nitrite oxidation at ≥65 °C is not quite as clear, the elevated NO$_2^-$ concentrations along with low NO$_3^-$ may result in a nitrite shunt at temperatures >65 °C. This high temperature nitrite shunt is complete, with no flow of nitrogen through nitrate at temperatures exceeding ~65 °C. Further investigations are needed to determine the effects of this thermal decoupling on the community structure and function with the GBS geothermal system.
Table 3.1. Locations, temperatures, and pH for each sampling site.

<table>
<thead>
<tr>
<th>Site</th>
<th>Sampling Date</th>
<th>Temperature (°C)</th>
<th>pH</th>
<th>GPS location</th>
<th>Distance From Source (m)</th>
</tr>
</thead>
<tbody>
<tr>
<td>SSW 70</td>
<td>7/11/2010</td>
<td>71.9</td>
<td>7.67</td>
<td>n.d.</td>
<td>n.d.</td>
</tr>
<tr>
<td>RHC Source</td>
<td>7/19/2010</td>
<td>92.6</td>
<td>7.25</td>
<td>n.d.</td>
<td>n.d.</td>
</tr>
<tr>
<td>RHC 80</td>
<td>7/19/2010</td>
<td>79.6</td>
<td>7.61</td>
<td>n.d.</td>
<td>17.07</td>
</tr>
<tr>
<td>RHC 70</td>
<td>7/19/2010</td>
<td>69.3</td>
<td>7.98</td>
<td>n.d.</td>
<td>39.93</td>
</tr>
<tr>
<td>RHC 60</td>
<td>7/19/2010</td>
<td>57.9</td>
<td>8.37</td>
<td>n.d.</td>
<td>103.94</td>
</tr>
<tr>
<td>RHC 50</td>
<td>7/19/2010</td>
<td>48.3</td>
<td>8.47</td>
<td>n.d.</td>
<td>222.81</td>
</tr>
<tr>
<td>SSW Source</td>
<td>9/28/2012</td>
<td>82.9</td>
<td>7.18</td>
<td>N40° 39.184', W119° 22.499'</td>
<td>-</td>
</tr>
<tr>
<td>SSW 75</td>
<td>9/28/2012</td>
<td>75.2</td>
<td>7.29</td>
<td>N40° 39.180', W119° 22.494'</td>
<td>4.55</td>
</tr>
<tr>
<td>SSW 70</td>
<td>9/28/2012</td>
<td>70.1</td>
<td>7.38</td>
<td>N40° 39.173', W119° 22.486'</td>
<td>9.34</td>
</tr>
<tr>
<td>SSW 65</td>
<td>9/28/2012</td>
<td>65.3</td>
<td>7.51</td>
<td>N40° 39.168', W119° 22.470'</td>
<td>14.45</td>
</tr>
<tr>
<td>SSW 60</td>
<td>9/28/2012</td>
<td>60.2</td>
<td>7.57</td>
<td>N40° 39.160', W119° 22.469'</td>
<td>19.65</td>
</tr>
<tr>
<td>SSW 50</td>
<td>9/28/2012</td>
<td>49.5</td>
<td>7.78</td>
<td>N40° 39.156', W119° 22.461'</td>
<td>23.92</td>
</tr>
<tr>
<td>RHC Source</td>
<td>9/27/2012</td>
<td>93.9</td>
<td>7.02</td>
<td>N40° 40.438', W119° 21.854'</td>
<td>-</td>
</tr>
<tr>
<td>RHC 85</td>
<td>9/28/2012</td>
<td>84.6</td>
<td>7.32</td>
<td>N40° 40.450', W119° 21.855'</td>
<td>10.36</td>
</tr>
<tr>
<td>RHC 80</td>
<td>9/28/2012</td>
<td>80.3</td>
<td>7.28</td>
<td>N40° 40.455', W119° 21.852'</td>
<td>21.03</td>
</tr>
<tr>
<td>RHC 75</td>
<td>9/28/2012</td>
<td>75.1</td>
<td>7.34</td>
<td>N40° 40.461', W119° 21.848'</td>
<td>33.22</td>
</tr>
<tr>
<td>RHC 70</td>
<td>9/28/2012</td>
<td>70.4</td>
<td>7.54</td>
<td>N40° 40.466', W119° 21.837'</td>
<td>51.21</td>
</tr>
<tr>
<td>RHC 65</td>
<td>9/28/2012</td>
<td>65.0</td>
<td>7.67</td>
<td>N40° 40.471', W119° 21.826'</td>
<td>68.28</td>
</tr>
<tr>
<td>RHC 60</td>
<td>9/28/2012</td>
<td>59.8</td>
<td>7.68</td>
<td>N40° 40.490', W119° 21.810'</td>
<td>111.86</td>
</tr>
<tr>
<td>RHC 55</td>
<td>9/28/2012</td>
<td>54.5</td>
<td>7.71</td>
<td>N40° 40.505', W119° 21.800'</td>
<td>145.69</td>
</tr>
<tr>
<td>RHC 50</td>
<td>9/28/2012</td>
<td>50.5</td>
<td>7.75</td>
<td>N40° 40.520', W119° 21.791'</td>
<td>173.43</td>
</tr>
</tbody>
</table>

n.d. = not determined
**Table 3.2.** OTUs from the outflows of SSW and RHC at 60°C and 50°C.

<table>
<thead>
<tr>
<th>Representative Sequence</th>
<th># in OTU</th>
<th>Cultured microorganism whose 16S rRNA gene has the highest BLASTn hit to the OTU representative sequence</th>
</tr>
</thead>
<tbody>
<tr>
<td><strong>SSW 60 °C</strong></td>
<td></td>
<td></td>
</tr>
<tr>
<td>SSW_60_23</td>
<td>18</td>
<td><em>Nitrospira moscoviensis</em> SBR2046</td>
</tr>
<tr>
<td>SSW_60_10</td>
<td>14</td>
<td><em>Nitrospira calida NS10</em></td>
</tr>
<tr>
<td>SSW_60_36</td>
<td>4</td>
<td><em>Nitrospira calida NS10</em></td>
</tr>
<tr>
<td><strong>SSW 50 °C</strong></td>
<td></td>
<td></td>
</tr>
<tr>
<td>SSW_50_13</td>
<td>14</td>
<td><em>Nitrospira moscoviensis</em> SBR2046</td>
</tr>
<tr>
<td>SSW_50_31</td>
<td>12</td>
<td><em>Nitrospira calida NS10</em></td>
</tr>
<tr>
<td>SSW_50_23</td>
<td>9</td>
<td><em>Nitrospira calida NS10</em></td>
</tr>
<tr>
<td>SSW_50_27</td>
<td>1</td>
<td><em>Nitrospira moscoviensis</em> SBR2046</td>
</tr>
<tr>
<td><strong>RHC 60 °C</strong></td>
<td></td>
<td></td>
</tr>
<tr>
<td>RHC_4_20</td>
<td>28</td>
<td><em>Nitrospira calida NS10</em></td>
</tr>
<tr>
<td>RHC_4_32</td>
<td>8</td>
<td><em>Nitrospira calida NS10</em></td>
</tr>
<tr>
<td><strong>RHC 50 °C</strong></td>
<td></td>
<td></td>
</tr>
<tr>
<td>RHC_5_30</td>
<td>19</td>
<td><em>Nitrospira calida NS10</em></td>
</tr>
<tr>
<td>RHC_5_17</td>
<td>11</td>
<td><em>Nitrospira moscoviensis</em> SBR2046</td>
</tr>
<tr>
<td>RHC_5_37</td>
<td>3</td>
<td><em>Nitrospira moscoviensis</em> SBR2046</td>
</tr>
<tr>
<td>RHC_5_24</td>
<td>3</td>
<td><em>Nitrospira calida NS10</em></td>
</tr>
<tr>
<td>RHC_4_35</td>
<td>1</td>
<td><em>Nitrospira calida NS10</em></td>
</tr>
</tbody>
</table>

*a* Number of sequences in the 16S rRNA gene library represented by the OTU.

*b* Percent identity to the OTU representative sequence.
Figure 3.1. Abundance of “Ca. N. yellowstonii” and *Nitrospira* and concentrations of dissolved inorganic nitrogen along geothermal outflow channels of SSW and RHC in July 2010. Ammonia, nitrite, and nitrate were measured colorimetrically and 16S rRNA gene copy numbers were measured using qPCR.
Figure 3.2. Concentrations of dissolved inorganic nitrogen along geothermal outflow channels of SSW and RHC in September 2012. $\Sigma NH_3$, $NO_2^-$, and $NO_3^-$ were measured colorimetrically.
Figure 3.3. Inferred rates of metabolic activities along the outflow of SSW and RHC. Rates of ammonia and nitrite oxidation were calculated using the flow rates and changes in NO$_2^-$ + NO$_3^-$ (ammonia oxidation) and NO$_3^-$ (nitrite oxidation).
CHAPTER 4

CONCLUSIONS

The results presented here demonstrate the wide geographic range of *Nitrospira calida* in terrestrial geothermal systems. Organisms with high (>95%) 16S rRNA gene identities to *Nitrospira calida* were found in both SSW and RHC. Similarly, enrichment cultures from geothermal springs located in geothermal springs in Tengchong, China, and Armenia all contained organisms that were closely related to *Nitrospira calida* (≥97% 16S rRNA identities). The physiological properties were similar for all enrichments characterized, regardless of their origin, with an optimal initial NO$_2^-$ concentration of 500 - 750 µM, and an optimal temperature for nitrite oxidation activity of 45 - 50 °C.

Data obtained from characterization of enrichment cultures, geochemical analyses, and qPCR all consistently indicated an upper temperature limit for thermophilic NOB of 60 - 65 °C. While attempts to enrich for NOB from environments with temperatures >61 °C were unsuccessful, enrichments were successfully inoculated from 5 springs located in Great Basin, U.S., Tengchong China, and Armenia at sites <58 °C. In addition, while active at 60 °C, successful enrichments never showed any activity when incubated at temperatures >60 °C. Measurements of dissolved inorganic nitrogen concentrations in two GBS springs showed NO$_2^-$ accumulation along with consistently low NO$_3^-$ (<2 µM) suggesting a lack of nitrite oxidation activity at temperatures >60 °C. Furthermore, inferred rates of *in situ* nitrite oxidation activity were not significant at temperatures >65 °C. Consistent with these results, *Nitrospira* 16S rRNA gene copy numbers were not abundant at temperatures >57.9 °C in either spring. Similarly, 16S rRNA sequences belonging to known NOB were absent from 557,076 pyrotag sequences.
obtained from hot springs located in Great Basin, US and Tengchong, China at temperatures ≥55 °C. In addition, the 250 Mbp metagenomic data from GBS environmental samples at temperatures of 77 - 85 °C, and the 577 Mbp metagenomic data from Yellowstone environmental samples at temperatures of 52.9 - 90 °C did not contain any significant genome matches to known NOB.

In contrast to results for NOB, the data presented here suggests an upper temperature limit for AOA of 75 - 84 °C. Inferred rates of ammonia oxidation activity were significant at temperatures <84.6 °C and “Ca. N. yellowstonii” 16S rRNA genes were abundant at ≤84.4 °C. Since NH₃ is in its most reduced oxidation state, ammonia oxidation is required to provide substrate for reductive N processes. Therefore, these results indicate that the dissimilatory N cycling processes would not function within the GBS geothermal system at temperatures >82 °C (Figure 4.1). In addition, these results suggest a spatial and thermal (~15 - 20 °C) difference between the two steps in nitrification within the GBS geothermal system. While the effects of this thermal decoupling on other N cycling processes within this system have yet to be explored, it is possible that a high temperature nitrite-shunt (as shown in Figure 4.1) exists within these geothermal environments. Further research is needed to determine whether this decoupling is ubiquitous in high temperature environments. In addition, the greater implications of this decoupling of nitrification in terms of community structure and function needs to be addressed.
Figure 4.1. Conceptual model of nitrogen cycle activities in GBS system springs at temperatures: A) <65 °C B) ~65 – 82 °C C) >~82 °C.
### APPENDIX A: ADDITIONAL TABLES

**Table A.1.** Concentrations of major anions and O2 along SSW and RHC outflow systems in July, 2010.

<table>
<thead>
<tr>
<th>Site</th>
<th>Sampling Date</th>
<th>Temperature (°C)</th>
<th>Fluoride (µM)</th>
<th>Chloride (mM)</th>
<th>Bromide (µM)</th>
<th>Sulfate (mM)</th>
<th>O₂ (ppm)</th>
</tr>
</thead>
<tbody>
<tr>
<td>SSW Source</td>
<td>7/11/2010</td>
<td>84.4</td>
<td>350.00</td>
<td>61.6</td>
<td>65.06</td>
<td>3.94</td>
<td>0.40</td>
</tr>
<tr>
<td>SSW 70</td>
<td>7/11/2010</td>
<td>71.9</td>
<td>104.95</td>
<td>63.3</td>
<td>60.79</td>
<td>4.1</td>
<td>1.80</td>
</tr>
<tr>
<td>SSW 60</td>
<td>7/11/2010</td>
<td>61.4</td>
<td>b.d.</td>
<td>63.8</td>
<td>57.07</td>
<td>4.1</td>
<td>3.00</td>
</tr>
<tr>
<td>SSW 50</td>
<td>7/11/2010</td>
<td>49.8</td>
<td>451.04</td>
<td>65.4</td>
<td>51.22</td>
<td>4.2</td>
<td>5.20</td>
</tr>
<tr>
<td>RHC Source</td>
<td>7/19/2010</td>
<td>92.6</td>
<td>b.d.</td>
<td>61.52</td>
<td>82.31</td>
<td>3.8</td>
<td>0.00</td>
</tr>
<tr>
<td>RHC 80</td>
<td>7/19/2010</td>
<td>79.6</td>
<td>151.81</td>
<td>63.2</td>
<td>69.36</td>
<td>4.0</td>
<td>1.20</td>
</tr>
<tr>
<td>RHC 70</td>
<td>7/19/2010</td>
<td>69.3</td>
<td>b.d.</td>
<td>65.1</td>
<td>68.79</td>
<td>4.1</td>
<td>1.80</td>
</tr>
<tr>
<td>RHC 60</td>
<td>7/19/2010</td>
<td>57.9</td>
<td>319.52</td>
<td>65.2</td>
<td>64.15</td>
<td>4.1</td>
<td>2.22</td>
</tr>
<tr>
<td>RHC 50</td>
<td>7/19/2010</td>
<td>48.3</td>
<td>18.19</td>
<td>66.3</td>
<td>58.44</td>
<td>4.2</td>
<td>2.80</td>
</tr>
</tbody>
</table>

b.d. is below detection
Figure B.1. Lag of NO$_2^-$ oxidation activity with respect to temperature in enrichment cultures from Great Basin, Tengchong, and Armenia. NO$_2^-$ was monitored colorimetrically. Errors bars, standard deviation (n=3). * Samples incubated at room temperature (24 °C to 26 °C).
Figure B.2. Lag of NO$_2^-$ oxidation activity with respect to initial NO$_2^-$ concentration in enrichment cultures from Great Basin, Tengchong, and Armenia. NO$_2^-$ was monitored colorimetrically. Errors bars, standard deviation (n=3).
REFERENCES


Felsenstein, J. 2005. PHYLIP (Phylogeny Inference Package) version 3.6 Distributed by the author, Department of Genetics, University of Washington, Seattle.


**Zhang G, C. Liu, H. Liu, Z. Jin, G. Han, and L. Li.** 2008b. Geochemistry of the Rehai and Ruidian geothermal waters, Yunnan Province, China. Geothermics **37:**73-83

VITA

Graduate College
University of Nevada, Las Vegas

Tara A. Edwards

Degree:
Associate of Science, 2006
College of Southern Nevada

Bachelor of Science, Biology, 2010
University of Nevada – Las Vegas

Bachelor of Science, Biochemistry 2010
University of Nevada, Las Vegas

Thesis Title: Nitrite oxidation in geothermal springs: evidence of an upper temperature limit for thermophilic nitrite-oxidizing bacteria of ~60 - 65 °C

Thesis Examination Committee:
   Committee Chair, Brian Hedlund, Ph.D.
   Committee Member, Eduardo Robleto, Ph.D.
   Committee Member, Penny Amy, Ph.D.
   Graduate Faculty Representative, Elisabeth Hausrath, PhD.